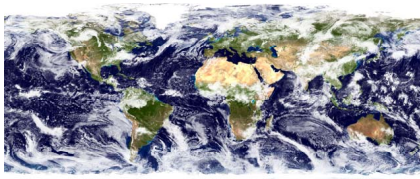
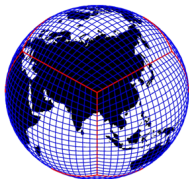


Conservative Semi-Lagrangian Multi-tracer scheme (CSLAM): a semi-implicit shallow water and a fully compressible non-hydrostatic solver with fully consistent transport

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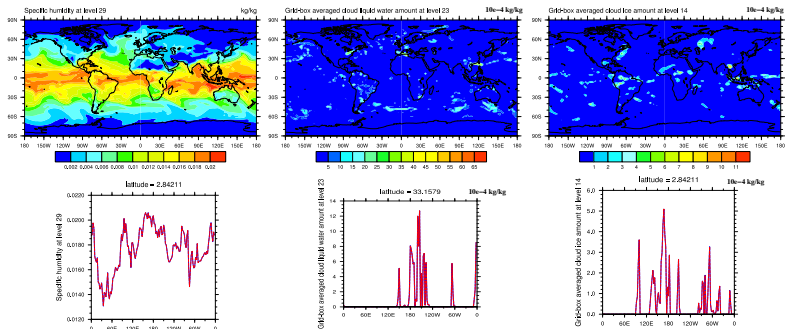
- 1 CSLAM transport scheme
 - formulation
 - air-tracer mass coupling
 - conditions for local mass-conservation
 - extension to there sphere: gnomonic coordinates
 - CSLAM in NCAR's CAM-SE (Community Atmosphere Model - Spectral Element)
- 2 CSLAM-SW: shallow water model
 - semi-implicit time-stepping of tracers
 - treatment of divergence
- 3 CSLAM-NH: non-hydrostatic $x - z$ plane solver
 - preliminary results

- Multi-tracer efficiency
 - CAM5 has 31 continuity equations
(micro-physics and convection scheme developers are eager to add more!)
 - CAM-Chem has approximately 107 continuity equations
- Shape-preservation (large gradients, physics)
- Consistency (air ↔ tracer)
- Correlation accuracy
- Efficiency on 'traditional'¹ massively parallel machines:
 - *minimize frequency of message passing*: e.g. long Δt 's (semi-Lagrangian)
 - *minimize message sizes*: local computational stencil
 - *minimize memory usage*: e.g., 2-time-level, no multi-moment
- Accuracy on non-orthogonal grids (splitting errors) ⇒ fully two-dimensional methods

¹GPUs and Intel MIC architectures??

Design (see fuller discussion in Lauritzen et al., 2011)

- Multi-tracer efficiency
- Shape-preservation (large gradients, physics)



- Consistency (air \leftrightarrow tracer)
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- Multi-tracer efficiency
- Shape-preservation (large gradients, physics)
- Consistency (air \leftrightarrow tracer)

Consider flux-form continuity equations for air mass and tracer mass:

$$\frac{\partial \rho}{\partial t} + \nabla \cdot (\rho \vec{v}) = 0, \quad (1)$$

$$\frac{\partial (\rho q)}{\partial t} + \nabla \cdot (\rho q \vec{v}) = 0, \quad (2)$$

where ρ is air density and q is tracer mixing ratio. 'Free-stream' preservation implies that the discretization scheme for (2) reduces to (1) when $q = 1$.

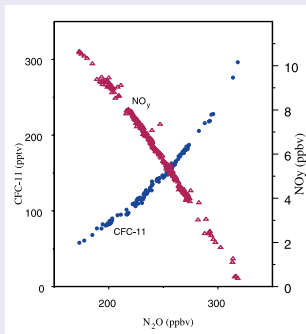
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Relationships between long-lived stratospheric tracers, manifested in similar spatial structures on scales ranging from a few to several thousand kilometers, are displayed most strikingly if the mixing ratio of one is plotted against another, when the data collapse onto remarkably compact curves. - Plumb (2007)

E.g., when plotting nitrous oxide (N_2O) against 'total odd nitrogen' (NO_y) or chlorofluorocarbon (CFC's)



It is therefore highly desirable that the transport schemes used in chemistry and chemistry-climate models should not disrupt such functional relations in unphysical ways through numerical mixing or, indeed, unmixing.' -(Lauritzen and Thuburn, 2012)

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NCAR's Yellowstone is a 1.5-petaflops high-performance computing system with 72,288 processor cores.

- Accuracy on non-orthogonal grids (splitting errors) \Rightarrow fully two-dimensional methods

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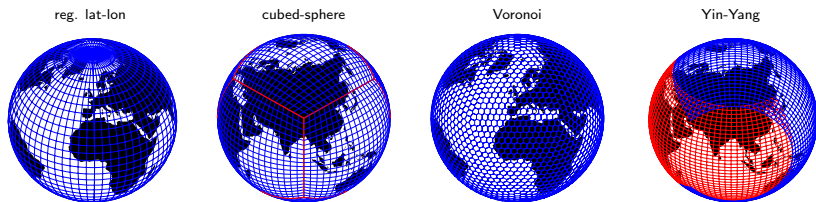


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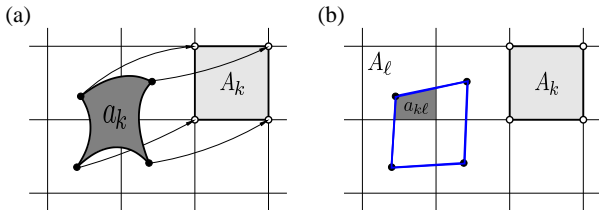
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Part I
New geometrically flexible multi-tracer scheme

CSLAM is based on pioneering work by Dukowicz (1984), Ramshaw (1985), Dukowicz and Baumgardner (2000), and Margolin and Shashkov (2003)!



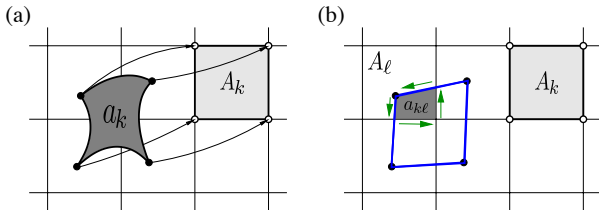
Finite-volume Lagrangian form of continuity equation for $\psi = \rho, \rho \phi$:

$$\int_{A_k} \psi_k^{n+1} dx dy = \int_{a_k} \psi_k^n dx dy = \sum_{\ell=1}^{L_k} \iint_{a_{k\ell}} f_{\ell}(x, y) dx dy,$$

where the $a_{k\ell}$'s are non-empty overlap regions:

$$a_{k\ell} = a_k \cap A_{\ell}, \quad a_{k\ell} \neq \emptyset; \quad \ell = 1, \dots, L_k. \quad (1)$$

For higher-order upstream cell-edge approximations see Ullrich et al. (2013).

