



**SciDAC**  
Scientific Discovery through  
Advanced Computing

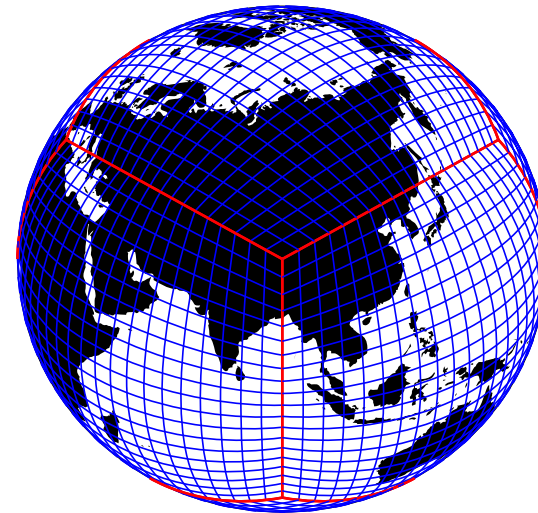
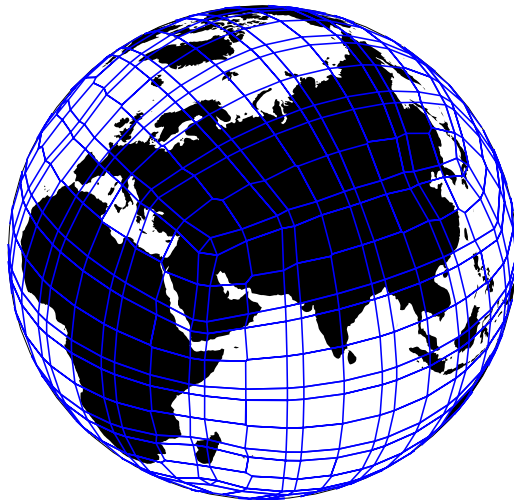


**CGD**  
Climate & Global Dynamics



# CAM dynamics update

**Peter Hjort Lauritzen<sup>§</sup>, M.A. Taylor<sup>\*</sup>, J.-F. Lamarque<sup>§</sup>, J. Bacmeister<sup>§</sup>,  
F. Vitt<sup>§</sup>, S. Goldhaber<sup>§</sup>, C. Hannay<sup>§</sup>,  
A. Conley<sup>§</sup>, B. Eaton<sup>§</sup>, R.B. Neale<sup>§</sup>, ...**



**AMWG winter meeting, February 10-12, 2014**

**§ NCAR    \*Sandia National Laboratories**



# Overview

## CAM-SE:

- New default time-stepping
- Axial angular momentum conservation
- Physics-grid and CSLAM transport
- Capability of doing offline simulations driven by meteorological analysis for chemistry applications

Leaving out lots of HOMME development:  
non-hydrostatic DG (R.D. Nair & R. Kloefkorn), non-hydrostatic SE  
(R.D. Nair & D. Hall), implicit time-stepping (K. Evans), ...

## Other:

- Energy definition in CAM (covered by D.L. Williamson)
- Nudging (“on the physics time-step”) – J.T. Bacmeister
- MPAS in CAM





# Trunk changes to time-stepping



- `tstep_type=5`:  
Switched to a 5-stage Runge-Kutta time-stepping;  
based on Kinnmark and Gray (1984) with a modification (Ullrich; unpublished) to  
make it non-linearly 3<sup>rd</sup>-order in time (implemented by M.A. Taylor)
- User confusion on CAM namelist:  
(e.g., split namelist variables do not mean the same thing in CAM-SE as CAM-FV)

