



# CESM2 release of CAM-SE (& CAM-SE-CSLAM)

Peter Hjort Lauritzen National Center for Atmospheric Research Boulder, Colorado, USA

> Collaborators: S. Goldhaber (NCAR), J. Bacmeister (NCAR), R.D.Nair (NCAR), M.A. Taylor (Sandia)

CESM Atmosphere Model Working Group Meeting 27 February – 1 March 2017 Boulder, Colorado, USA

 NCAR
 National Center for Atmospheric Research

 UCAR
 Climate & Global Dynamics
 climate 

 Models
 Society

# **New SE dynamical core for CESM2**

#### Science changes:

- Dry mass vertical coordinates
- Condensate loading in dynamical core (recommended as default)
- Separate physics grid and CSLAM options (not scientifically supported "yet")
- Eulerian vertical advection no longer supported! Moist vertical coordinates not supported! (to keep code base simpler)

#### **Other:**

- Out-of-the-box CESM configurations for idealized setups (Held-Suarez, moist baroclinic wave with Kessler physics, terminator chemistry, ...)
- Performance upgrades from CISL (threading, more efficient SE transport)
- Dynamical core is no longer imported from HOMME (High-Order Methods Modeling Environment) ; code must go through CAM code review
- Cleaned up code base: trunk SE has ~61000 lines of code; new SE has ~39000 lines of code (further cleanup in progress ...)

Consider a 'moist'  $\eta\text{-coordinate system:}$  The pressure is given by

$$p(\eta) = A(\eta)p_0 + B(\eta)ps,$$

where ps is 'moist' surface pressure.

In a floating  $\eta$ -coordinate system,  $\dot{\eta} = 0$ , the continuity equation for p can be written as

$$\frac{\partial}{\partial t} \left[ \left( \frac{\partial p}{\partial \eta} \right) \right] + \nabla \cdot \left[ \left( \frac{\partial p}{\partial \eta} \right) \vec{v} \right] = S^p,$$

where  $S^p(q_v)$  is the source/sink term for pressure  $(q_v \equiv \text{specific humidity}).$ 

- This source/sink term:
  - makes the handling of tracers more complicated

An inert tracer will have source/sink terms (i.e. if there are moisture changes all "wet" mixing ratios must be changed accordingly)

- makes it harder to move towards conserving a more comprehensive total energy
- makes it harder to represent condensate loading in the dynamical core
- Complicates CSLAM-SE coupling in a moist atmosphere

$$\frac{\partial}{\partial t} \left[ \left( \frac{\partial p}{\partial \eta} \right) \right] + \nabla \cdot \left[ \left( \frac{\partial p}{\partial \eta} \right) \vec{v} \right] = S^p,$$

where  $S^p(q_v)$  is the source/sink term for pressure  $(q_v \equiv \text{specific humidity}).$ 

he

If one uses a dry mass vertical coordinate

$$p(\eta_d) = A(\eta_d)p_0 + B(\eta_d)ps_d,$$

where  $ps_d$  is dry surface pressure, then the continuity equation for pressure does not have sources/sinks

$$\frac{\partial}{\partial t} \left[ \left( \frac{\partial p_d}{\partial \eta_d} \right) \right] + \nabla \cdot \left[ \left( \frac{\partial p_d}{\partial \eta_d} \right) \vec{v} \right] = 0.$$

Model levels do not move during physics-dynamics coupling!