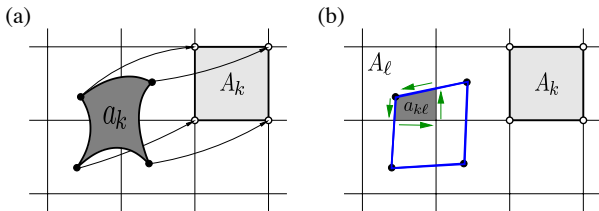


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where $\partial a_{k\ell}$ is the boundary of $a_{k\ell}$ and

$$f_\ell(x, y) = \sum_{i+j \leq 2} c_\ell^{(i,j)} x^i y^j.$$

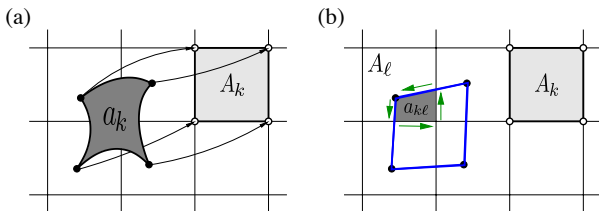


Finite-volume Lagrangian form of continuity equation for $\psi = \rho, \rho \phi$:

$$\int_{A_k} \psi_k^{n+1} dx dy = \int_{a_k} \psi_k^n dx dy = \sum_{\ell=1}^{L_k} \oint_{\partial a_{k\ell}} [P dx + Q dy],$$

where $\partial a_{k\ell}$ is the boundary of $a_{k\ell}$ and

$$\sum_{i+j \leq 2} \left[-\frac{\partial P^{(i,j)}}{\partial y} + \frac{\partial Q^{(i,j)}}{\partial x} \right] = f_\ell(x, y) = \sum_{i+j \leq 2} c_\ell^{(i,j)} x^i y^j.$$



Finite-volume Lagrangian form of continuity equation for $\psi = \rho, \rho \phi$:

$$\int_{A_k} \psi_k^{n+1} dx dy = \int_{a_k} \psi_k^n dx dy = \sum_{\ell=1}^{L_k} \left[\sum_{i+j \leq 2} c_\ell^{(i,j)} w_{k\ell}^{(i,j)} \right],$$

where weights $w_{k\ell}^{(i,j)}$ are functions of the coordinates of the vertices of $a_{k\ell}$.

$w_{k\ell}^{(i,j)}$ can be re-used for each additional tracer \Rightarrow *multi-tracer efficiency!*

Air density ρ and tracer mixing ratio q must be coupled carefully to ensure:

- mass-conservation
- shape-preservation (q is invariant following parcels; not ρq)
- 'free-stream' preservation

A 'Lagrangian' solution (Appendix B of Nair and Lauritzen, 2010)

- In cell k reconstruct sub-grid-scale distribution for ρ and q separately:

$$\rho(x, y) = \sum_{i+j \leq 2} \rho^{(i,j)} x^i y^j \quad \text{and} \quad q(x, y) = \sum_{i+j \leq 2} q^{(i,j)} x^i y^j.$$

- Apply shape-preserving reconstruction filter to $q(x, y)$ (see next slide)
- In Eulerian cell k tracer mass sub-grid-scale reconstruction is:

$$\rho q(x, y) = \bar{\rho}_k q_k(x, y) + \bar{q}_k [\rho_k(x, y) - \bar{\rho}_k], \quad (1)$$

where $\bar{(\cdot)}$ is cell average value; $\bar{q}_k = \overline{\rho q_k} / \bar{\rho}_k$

$\Rightarrow q = 1 \Rightarrow$ reconstruction (1) reduces to reconstruction of air density!

$\Rightarrow \rho q(x, y)$ is degree 2 (with $\rho q(x, y) = \rho(x, y) \times q(x, y)$ it would have been 4)

Fully 2D reconstruction filter/limiter (Barth and Jespersen, 1989)

Scale the reconstruction function $\psi(x, y)$ so that extreme values lie within the adjacent cell-average values (can be applied selectively for less diffusion, Harris et al., 2010)

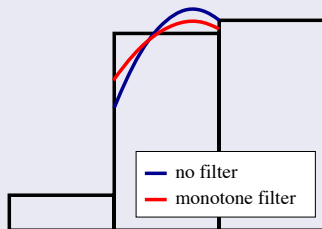


Figure: One-dimensional illustration of fully two-dimensional filter

Note that enforcing shape-preservation is 'harder/stricter' than for flux-form schemes:

- For Lagrangian schemes we can't mix low and high-order fluxes (e.g. Zalesak, 1979)
- Reconstruction functions must satisfy mass-conservation constraint:

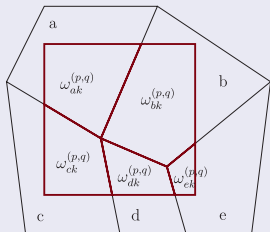
$$\int_{A_k} \psi_k(x, y) dA = \bar{\psi}_k \Delta A_k,$$

where $\bar{\psi}_k$ is cell average value over A_k with area ΔA_k . (more on this in a moment)

Conditions for local mass-conservation (Erath et al., 2013)

- Line-integrals must span the domain without 'cracks/overlaps':

$$\sum_{i \in \mathcal{E}} \Delta a_{ik} = \Delta A_k \quad \text{where} \quad \mathcal{E}_k = \{\ell \mid a_{\ell k} \cap A_k \neq \emptyset\} = \{(a, b, c, d, e)\} \quad (2)$$



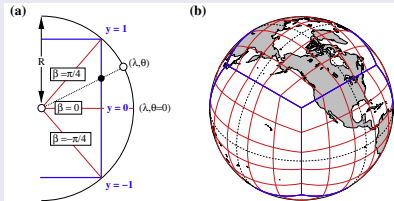
- 'Interior' line-integrals cancel.
- Line-integrals along boundary of Eulerian cell do not cancel since the reconstruction function is not continuous across cell boundaries.

- Boundary line-integrals must integrate $f_\ell(x, y)$ exactly!:**

$$\sum_{\ell \in \mathcal{E}_k} \left[\sum_{i+j \leq 2} c_\ell^{(i,j)} x^i y^j \right] = \bar{\psi}_k \Delta A_k, \quad (3)$$

Satisfying (3) on sphere can be tricky - next slide!

Line-integrals on the sphere: gnomonic projection



'Cartesian-like' coordinates:

$$(x, y) = R (\tan \alpha, \tan \beta) \quad (4)$$

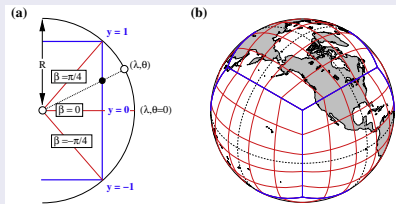
where R radius and $\alpha, \beta = [-\frac{\pi}{4}, \frac{\pi}{4}]$ central angles.