$$\Delta t_{phys} = dtime,$$

$$\Delta t_{remap} = \frac{\Delta t_{phys}}{se\_nsplit},$$

$$\Delta t_{tracer} = \frac{\Delta t_{remap}}{rsplit},$$

$$\Delta t_{dyn} = \frac{\Delta t_{tracer}}{qsplit},$$

resolution	dtime	se_nsplit	rsplit	hypervis_ subcycle	$\Delta t_{remap}$ [s]	$\Delta t_{tracer}$ [s] = $\Delta t_{dyn}$ [s]	$\Delta t_{hypervis}$ [s]	$\nu$ [m <sup>4</sup> /s]
ne11np4 <sup>a</sup>	1800 (1800)	1 (5)	2 (2)	3 (1)	1800 (360)	900 (180)	300 (180)	$2.0 \times 10^{16}$ ( $2.0 \times 10^{16}$ )
ne16np4 <sup>b</sup>	1800 (1800)	1 (5)	3 (3)	3 (1)	1800 (360)	600 (120)	200 (120)	$7.0 \times 10^{15}$ ( $7.0 \times 10^{15}$ )
ne30np4	1800 (1800)	2 (10)	3 (3)	3 (1)	900 (180)	300 (60)	100 (60)	$1.0 \times 10^{15}$ ( $1.0 \times 10^{15}$ )
ne60np4	1800 (1800)	4 (20)	3 (3)	4 (1)	450 (90)	150 (30)	37.5 (30)	$1.0 \times 10^{14}$ ( $1.0 \times 10^{14}$ )
ne120np4 <sup>c</sup>	900 (900)	4 (20)	3 (3)	4 (1)	225 (45)	75 (15)	18.75 (15)	$1.0 \times 10^{13}$ ( $1.0 \times 10^{13}$ )
ne240np4	600 <sup>d</sup> (600)	5 (25)	3 (3)	4 (1)	120 (24)	40 (8)	10 (8)	$1.1 \times 10^{12}$ ( $1.1 \times 10^{12}$ )

<sup>a</sup>untested

<sup>b</sup>untested

<sup>c</sup>if winds are maximum 600 m/s; for CAM it is 120 m/s

<sup>d</sup>900 works, however, gravity wave noise!

qsplitt=1

<http://www.cgd.ucar.edu/cms/pel/software/cam-se-dt-table.pdf>

- assuming Lagrangian vertical coordinate





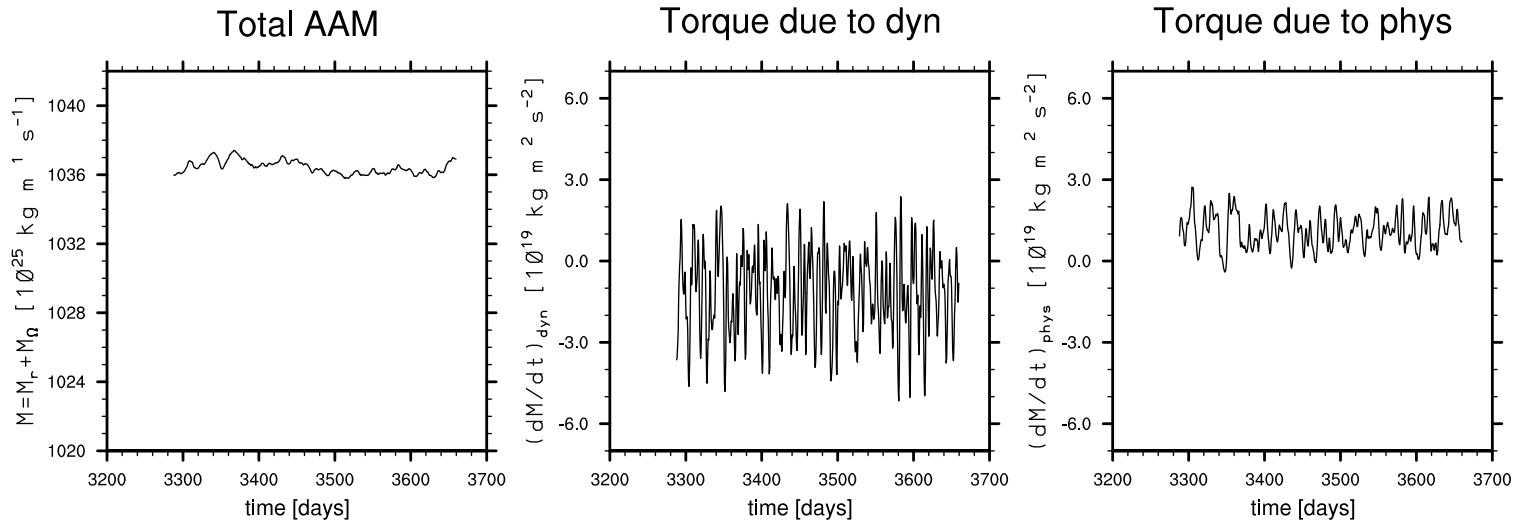
# Axial angular momentum analysis

Held-Suarez forcing (flat Earth => no mountain torque)



CAM-FV: finite-volume (Lin, 2004) dynamical core in CAM

Axial Angular Momentum (AAM) diagnostics for CAM-FV



M=global integrated  
axial angular  
momentum

$$0 \sim \left( \frac{dM}{dt} \right)_{dyn} \ll \left( \frac{dM}{dt} \right)_{phys}$$

NOT fulfilled!