The  $\eta_d$ -coordinate atmospheric primitive equations assuming floating Lagrangian vertical coordinates [*Starr*, 1945; *Lin*, 2004] can be written as

$$\frac{\partial \vec{v}}{\partial t} + (\zeta + f) \,\hat{\vec{k}} \times \vec{v} + \nabla_{\eta_d} \left( \frac{1}{2} \vec{v}^2 + \Phi \right) + \frac{1}{\rho} \nabla_{\eta_d} p = \nu \nabla^4 \vec{u},\tag{16}$$

$$\frac{\partial T}{\partial t} + \vec{v} \cdot \nabla_{\eta_d} T - \frac{1}{c_p \rho} \omega = \nu \nabla_{\eta_d}^4 T, \qquad (17)$$

$$\frac{\partial}{\partial t} \left( \frac{\partial p_d}{\partial \eta_d} m_i \right) + \nabla_{\eta_d} \cdot \left( \frac{\partial p_d}{\partial \eta_d} m_i \vec{v} \right) = v \nabla_{\eta_d}^4 (m_i), \quad i = d, v, cl, ci, \dots$$
(18)

where  $\Phi$  is the geopotential height ( $\Phi = g z$ , where g is the gravitational constant),  $c_p$  is the specific heat constant for dry air,  $\hat{\vec{k}}$  is the unit vector normal to the surface of the sphere,  $\zeta = \hat{\vec{k}} \cdot \nabla \times \vec{v}$  is vorticity, f Coriolis parameter, and  $\omega = Dp/Dt$  is the pressure vertical velocity.

The  $\eta_d$ -coordinate atmospheric primitive equations assuming floating Lagrangian vertical coordinates [*Starr*, 1945; *Lin*, 2004] can be written as

$$\frac{\partial \vec{v}}{\partial t} + (\zeta + f) \,\hat{\vec{k}} \times \vec{v} + \nabla_{\eta_d} \left( \frac{1}{2} \vec{v}^2 + \Phi \right) + \frac{1}{\rho} \nabla_{\eta_d} p = \nu \nabla^4 \vec{u},\tag{16}$$

$$\frac{\partial T}{\partial t} + \vec{v} \cdot \nabla_{\eta_d} T - \frac{1}{c_p \rho} \omega = v \nabla_{\eta_d}^4 T, \qquad (17)$$

$$\frac{\partial}{\partial t} \left( \frac{\partial p_d}{\partial \eta_d} m_i \right) + \nabla_r \left( \frac{\partial p_d}{\partial \eta_d} m_i \vec{v} \right) = v \nabla_{\eta_d}^4 (m_i), \quad i = d, v, cl, ci, \dots$$
(18)

where  $\Phi$  is the geopotential height ( $\vec{\sigma}$  g z, where  $\xi$  is the gravitational constant),  $c_p$ is the specific heat constant for dr is the unit vertical velocity. where  $\xi = \hat{\vec{k}} \cdot \nabla \times \vec{v}$  is vorticit is the unit vertical velocity.

$$\rho = \rho_d\left(\sum_i m_i\right), i = `d`, `v`, `cl`, `ci`.$$

$$c_p = \frac{\sum_i \left[ m_i \, c_{pi} \right]}{\sum_i m_i}$$

dry air 'd', water vapor 'v', cloud liquid 'cl' and cloud ice 'ci'

# "Correct" Internal Energy

The total internal energy integrated over the entire atmosphere is given by

$$I_{tot} = \iiint \rho c_p T \, dz \, \cos(\theta) r \, d\lambda \, d\theta$$

Using the hydrostatic balance this equation can be written as

$$I_{tot} = \sum_{i} I_i,$$

where  $I_d$  is the total internal energy of dry air,  $I_v$  the total internal energy of water vapor, etc.:

$$I_{d} = -\frac{1}{g} \iiint c_{pd} T\left(\frac{\partial p_{d}}{\partial \eta_{d}}\right) d\eta_{d} \cos(\theta) r \, d\lambda \, d\theta,$$
  

$$I_{v} = -\frac{1}{g} \iiint c_{pv} m_{v} T\left(\frac{\partial p_{d}}{\partial \eta_{d}}\right) d\eta_{d} \cos(\theta) r \, d\lambda \, d\theta,$$
  

$$I_{cl} = -\frac{1}{g} \iiint c_{pcl} m_{cl} T\left(\frac{\partial p_{d}}{\partial \eta_{d}}\right) d\eta_{d} \cos(\theta) r \, d\lambda \, d\theta,$$
  

$$I_{ci} = -\frac{1}{g} \iiint c_{pci} m_{ci} T\left(\frac{\partial p_{d}}{\partial \eta_{d}}\right) d\eta_{d} \cos(\theta) r \, d\lambda \, d\theta.$$

# "Correct" Internal Energy

The total internal energy integrated over the entire atmosphere is given by

$$I_{tot} = \iiint \rho c_p T \, dz \, \cos(\theta) r \, d\lambda \, d\theta$$

Using the hydrostatic balance

equation can be written as

$$I_{tot} = \sum_i I_i,$$

where  $I_d$  is the total interest.