Line-integrals for β constant can be computed exactly (Ullrich et al., 2009):
(note: mass-conservation relies on this!)

$$\begin{aligned} I^{(0,0)} &= -\arctan\left(\frac{xy}{\rho}\right), & I^{(2,0)} &= -y \operatorname{arcsinh}\left(\frac{x}{\sqrt{1+y^2}}\right) - \arccos\left(\frac{x}{\sqrt{1+x^2}} \frac{y}{\sqrt{1+y^2}}\right), \\ I^{(1,0)} &= \operatorname{arcsinh}\left(\frac{y}{\sqrt{1+x^2}}\right), & I^{(0,2)} &= -x \operatorname{arcsinh}\left(\frac{y}{\sqrt{1+x^2}}\right) - \arccos\left(\frac{x}{\sqrt{1+x^2}} \frac{y}{\sqrt{1+y^2}}\right), \\ I^{(0,1)} &= \operatorname{arcsinh}\left(\frac{x}{\sqrt{1+y^2}}\right), & I^{(1,1)} &= \rho, \end{aligned}$$

→ some integrals 'ill-conditioned', in particular, at high resolution!

Line-integrals on the sphere: gnomonic projection

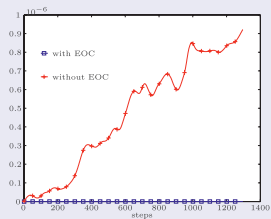


'Cartesian-like' coordinates:

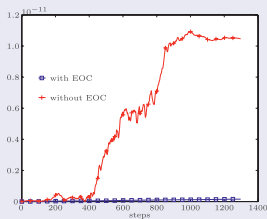
$$(x, y) = R (\tan \alpha, \tan \beta) \quad (4)$$

where R radius and $\alpha, \beta = [-\frac{\pi}{4}, \frac{\pi}{4}]$ central angles.

Performing line-integrals along cell boundaries with Gaussian quadrature is much more robust, however, integrals are not exact!



(a) Two Gaussian points for line integrals.



(b) Four Gaussian points for line integrals.

Enforce consistency: Locally scale weights so that

$$\sum_{i \in \mathcal{E}} a_{ik} = A_k \quad (5)$$

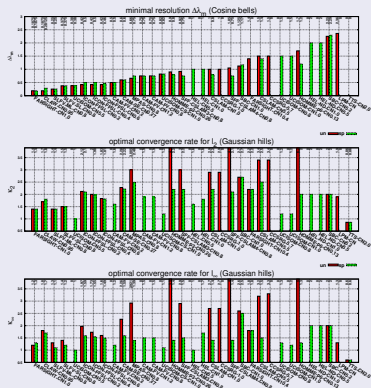
and similarly for higher-order moments.

(Erath et al., 2013)

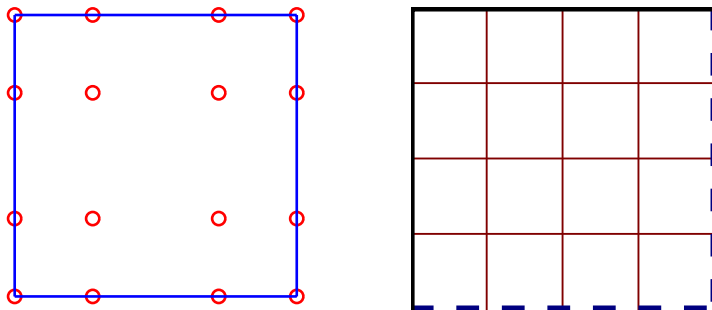
New diagnostics/test case suite designed to assess:

- 1 numerical order of convergence,
- 2 'minimal' resolution,
- 3 ability of the transport scheme to preserve filaments,
- 4 ability of the transport scheme to transport 'rough' distributions,
- 5 ability of the transport scheme to preserve pre-existing functional relations between tracers,
- 6 ability of transport scheme to deal with divergent flows (Nair and Lauritzen, 2010).

Manuscript comparing 17 state-of-the-art schemes using new standard test case suite is almost complete (Lauritzen et al., 2013)



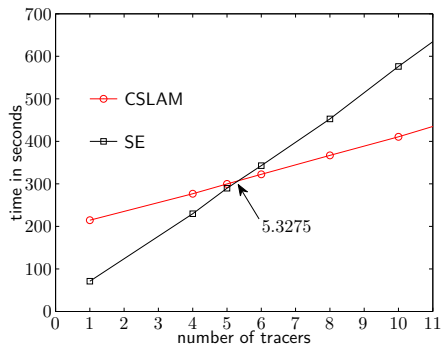
- SE = spectral element dynamical core in CAM/HOMME (Dennis et al., 2012)
- SE uses elements and each element has a quadrature grid
- CSLAM uses an equi-angular gnomonic finite-volume grid



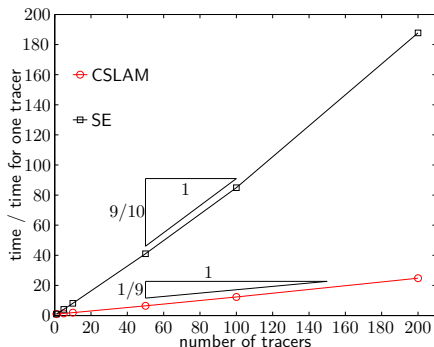
- Figure: (left) 'CSLAM grid' and (right) spectral element quadrature grid
- Infrastructure is being implemented to support coarser or finer finite-volume physics grid (physics grid may, of course, also simply coincide with CSLAM grid).

CSLAM implemented in CAM-SE for 'offline' transport (Erath et al., 2012)

- SE = spectral element dynamical core in CAM/HOMME (Dennis et al., 2012)
- SE max. Courant number (CN): $CN < 0.28$
- CSLAM max. Courant number for SE implementation: $CN < 1$



NCAR's Cray XT5m



These performance numbers are for exact trajectories!

'Online' coupling CAM-SE with CSLAM - ongoing work!

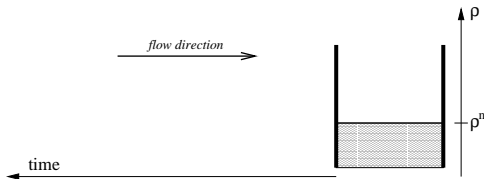
Constraints: mass-conservation, consistency, shape-preservation

CAM-SE predicted ρ_{SE} does, obviously, not match 'offline' ρ_{CSLAM} computed by CSLAM!

Possible solutions:

- overwrite ρ_{SE} with ρ_{CSLAM} ; unstable?
- nudge ρ_{SE} towards ρ_{CSLAM} ; unstable?
- switch to flux-form version of CSLAM (Harris et al., 2010) and use well-known finite-volume method for coupling: SE provides accumulated background flux of air mass and CSLAM provides average flux of q (satisfies all constraints!):

$$\text{Tracer mass flux} = \langle q \rangle_{CSLAM} \sum_{j=1}^{nsplit} \rho_{SE}^{(n+j/nsplit)}$$



'Online' coupling CAM-SE with CSLAM - ongoing work!