Lebonnois et al., (2012)



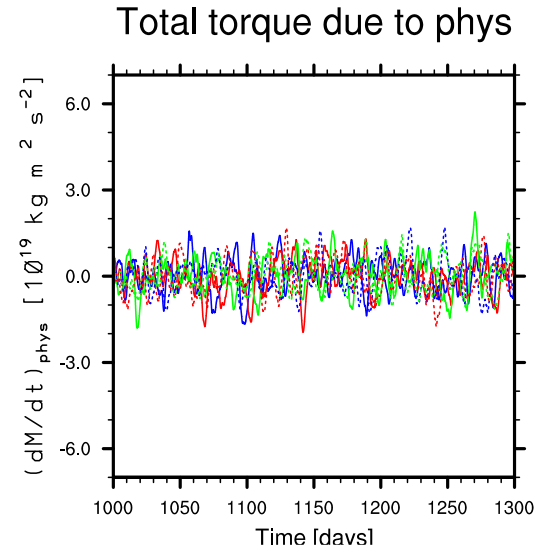
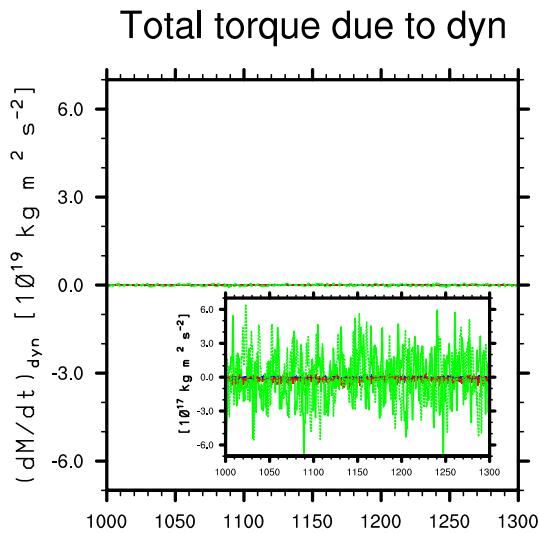


# Axial angular momentum analysis

Held-Suarez forcing (flat Earth => no mountain torque)



CAM-SE results:



np=2  
np=3  
np=4

$$0 \sim \left( \frac{dM}{dt} \right)_{dyn} \ll \left( \frac{dM}{dt} \right)_{phys}$$

Lauritzen et al. (2014; in press)