Iry air,  $I_{\nu}$  the total internal energy of water vapor,

The internal energy in CAM physics is defined as

$$I_{tot}^{(CAM)} = -\frac{1}{g} \iiint c_{pd} T (1 + m_{v}) \left(\frac{\partial p_{d}}{\partial \eta_{d}}\right) d\eta_{d} \cos(\theta) r \, d\lambda \, d\theta$$

$$I_{cl} = -\frac{1}{g} \iiint c_{pcl} m_{cl} T \left(\frac{\partial p_{d}}{\partial \eta_{d}}\right) d\eta_{d} \cos(\theta) r \, d\lambda \, d\theta,$$

$$I_{ci} = -\frac{1}{g} \iiint c_{pci} m_{ci} T \left(\frac{\partial p_{d}}{\partial \eta_{d}}\right) d\eta_{d} \cos(\theta) r \, d\lambda \, d\theta.$$

# "Correct" Internal Energy

The total internal energy integrated over the entire atmosphere is given by

$$I_{tot} = \iiint \rho c_p T \, dz \, \cos(\theta) r \, d\lambda \, d\theta$$

**Enforcing the correct energy in CAM physics is NON-TRIVIAL:** 

- If a parameterization alters water vapor, cloud liquid, and/or cloud ice then internal energy (and kinetic energy) changes
- The assumption that pressure levels stay fixed during physics updates is violated unless we switch to dry pressure levels

of water vapor,

$$I_{\nu} = -\frac{1}{g} \iiint c_{p\nu} m_{\nu} T\left(\frac{\partial p_{d}}{\partial \eta_{d}}\right) d\eta_{d} \cos(\theta) r \, d\lambda \, d\theta,$$
  

$$I_{cl} = -\frac{1}{g} \iiint c_{pcl} m_{cl} T\left(\frac{\partial p_{d}}{\partial \eta_{d}}\right) d\eta_{d} \cos(\theta) r \, d\lambda \, d\theta,$$
  

$$I_{ci} = -\frac{1}{g} \iiint c_{pci} m_{ci} T\left(\frac{\partial p_{d}}{\partial \eta_{d}}\right) d\eta_{d} \cos(\theta) r \, d\lambda \, d\theta.$$

# **Two SE configurations**

controlled with namelist: qsize\_condensate\_loading = 1,3

- 1. Set  $c_p = c_{pd}$  and  $\rho = \rho_d + \rho_v$
- 2. Use 'correct'  $c_p$  and  $\rho = \rho_d + \rho_v + \rho_{ci} + \rho_{cl}$

In dynamics-physics coupling we pass  $\Delta p = \Delta p_d(1 + m_v)$  to remain consistent with the physics definition of total energy.

Both configurations have pros and cons:

- 1a. Only difference between #1 and trunk is vertical coordinate
- 1b. The continuous equations of motion conserve the same "wrong" energy as CAM physics.
- 1c. The adiabatic momentum equations and thermodynamic equations do not "feel" the condensates
- 2a. The continuous equations of motion conserve the "correct" energy but CAM physics will (through the energy fixer) enforce the "wrong" energy
- 2b. The adiabatic momentum equations and thermodynamic equations "feel" the condensates (a.k.a. condensate loading; may be significant at higher resolution)