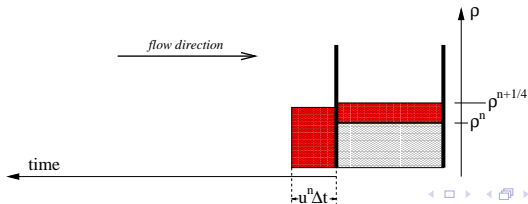
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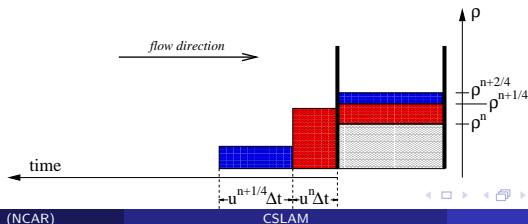
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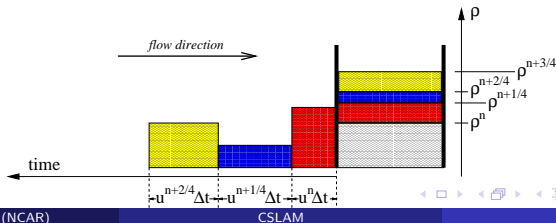
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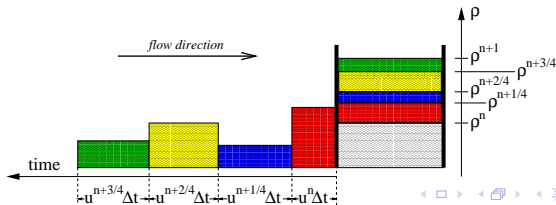
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'Online' coupling CAM-SE with CSLAM - ongoing work!

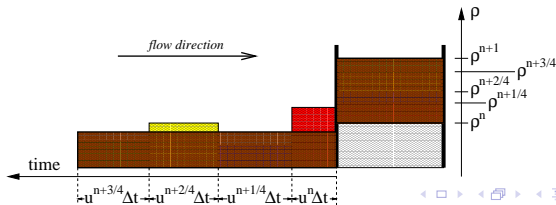
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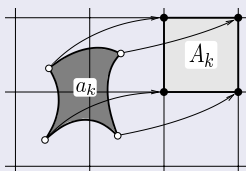
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semi-Lagrangian form



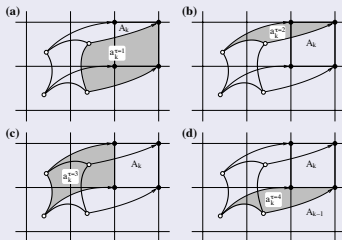
$$\frac{D}{Dt} \int_{A(t)} \psi dA = 0.$$

where $A(t)$ is a Lagrangian[†] control volume and

$$\frac{D}{Dt} = \frac{\partial}{\partial t} + \vec{v} \cdot \nabla,$$

is the material/total derivative.

Eulerian (flux-form) form



Integrate

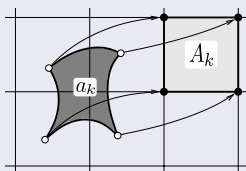
$$\frac{\partial \psi}{\partial t} + \nabla \cdot (\psi \vec{v}) = 0$$

over an Eulerian control volume A_k :

$$\frac{\partial}{\partial t} \int_{A_k} \psi dA + \int_{A_k} \nabla \cdot (\psi \vec{v}) dA = 0.$$

[†] volume whose bounding surface moves with the local fluid velocity \Leftrightarrow volume which always contains the same material particles

semi-Lagrangian form



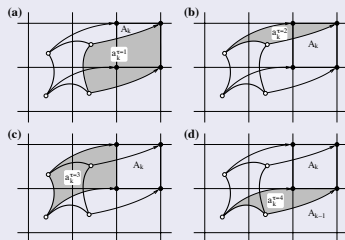
$$\frac{D}{Dt} \int_{A(t)} \psi dA = 0.$$

where $A(t)$ is a Lagrangian[†] control volume and

$$\frac{D}{Dt} = \frac{\partial}{\partial t} + \vec{v} \cdot \nabla,$$

is the material/total derivative.

Eulerian (flux-form) form



Apply divergence theorem on second term:

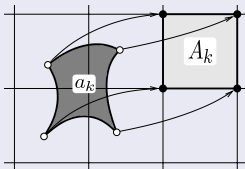
$$\frac{\partial}{\partial t} \int_{A_k} \psi dA + \oint_{\partial A_k} (\psi \vec{v}) \cdot \vec{n} dS = 0,$$

where ∂A_k is the boundary of A_k and \vec{n} the outward normal vector to ∂A_k .

→ instantaneous flux of tracer mass through boundaries of A_k

[†] volume whose bounding surface moves with the local fluid velocity ⇔ volume which always contains the same material particles

semi-Lagrangian form



$$\int_{A(t+\Delta t)} \psi dA = \int_{A(t)} \psi dA,$$

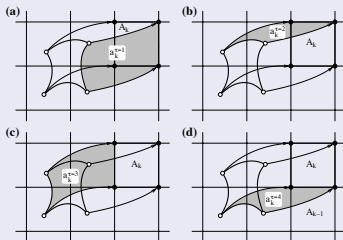
where Δt is time-step and $t = n \Delta t$.

Upstream semi-Lagrangian approach:

$$\overline{\psi}_k^{n+1} \Delta A_k = \overline{\psi}_k^n \Delta a_k,$$

where $\overline{(\)}$ is average value over cell.

Eulerian (flux-form) form



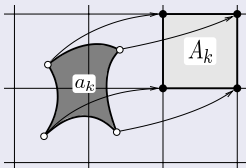
Apply divergence theorem on second term:

$$\frac{\partial}{\partial t} \int_{A_k} \psi dA + \oint_{\partial A_k} (\psi \vec{v}) \cdot \vec{n} dS = 0,$$

where ∂A_k is the boundary of A_k and \vec{n} the outward normal vector to ∂A_k .

→ instantaneous flux of tracer mass through boundaries of A_k

semi-Lagrangian form



$$\int_{A(t+\Delta t)} \psi dA = \int_{A(t)} \psi dA,$$

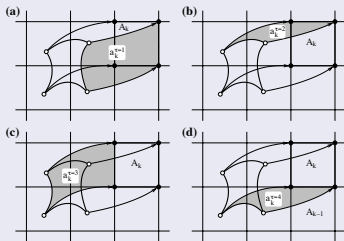
where Δt is time-step and $t = n \Delta t$.

Upstream semi-Lagrangian approach:

$$\overline{\psi}_k^{n+1} \Delta A_k = \overline{\psi}_k^n \Delta a_k,$$

where $\overline{(\)}$ is average value over cell.