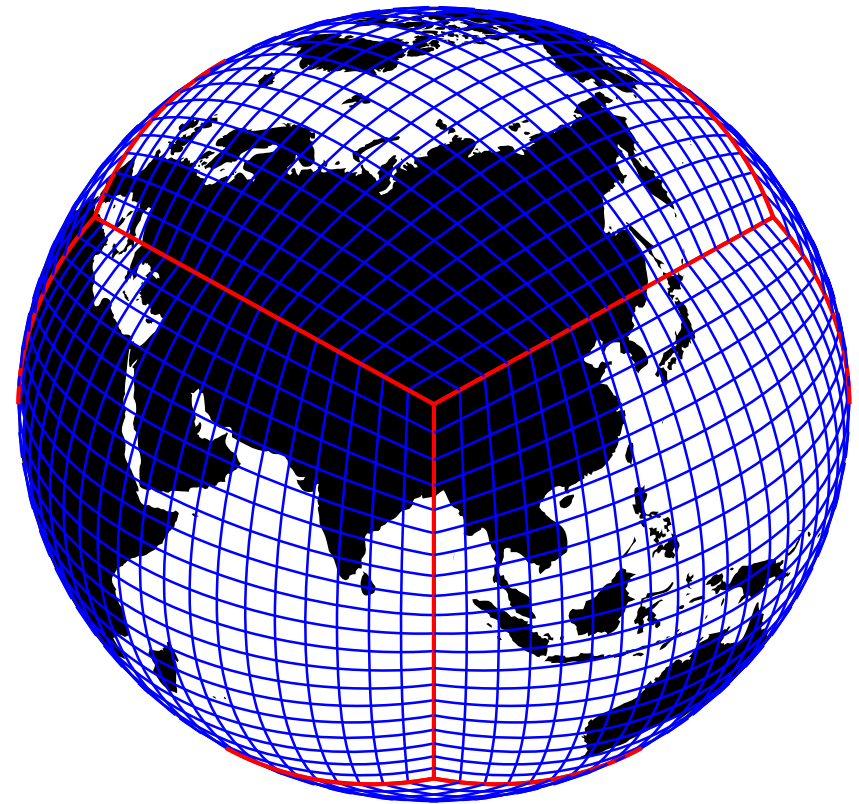


# Separating physics and dynamics grids

Current physics/"coupler" grid



Finite-volume equi-angular gnomonic grid





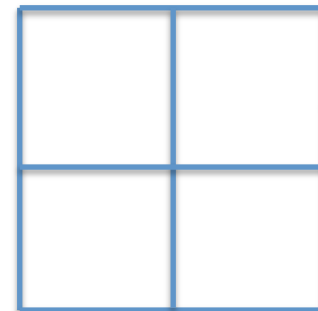
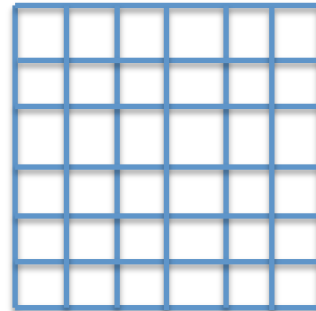
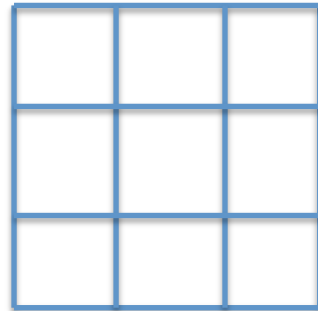
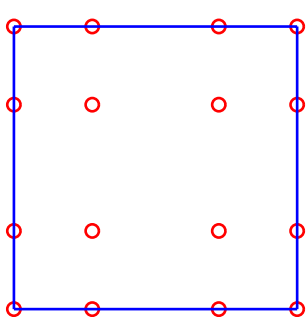
# Separating physics and dynamics grids

**6 month+ of re-engineering of CAM history output .... (S. Goldhaber)**

Main tasks:

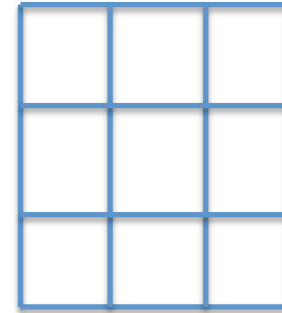
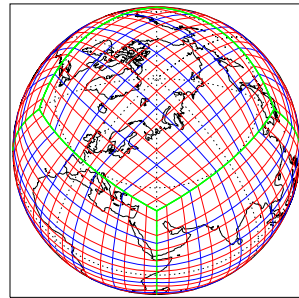
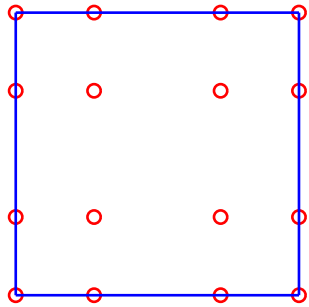
- enable output on an arbitrary physics grid (different from dynamics grid)
- remove assumptions in physics assuming dynamics-physics points co-located
- interface code for SE stored entire grid on every MPI task – fixed! (should help scalability on small memory massively parallel machines)
- dp\_coupling is now able to support physics grid
- tools to create IC files with different grids in one file

Current physics grid is equal-area finite-volume-type grid in each element  
(support coarser, finer, or similar resolution within each element)



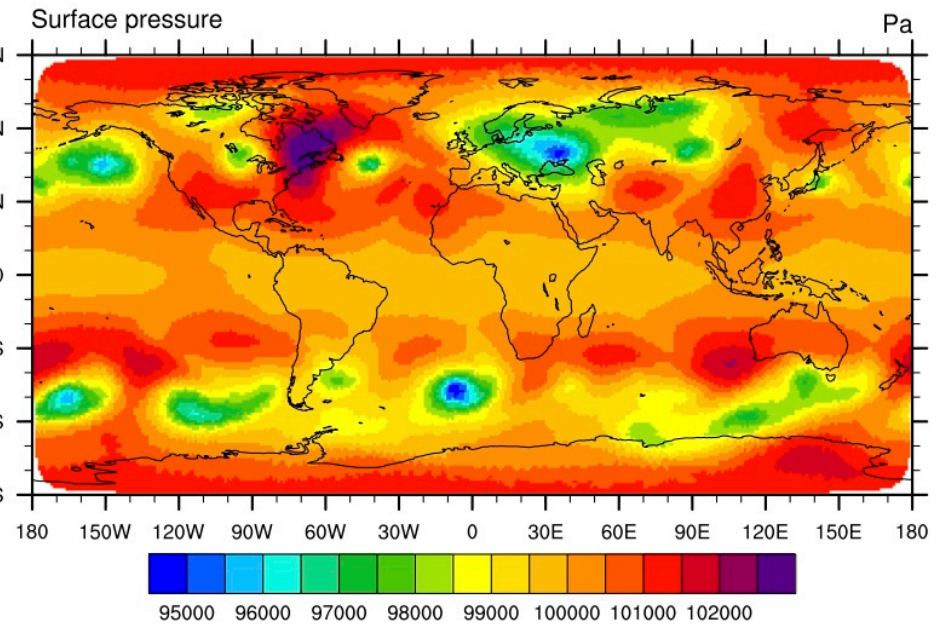
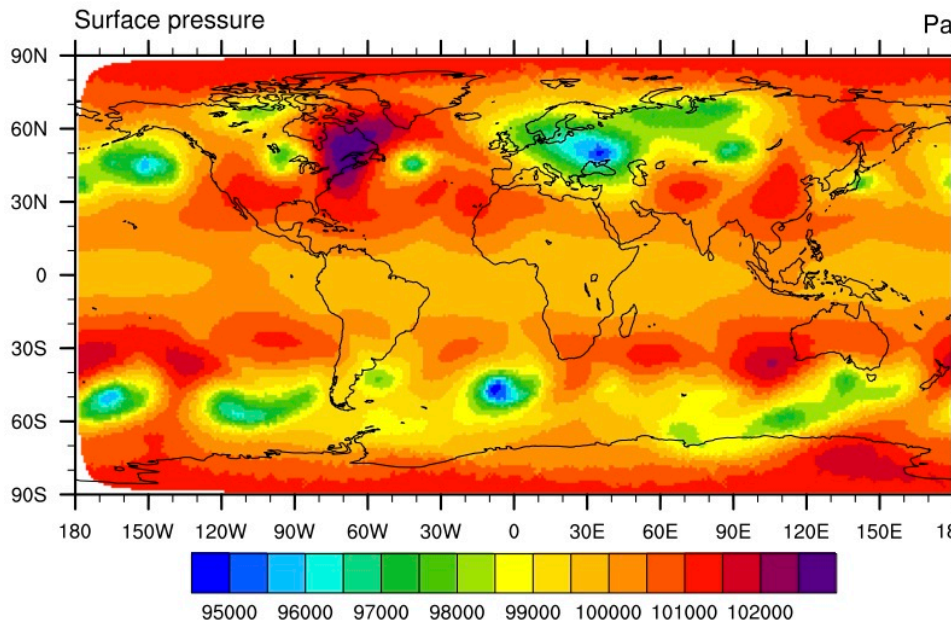