# **Energy budgets in CAM-FV (0.9x0.9)**

CAM5.3 physics; 10 year run

pEFIX = dE/dt energy fixer pDMEA = dE/dt dme\_adjust  $\begin{array}{rll} (pBP\text{-}pBF) & : & -0.8103 \ \text{W/m}^2 \\ (pAM\text{-}pAP) & : & 0.2636 \ \text{W/m}^2 \end{array}$ 

Energy fixer fixes dme\_adjust (pDMEA), lack of energy conservation in adiabatic dynamical core (dADIA) and energy lost/gained in physics-dynamics coupling (dPDC):

-pEFIX = pDMEA+dADIA+dPDC

CAM-FV uses updated state (no "drippling" of tendencies) from physics so dPDC=0

 $\Rightarrow$  dADIA = dE/dt adiabatic dynamical core = -pEFIX-pDMEA : -1.0739 W/m<sup>2</sup>

Aside:

At 2 degree horizontal resolution dE/dt adiabatic dynamical : -1.2738 W/m<sup>2</sup>

#### Energy budgets in CAM-SE configuration 1 CAM5.9999 physics; 6 year run; qsize condensate loading = 1 http://webext.cgd.ucar.edu/FCLIMO/f.e20.F2000 DEV.ne30 ne30.physgrid25 cam5 4 96 se gsize 1/atm/

pEFIX = dE/dt energy fixer pDMEA = dE/dt dme\_adjust (pBP-pBF) : -0.1913 W/m<sup>2</sup> (pAM-pAP) : 0.3064 W/m<sup>2</sup>

Energy fixer fixes dme\_adjust (pDMEA), lack of energy conservation in adiabatic dynamical core (dADIA) and energy lost/gained in physics-dynamics coupling (dPDC):

-pEFIX = pDMEA+dADIA+dPDC

dADIA = dE/dt adiabatic dynamical core = -0.0732 W/m<sup>2</sup> (-0.1604 W/m<sup>2</sup> vertical remapping, 0.0872 W/m<sup>2</sup> Lagrangian dyn, hypervis V added as heating = 0.7110 W/m<sup>2</sup>)

dPDC = dE/dt physics-dynamics coupling (ftype=0) = -0.0419 W/m<sup>2</sup>

dADIA (SE configuration 1) = -0.0732 W/m<sup>2</sup> << dADIA (FV) = -1.0739 W/m<sup>2</sup> pDMEA is about the same for CAM-FV and CAM-SE

#### Energy budgets in CAM-SE configuration 2 CAM5.9999 physics; 6 year run; gsize condensate loading = 3

CAM5.9999 pnysics; 6 year run; qsize\_condensate\_loading = 3 http://webext.cgd.ucar.edu/FCLIMO/f.e20.F2000\_DEV.ne30\_ne30.physgrid25\_cam5\_4\_96\_se\_gsize\_3/atm/

pEFIX = dE/dt energy fixer pDMEA = dE/dt dme\_adjust (pBP-pBF) : -0.7070 W/m<sup>2</sup> (pAM-pAP) : 0.3102 W/m<sup>2</sup>

Energy computations in dynamical use "correct" energy formula Energy computations in physics use "wrong" energy formula

=> We can not mix computations done in physics and dynamics but ....

 $\begin{array}{ll} \text{dADIA} = \text{dE}/\text{dt} \text{ adiabatic dynamical core} & = -0.0928 \ \text{W}/\text{m}^2 \\ \text{pEFIX configuration 2} - \text{pEFIX configuration 1} & = 0.5257 \ \text{W}/\text{m}^2 \\ \text{i.e. the inconsistency in energy formula is ca. } 0.5 \ \text{W}/\text{m}^2 \ \text{consistent with} \\ \text{Mark Taylors findings.} \end{array}$ 

Even with this inconsistency the energy fixer fixes less than for CAM-FV => I recommend configuration 2 as default in CESM2





## **CAM-SE-CSLAM without moisture**

**MARCH 2017** 

LAURITZEN ET AL.