Eulerian (flux-form) form



$$\frac{\partial}{\partial t} \int_{A_k} \psi dA + \oint_{\partial A_k} (\psi \vec{v}) \cdot \vec{n} dS = 0,$$

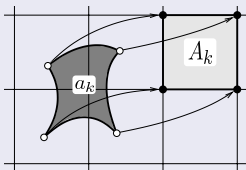
$$\overline{\psi}^{n+1} \Delta A_k = \overline{\psi}^n \Delta A_k +$$

$$\int_{n\Delta t}^{(n+1)\Delta t} \left[\oint_{\partial A_k} (\psi \vec{v}) \cdot \vec{n} dS \right] dt = 0,$$

→ flux of tracer mass through boundaries of A_k during $t \in [n\Delta t, (n+1)\Delta t]$

Finite-volume approach:

semi-Lagrangian form



$$\int_{A(t+\Delta t)} \psi dA = \int_{A(t)} \psi dA,$$

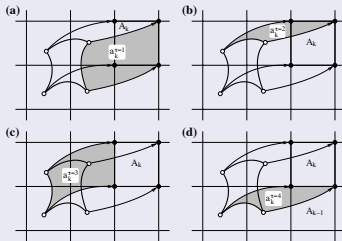
where Δt is time-step and $t = n \Delta t$.

Upstream semi-Lagrangian approach:

$$\overline{\psi}_k^{n+1} \Delta A_k = \overline{\psi}_k^n \Delta a_k,$$

where $\overline{(\)}$ is average value over cell.

Eulerian (flux-form) form



$$\overline{\psi}_k^{n+1} \Delta A_k = \overline{\psi}_k^n \Delta A_k - \sum_{\tau=1}^4 F_k^{(\tau)},$$

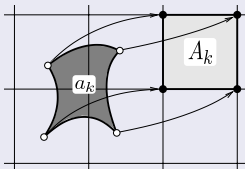
where

$$F_k^{(\tau)} = s_k^{(\tau)} \int_{a_k^{\tau}} \psi^n(x, y) dA.$$

is flux of mass through face τ during Δt ,
and $s_k^{(\tau)} = \pm 1$

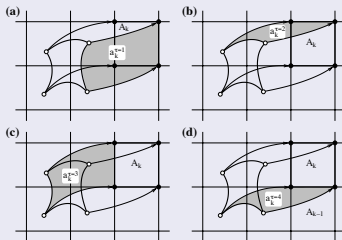
for simplicity assume s_k^{τ} is NOT multi-valued; for multi-valued case see, e.g., Harris et al. (2010).

semi-Lagrangian form



$$\bar{\psi}_k^{n+1} \Delta A_k = \bar{\psi}_k^n \Delta a_k,$$

Eulerian (flux-form) form



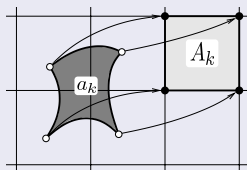
$$\bar{\psi}_k^{n+1} \Delta A_k = \bar{\psi}_k^n \Delta A_k - \sum_{\tau=1}^4 F_k^{(\tau)},$$

Note equivalence between Lagrangian cell-integrated and Eulerian flux-form continuity equations:

$$\Delta A_k - \sum_{\tau=1}^4 \left(s_k^{(\tau)} \Delta a_k^{(\tau)} \right) = \Delta a_k.$$

i.e. the areas involved in Eulerian forecast equals upstream Lagrangian area a_k .

semi-Lagrangian form



$$\bar{\psi}_k^{n+1} \Delta A_k = \bar{\psi}_k^n \Delta a_k,$$

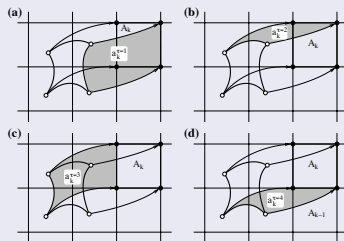
Define a global piecewise continuous reconstruction function

$$\psi(x, y) = \sum_{k=1}^N I_{A_k} \psi_k(x, y),$$

where I_{A_k} is the indicator function

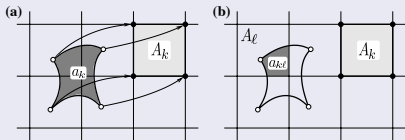
$$I_{A_k} = \begin{cases} 1, & (x, y) \in A_k, \\ 0, & (x, y) \notin A_k. \end{cases}$$

Eulerian (flux-form) form



$$\bar{\psi}_k^{n+1} \Delta A_k = \bar{\psi}_k^n \Delta A_k - \sum_{\tau=1}^4 F_k^{(\tau)},$$

semi-Lagrangian form



$$\bar{\psi}_k^{n+1} \Delta A_k = \bar{\psi}_k^n \Delta a_k,$$

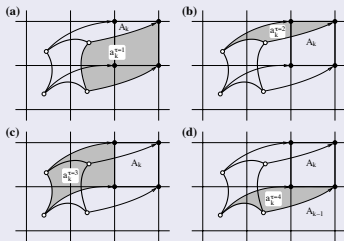
$$\bar{\psi}_k^{n+1} \Delta A_k = \sum_{\ell=1}^{L_k} \int_{a_{k\ell}} \psi_\ell^n(x, y) dA.$$

where $a_{k\ell}$ is the non-empty overlap area

$$a_{k\ell} = a_k \cap A_\ell, \quad a_{k\ell} \neq \emptyset; \quad \ell = 1, \dots, L_k,$$

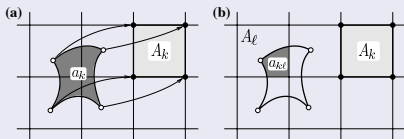
where N is the number of cells in the domain and L_k number of overlap areas.

Eulerian (flux-form) form



$$\bar{\psi}_k^{n+1} \Delta A_k = \bar{\psi}_k^n \Delta A_k - \sum_{\tau=1}^4 F_k^{(\tau)},$$

semi-Lagrangian form



$$\bar{\psi}_k^{n+1} \Delta A_k = \bar{\psi}_k^n \Delta a_k,$$

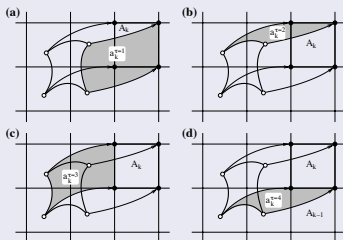
$$\bar{\psi}_k^{n+1} \Delta A_k = \sum_{\ell=1}^{L_k} \int_{a_{k\ell}} \psi_\ell^n(x, y) dA.$$

where $a_{k\ell}$ is the non-empty overlap area

$$a_{k\ell} = a_k \cap A_\ell, \quad a_{k\ell} \neq \emptyset; \quad \ell = 1, \dots, L_k,$$

where N is the number of cells in the domain and L_k number of overlap areas.