# Held-Suarez runs with physics grid



**Control, day 9**

**Physics grid (NC3), day 9**



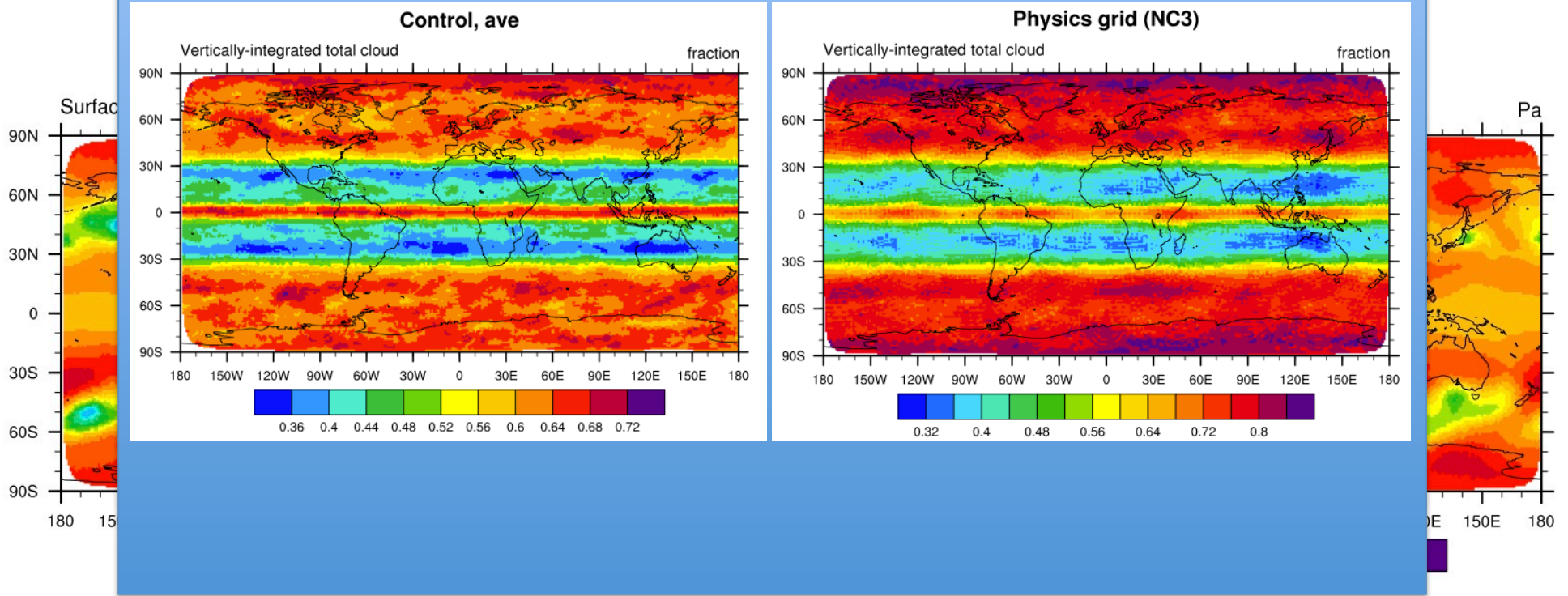




# Held-Suarez runs with physics grid



Currently debugging the Aqua-planet run





# Separating physics and dynamics grids

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## Main tasks:

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Current physics grid is equal-area finite-volume-type grid in each element

## Todo:

- flush out bugs ...
- longer term: generalize mapping to support arbitrary physics grids  
(for example, this will support mesh-refinement in the dynamical core and run physics on a uniform resolution grid)
- enforce total energy conservation in mapping process

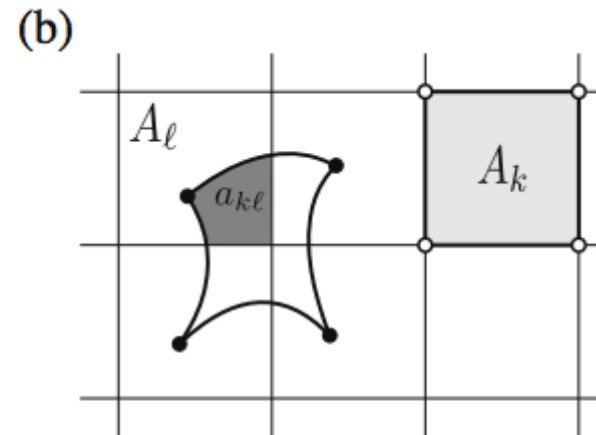
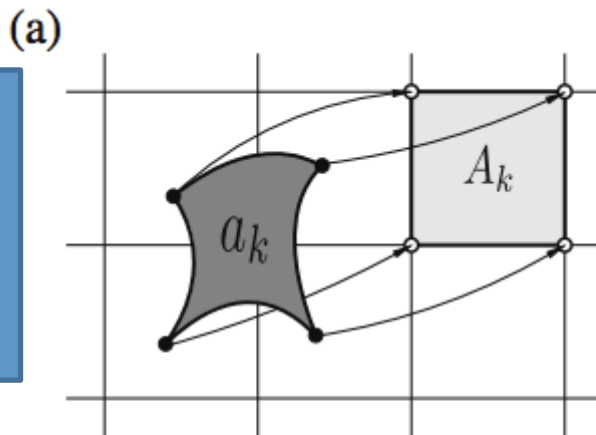




# Multi-tracer transport (CSLAM) in CAM-SE

- Stand-alone Lagrangian CSLAM scheme has been in HOMME for a while (Erath et al. 2012); also other option: SPELT (Erath and Nair, 2014)

Allows for long time-steps  
=> less MPI communication



Computing overlap areas is expensive but weights can be re-used for each additional tracer

**Fig. 1.** A schematic illustration of concepts used in the semi-Lagrangian finite-volume scheme. (a) The deformed departure cell  $a_k$  (dark shaded area) ends up, after being transported by the flow for one time-step, at the regular arrival cell  $A_k$  (light shaded area). The trajectories for the cell vertices are shown with arrows, and the departure and arrival cell vertices are marked with filled and open circles, respectively. (b) Illustrates the overlap region between the grid cell  $A_\ell$  and the departure cell  $a_k$  referred to as  $a_{k\ell}$  used for the upstream integral computation given in Eq. (4).

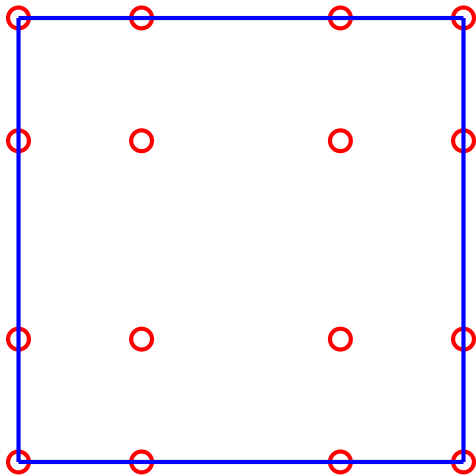




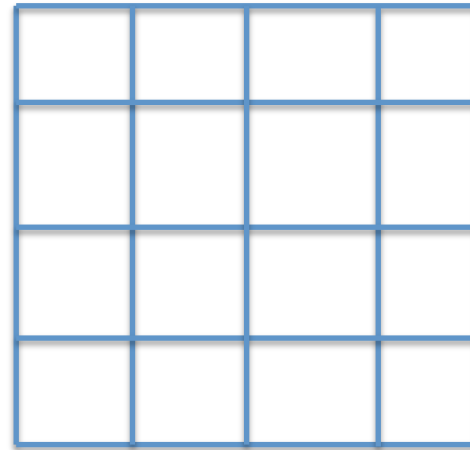
# Multi-tracer transport (CSLAM) in CAM-SE