#### Monthly Weather Review

#### CAM-SE-CSLAM: Consistent Coupling of a Conservative Semi-Lagrangian Finite-Volume Method with Spectral Element Dynamics

PETER HJORT LAURITZEN

National Center for Atmospheric Research,<sup>a</sup> Boulder, Colorado

MARK A. TAYLOR AND JAMES OVERFELT

Sandia National Laboratories, Albuquerque, New Mexico

PAUL A. ULLRICH

Department of Land, Air, and Water Resources, University of California, Davis, Davis, California

RAMACHANDRAN D. NAIR, STEVE GOLDHABER, AND RORY KELLY

National Center for Atmospheric Research, a Boulder, Colorado





833

### **CAM-SE-CSLAM without moisture**







## **CAM-SE-CSLAM with moisture**

#### "This is where the fun begins!" – Staniforth et al. (2006)







## Mapping u,v, T, omega from dynamics grids (GLL) to finite-volume (CSLAM) grid

Temperature (& omega): Integrate basis function representation of dp\*T over physics grid control volumes (high-order remapping; conserves internal energy)







## Mapping u,v, T, omega from dynamics grids (GLL) to finite-volume (CSLAM) grid



# **CAM-SE with "rougher" topography**

Held-Suarez forcing with real-world topography (6 months spin-up; 2 years and 9 months average) Note: dry test so no moist physics feedback

bnd\_topo = '/ home/pel/run\_scripts/topo/ne30np4\_nc3000\_Nsw042\_Nrs008\_Co060\_Fi001\_ZR\_test\_vX\_111416.nc



UCAR | Climate & Global Dynamics climate • models • society

## **CAM-SE-CSLAM**

Held-Suarez forcing with real-world topography (6 months spin-up; 2 years and 9 months average) Note: dry test so no moist physics feedback

bnd\_topo = '/ home/pel/run\_scripts/topo/ne30np4\_nc3000\_Nsw042\_Nrs008\_Co060\_Fi001\_ZR\_test\_vX\_111416.nc



NCAR National Center for Atmospheric Research UCAR Climate & Global Dynamics













# **Temperature tendency: FT**

-4e-	-05 -2e-05	<u>0 2e-0</u>	05 4e-05 6e-0	5 8e-05	0.0001
(7) 1d va	urs (13)2d v	vars (26) 3d	vars		
Dim:	Name:	Min:	Current:	Max:	Units:
Scan:	time	76.1	1-Apr-1989 04	76.1	days since 19
	lev	3.64347	912.645	992.556	hPa
			∑ int	erp_dir/cslar	n-cam5-ape.ave.no
		بالمحمور	Dera-		•
12					
الحملكر	<b>₽</b> ₽``	-		5	-8-
-15a.	CBS	3	- 380	<u>بالر</u>	States and a state
	2 2		3	Dente .	
	18 1	me the	1 1	104	State of the second
	- 24		$\gamma \gamma$		
200	the /	Sec.		6. A.	-
			And and the second second	1	
	18		5		$\langle \cdot \rangle$
Ľ	1			$\sim$	
		And a state			
Sec. 1					

#### CAM-SE-CSLAM with linear interpolation from phys to dyn: 5 month average

NCAR National Center for Atmospheric Research Climate & Global Dynamics Climate • models • society

# PRECT



CAM4 SE-CSLAM-physgrid: linear interpolation phys to dyn: 5 month average Plot looks similar for standard SE (but less noisy)

> NCAR National Center for Atmospheric Research UCAR Climate & Global Dynamics climate • models • society

# **Temperature tendency: FT**



#### CAM-SE-CSLAM with cubic tensor product interpolation from phys to dyn: 18 month average

NCAR National Center for Atmospheric Research UCAR Climate & Global Dynamics climate • models • society

# PRECT



CAM-SE-CSLAM with cubic tensor product interpolation from phys to dyn: 18 month average

> NCAR National Center for Atmospheric Research UCAR Climate & Global Dynamics climate • models • society























# CAM-S' SLAM c ;uration

We are getting very close to finally do science (beyond numerical methods research) with CAM-SE-CSLAM