Eulerian (flux-form) form



$$\bar{\psi}_k^{n+1} \Delta A_k = \bar{\psi}_k^n \Delta A_k - \sum_{\tau=1}^4 F_k^{(\tau)},$$

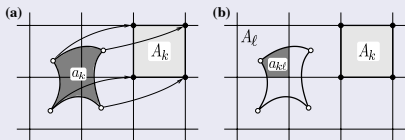
$$F_k^{(\tau)} = \sum_{\ell=1}^{L_k^{(\tau)}} \int_{a_{k\ell}} \psi_\ell^n(x, y) dA,$$

where $L_k^{(\tau)}$ is number of non-empty 'flux' overlap areas for face τ .

Note that in general: $L_k \ll \sum_{\tau=1}^4 L_k^{(\tau)}$

Finite-volume approach: Conditions for inherent mass-conservation

semi-Lagrangian form



$$\bar{\psi}_k^{n+1} \Delta A_k = \bar{\psi}_k^n \Delta a_k,$$

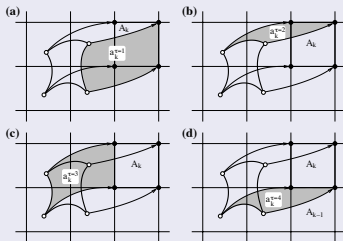
- a_k 's span Ω without gaps/overlaps

$$\bigcup_{k=1}^N a_k = \Omega, \text{ and } a_k \cap a_\ell = \emptyset \forall k \neq \ell.$$

- Sub-grid-scale representation of ψ must integrate to cell-average mass

$$\int_{A_k} \psi_k^n(x, y) dA = \bar{\psi}_k^n \Delta A,$$

Eulerian (flux-form) form



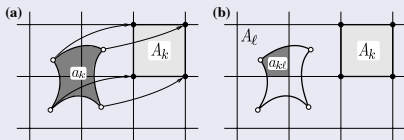
$$\bar{\psi}_k^{n+1} \Delta A_k = \bar{\psi}_k^n \Delta A_k - \sum_{\tau=1}^4 F_k^{(\tau)},$$

- Fluxes for 'shared' faces must cancel, e.g.,

$$F_k^{(3)} = -F_{k-1}^{(1)}$$

Any flux, even highly inaccurate fluxes, will NOT violate mass-conservation!

semi-Lagrangian form

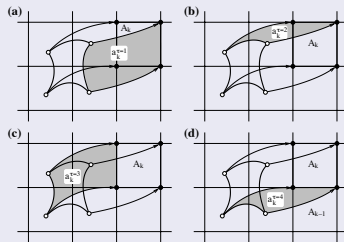


$$\bar{\psi}_k^{n+1} \Delta A_k = \bar{\psi}_k^n \Delta a_k,$$

The *only* direct way of enforcing shape-preservation is to filter the sub-grid-scale distribution $\psi_k^n(x, y)$:

- fully 2D filters (Barth and Jespersen, 1989)
- 1D filters for cascade schemes (Colella and Woodward, 1984; Zerroukat et al., 2005; Lin and Rood, 1996)

Eulerian (flux-form) form



$$\bar{\psi}_k^{n+1} \Delta A_k = \bar{\psi}_k^n \Delta A_k - \sum_{\tau=1}^4 F_k^{(\tau)},$$

Shape-preservation can be enforced by

- blending monotone and high-order fluxes (e.g., Flux-Corrected Transport Zalesak, 1979)
- making $\psi_k^n(x, y)$ shape-preserving (Barth and Jespersen, 1989)

Part II

Beyond linear transport: shallow-water model

- Shallow water equations on an f -plane:

$$\frac{\partial u}{\partial t} + u \frac{\partial u}{\partial x} + v \frac{\partial u}{\partial y} - f v - g \frac{\partial h}{\partial x} = 0$$

$$\frac{\partial v}{\partial t} + u \frac{\partial v}{\partial x} + v \frac{\partial v}{\partial y} + f u - g \frac{\partial h}{\partial y} = 0$$

$$\frac{\partial h}{\partial t} + \nabla \cdot (h \vec{v}) = 0$$

$$\frac{\partial (h q)}{\partial t} + \nabla \cdot (h q \vec{v}) = 0$$

where f Coriolis, h height, $\vec{v} = (u, v)$ horizontal velocity vector, g gravity.

- Momentum equation solved using traditional two-time-level semi-implicit scheme.
- Continuity equations solved with cell-integrated scheme: CSLAM
 - Semi-implicit time-stepping with cell-integrated Lagrangian schemes not straight forward (consistency, divergence discretization)

Traditionally: semi-Lagrangian advection of ρ is combined with semi-implicit time-stepping:

$$\bar{\rho}_k^{n+1} = (\bar{\rho}_k^{n+1})_{exp} - \frac{\Delta t}{2} \rho_{00} (\nabla \cdot \bar{v}_k^{n+1} - \nabla \cdot \tilde{v}_k^{n+1}),$$

where

- ρ_{00} a constant reference density
- $(\cdot)_{exp}$ is the explicit prediction
- \tilde{v}^{n+1} velocity extrapolated to time-level $(n + 1)$

What about tracers?

- Solving continuity equation for (ρq) explicitly

$$\bar{\rho} \bar{q}_k^{n+1} \Delta A_k = \bar{\rho} \bar{q}_k^n \Delta a_k$$

is NOT 'free-stream' preserving!

- Using 'traditional' semi-implicit approach for tracers

$$\bar{\rho} \bar{q}_k^{n+1} \Delta A_k = \bar{\rho} \bar{q}_k^n \Delta a_k - \frac{\Delta t}{2} (\rho q)_{00} (\nabla \cdot \bar{v}_k^{n+1} - \nabla \cdot \tilde{v}_k^{n+1}).$$

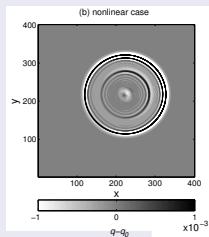
is 'free-stream' preserving but problematic (Lauritzen et al., 2008).

Traditionally: semi-Lagrangian advection of ρ is combined with semi-implicit time-stepping:

$$\bar{\rho}_k^{n+1} = (\bar{\rho}_k^{n+1})_{exp} - \frac{\Delta t}{2} \left\{ \nabla \cdot \left[(\bar{\rho}_k^{n+1})_{exp} \tilde{v}_k^{n+1} \right] - \nabla \cdot \left[(\bar{\rho}_k^n)_{exp} \tilde{v}_k^{n+1} \right] \right\}.$$

where

- ρ_{00} a constant reference density
- $(\cdot)_{exp}$ is the explicit prediction
- \tilde{v}^{n+1} velocity extrapolated to time-level $(n + 1)$



- Radially propagating gravity wave test (shallow water in Cartesian geometry; Wong et al., 2013b)
- Initial condition: $q = 1$
- Errors are $\mathcal{O}(10^{-3})$
- Problematic? Even when using a shape-preserving filter the semi-implicit correction term may render the scheme oscillatory and non-shape-preserving!