Now that physics grid infrastructure is maturing we can start focusing on CSLAM transport (CSLAM scheme is using a finite-volume (quasi equal-area) grid and needs new physics grid infrastructure in CAM)

SE (np=4)



CSLAM (nc=4)





# Multi-tracer transport (CSLAM) in CAM-SE

**Now that physics grid infrastructure is maturing we can start focusing on CSLAM transport (CSLAM scheme is using a finite-volume (quasi equal-area) grid and needs new physics grid infrastructure in CAM)**

- Stand-alone Lagrangian CSLAM scheme has been in HOMME for a while (Erath et al. 2012). **However, Erath et al. (2012) did not consider coupling with CAM-SE air density**
- To couple CSLAM scheme with CAM-SE we are using a conventional flux-form methodology (used in, for example, CAM-FV):
  - convert Lagrangian CSLAM to flux-form (90% done in HOMME; Lauritzen)
  - compute finite-volume type fluxes from CAM-SE (method derived by Taylor and Ullrich)

NCAR-Sandia is working on uniform resolution implementation (Lauritzen-Taylor)  
Argonne-Sandia are working on variable resolution implementation (PI: F. Hoffman)

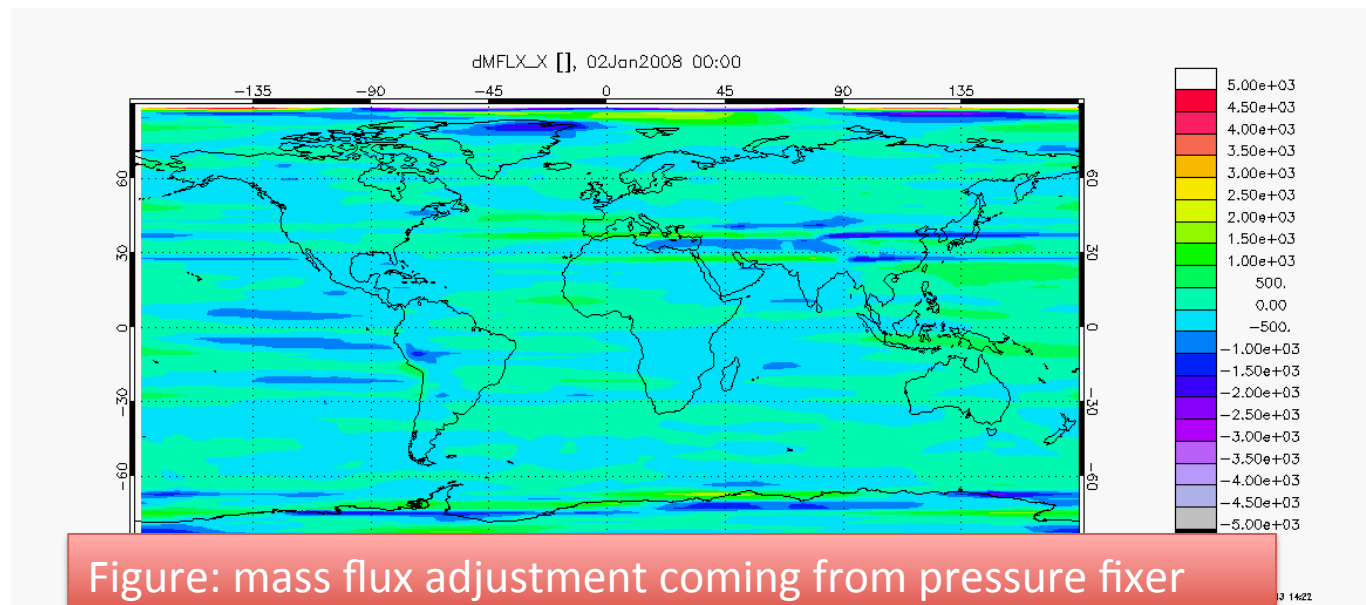


# Capability for doing offline simulations driven by (GEOS5) meteorological analysis in CAM-SE

## focus: chemistry applications

J.-F. Lamarque (PI), F. Vitt, A. Conley, P.H. Lauritzen

Current method in CAM-FV (CAM-Chem): Overwrite u,v,T,PS at every physics time-step and apply mass-fixer (directionally biased) to enforce consistency between internal mass-fluxes and driving data