Traditionally: semi-Lagrangian advection of ρ is combined with semi-implicit time-stepping:

$$\bar{\rho}_k^{n+1} = (\bar{\rho}_k^{n+1})_{exp} - \frac{\Delta t}{2} \left\{ \nabla \cdot \left[(\bar{\rho}_k^{n+1})_{exp} \bar{v}_k^{n+1} \right] - \nabla \cdot \left[(\bar{\rho}_k^n)_{exp} \tilde{v}_k^{n+1} \right] \right\}.$$

where

- ρ_{00} a constant reference density
- $(\cdot)_{exp}$ is the explicit prediction
- \tilde{v}^{n+1} velocity extrapolated to time-level $(n + 1)$

What about tracers?

- A solution is to formulate the semi-implicit terms in flux-form

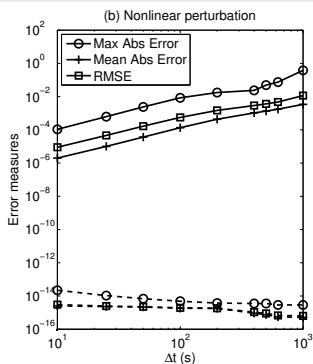
$$\overline{\rho q}_k^{n+1} = (\overline{\rho q}_k^{n+1})_{exp} - \frac{\Delta t}{2} \left\{ \nabla \cdot \left[(\overline{\rho q}_k^{n+1})_{exp} \bar{v}_k^{n+1} \right] - \nabla \cdot \left[(\overline{\rho q}_k^n)_{exp} \tilde{v}_k^{n+1} \right] \right\}.$$

so that reference states are eliminated (Wong et al., 2013b)

Time-stepping and coupling: mass-conservative semi-implicit approach

Traditionally: semi-Lagrangian advection of ρ is combined with semi-implicit time-stepping:

$$\bar{\rho}_k^{n+1} = (\bar{\rho}_k^{n+1})_{exp} - \frac{\Delta t}{2} \left\{ \nabla \cdot \left[(\bar{\rho}_k^{n+1})_{exp} \bar{v}_k^{n+1} \right] - \nabla \cdot \left[(\bar{\rho}_k^n)_{exp} \tilde{v}_k^{n+1} \right] \right\}.$$



- Radially propagating gravity wave test: error measures for q as a function of Δt
 - solid lines is 'problematic' formulation
 - dash lines is new formulation
- Initial condition: $q = 1$
- Errors in semi-implicit correction term increase with increasing Δt
- New formulation is 'free-stream preserving' and shape-preserving!
- Both formulations are stable for long Δt 's

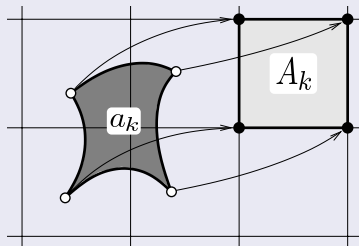
Discretization of divergence

In traditional semi-implicit semi-Lagrangian scheme, divergence is usually discretized with finite-differences:

$$\nabla_{eul} \cdot \vec{v} = \frac{u_{i+1j} - u_{ij}}{\Delta x} + \frac{v_{ij+1} - v_{ij}}{\Delta y}, \quad (6)$$

however, cell-integrated schemes 'see' a Lagrangian discretization of divergence based on area change:

$$\nabla_{lgr} \cdot \vec{v} = \frac{1}{\Delta A_k} \frac{\Delta A_k - \delta A_k}{\Delta t} \quad (7)$$



Simple example graphically illustrating difference between $\nabla_{eul} \cdot \vec{v}$ and $\nabla_{lgr} \cdot \vec{v}$

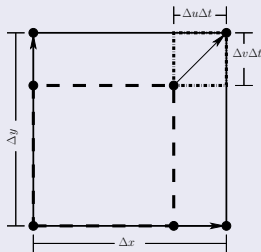


Figure: Assume the following velocity components at cell vertices

$$\vec{v}_{SW} = (0, 0),$$

$$\vec{v}_{NW} = (0, v),$$

$$\vec{v}_{SE} = (u, 0),$$

$$\vec{v}_{NE} = (u, v),$$

where standard compass notation has been used.

Eulerian discretization of divergence:

$$\nabla_{eul} \cdot \vec{v} = \frac{u}{\Delta x} + \frac{v}{\Delta y} \quad (6)$$

Lagrangian (cell-integrated) discretization of divergence:

$$\nabla_{lgr} \cdot \vec{v} = \frac{u}{\Delta x} + \frac{v}{\Delta y} - \Delta t \frac{u v}{\Delta x \Delta y}. \quad (7)$$

→ differ by non-linear term!

Simple example graphically illustrating difference between $\nabla_{eul} \cdot \vec{v}$ and $\nabla_{lgr} \cdot \vec{v}$

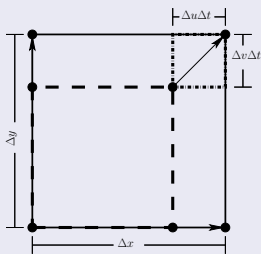


Figure: Assume the following velocity components at cell vertices

$$\vec{v}_{SW} = (0, 0),$$

$$\vec{v}_{NW} = (0, v),$$

$$\vec{v}_{SE} = (u, 0),$$

$$\vec{v}_{NE} = (u, v),$$

where standard compass notation has been used.

To have consistency with CSLAM (use $\nabla_{lgr} \cdot \vec{v}$) and retain a Helmholtz equation for the semi-implicit solve, the continuity equation is discretized as follows

$$\begin{aligned} \overline{\rho q_k}^{n+1} = & (\overline{\rho q_k}^{n+1})_{exp} - \frac{\Delta t}{2} \left\{ \nabla_{eul} \cdot \left[(\overline{\rho q_k}^{n+1})_{exp} \vec{v}_k^{n+1} \right] - \nabla_{lgr} \cdot \left[(\overline{\rho q_k}^n)_{exp} \tilde{v}_k^{n+1} \right] \right\} \\ & + \frac{\Delta t}{2} \overline{\left\{ \nabla_{eul} \cdot \left[(\overline{\rho q_k}^{n+1})_{exp} \vec{v}_k^{n+1} \right] - \nabla_{lgr} \cdot \left[(\overline{\rho q_k}^n)_{exp} \tilde{v}_k^{n+1} \right] \right\}} \frac{\delta a_k}{\Delta A_k}, \quad (6) \end{aligned}$$

(Lauritzen et al., 2006; Wong et al., 2013b)

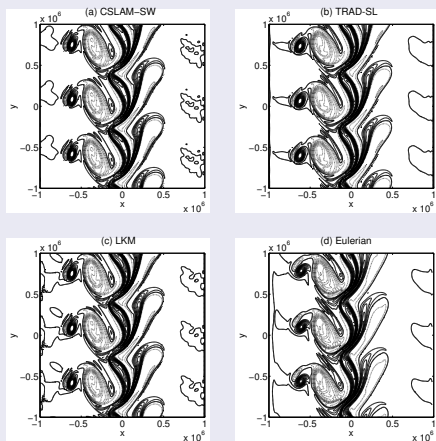
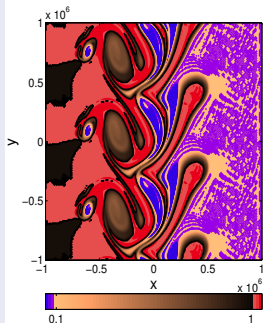


Figure: Vorticity

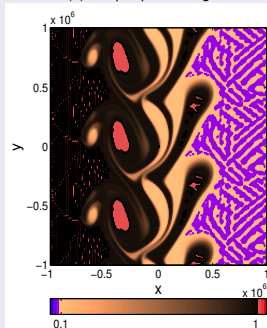
- (a) Traditional grid-point method for momentum equations and CSLAM for mass
- (b) Traditional grid-point method for momentum and continuity equations
- (c) Same as (a) but with 'problematic' semi-implicit method (LKM)
- (d) Eulerian discretization (semi-implicit leapfrog; Asselin filter)

(a) Non-shape-preserving CSLAM-SW



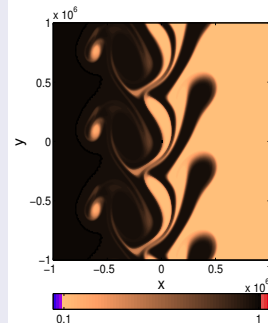
Specific concentration, q

(b) Shape-preserving LKM



Specific concentration, q

(c) Shape-preserving CSLAM-SW



Specific concentration, q

- Note: even with shape-preserving filter on explicit advection the semi-implicit correction terms render solution non-shape-preserving for the 'problematic formulation (LKM)

Part III

A compressible nonhydrostatic cell-integrated semi-Lagrangian semi-implicit solver (CSLAM-NH) with consistent and conservative transport

$$\frac{\partial u}{\partial t} + u \frac{\partial u}{\partial x} + w \frac{\partial u}{\partial z} = -\frac{\pi}{\rho_m} \gamma R_d \frac{\partial \Theta'_m}{\partial x} + F_u, \quad (7)$$

$$\frac{\partial w}{\partial t} + u \frac{\partial w}{\partial x} + w \frac{\partial w}{\partial z} = -\frac{\pi}{\rho_m} \gamma R_d \frac{\partial \Theta'_m}{\partial z} + \frac{g}{\rho_m} \left[\bar{\rho}_d \frac{\pi'}{\bar{\pi}} - \rho'_m \right] + F_w, \quad (8)$$

$$\frac{\partial \Theta_m}{\partial t} + \nabla \cdot (\Theta_m \mathbf{v}) = F_\Theta, \quad (9)$$

$$\frac{\partial \rho_d}{\partial t} + \nabla \cdot (\rho_d \mathbf{v}) = 0, \quad (10)$$

$$\frac{\partial Q_j}{\partial t} + \nabla \cdot (Q_j \mathbf{v}) = F_{Q_j}, \quad (11)$$

$$p = p_0 \left(\frac{R_d \Theta_m}{p_0} \right)^\gamma, \quad (12)$$

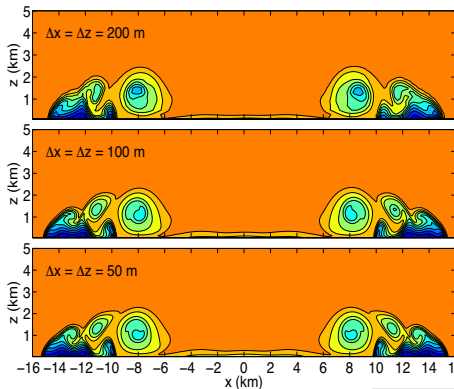
(Klemp et al., 2007)

- Equations linearized about hydrostatically balanced background state
- Momentum equations cast in their advective form
- All other equations (density, potential temperature, moist species, cast in their conservative flux-form).

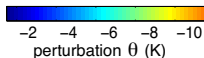
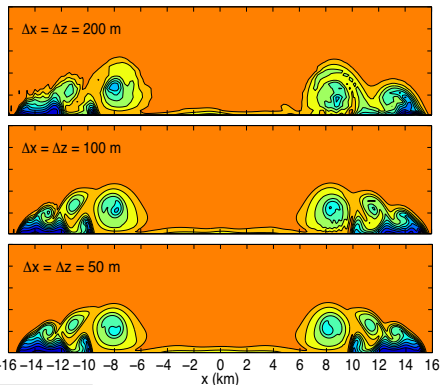
Density current with mean background flow (Straka et al., 1992)

Wong et al. (2013a):

CSLAM-NH



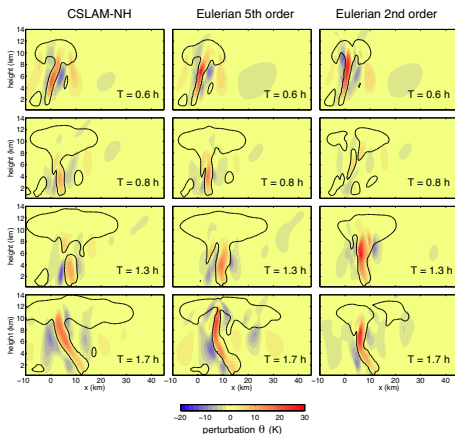
Eulerian Split-Explicit



- Symmetric solutions!
- Stable with $2\times$ split-explicit time-step

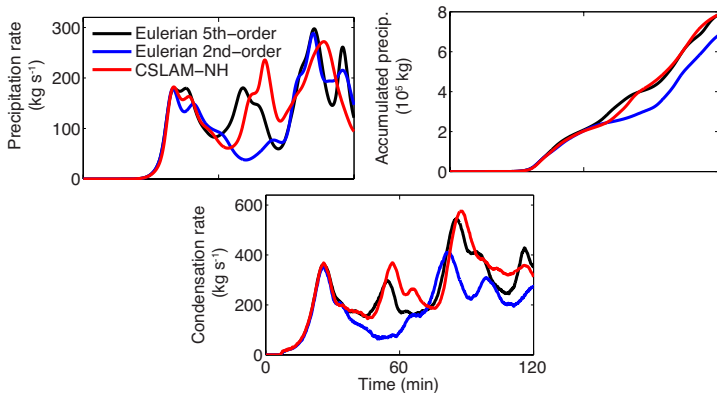
2D squall line (Weisman and Klemp, 1982)

Kessler microphysics scheme: diagnoses 'warm rain' (source/sink for water vapour, cloud water, and rainwater; latent heat release adjusts potential temperature).



- Vertical velocity (colored contours); solid contour - convective cloud structure
- CSLAM-NH looks more like 5th-order 'WRF' solution than 2nd-order

2D squall line (Weisman and Klemp, 1982)



- Moisture statistics (Wong et al., 2013a)
- M. Wong is currently working on 'adding' topography to the CSLAM-NH!



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