# Capability for doing offline simulations driven by (GEOS5) meteorological analysis in CAM-SE

## focus: chemistry applications

J.-F. Lamarque (PI), F. Vitt, A. Conley, P.H. Lauritzen

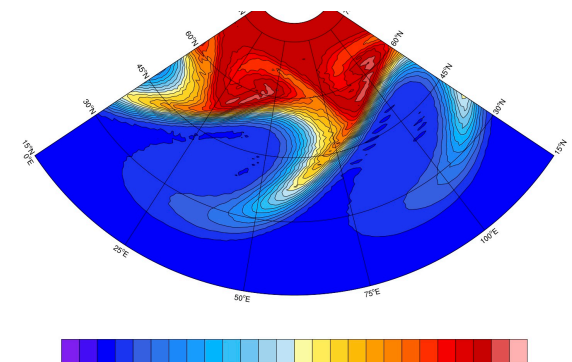
Method implemented in CAM-SE:

- Apply nudging to  $u, v, T$  (not PS)
- The nudging is implemented as a forcing term inside the dynamical core (at every Runge-Kutta step the dynamical core “feels” the nudging)

Does it work? Is PS nudged towards offline PS?

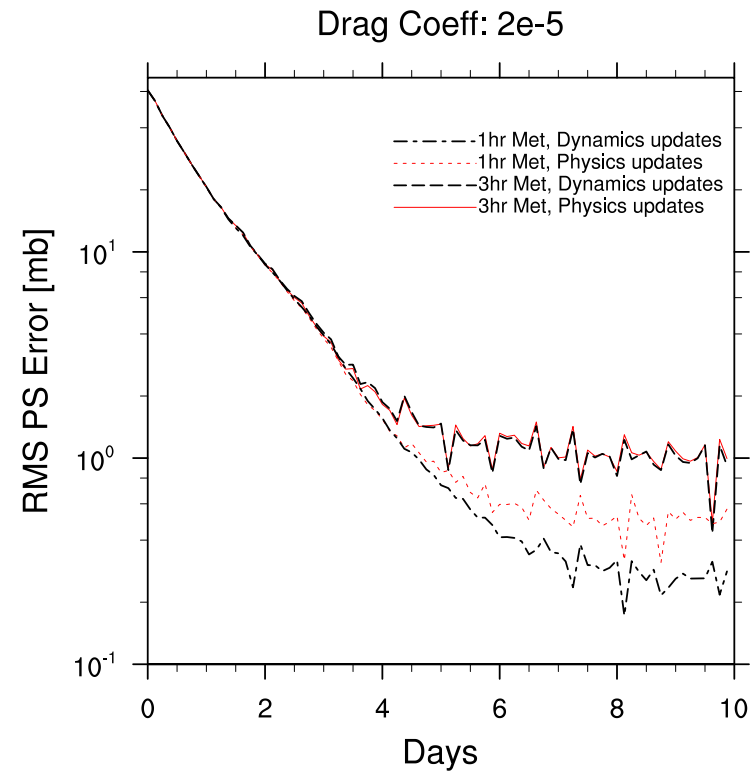
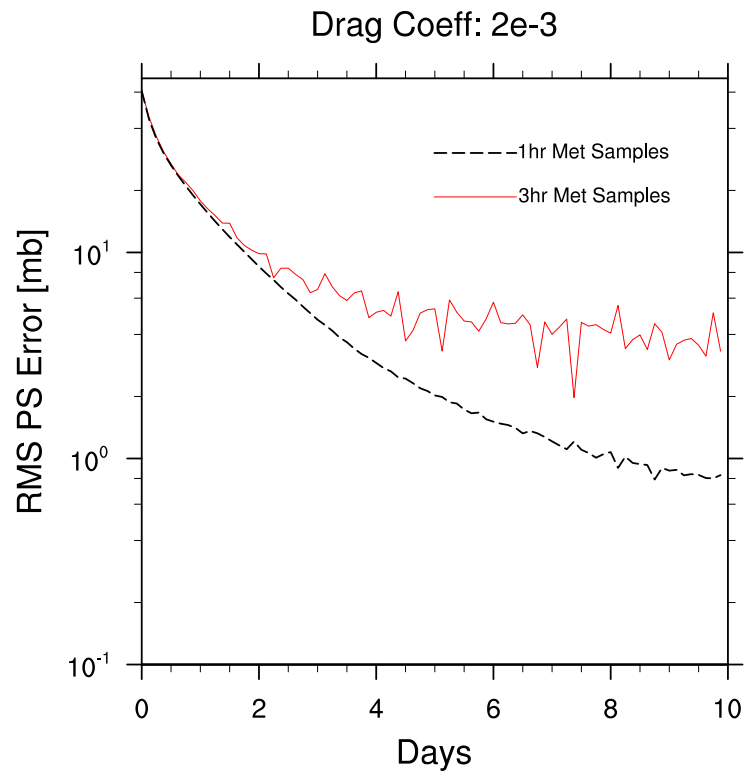
Idealized test:

Initial condition: Polvani baroclinic wave at time  $T=10$  days  
Force it to time-evolving solution starting from day 0





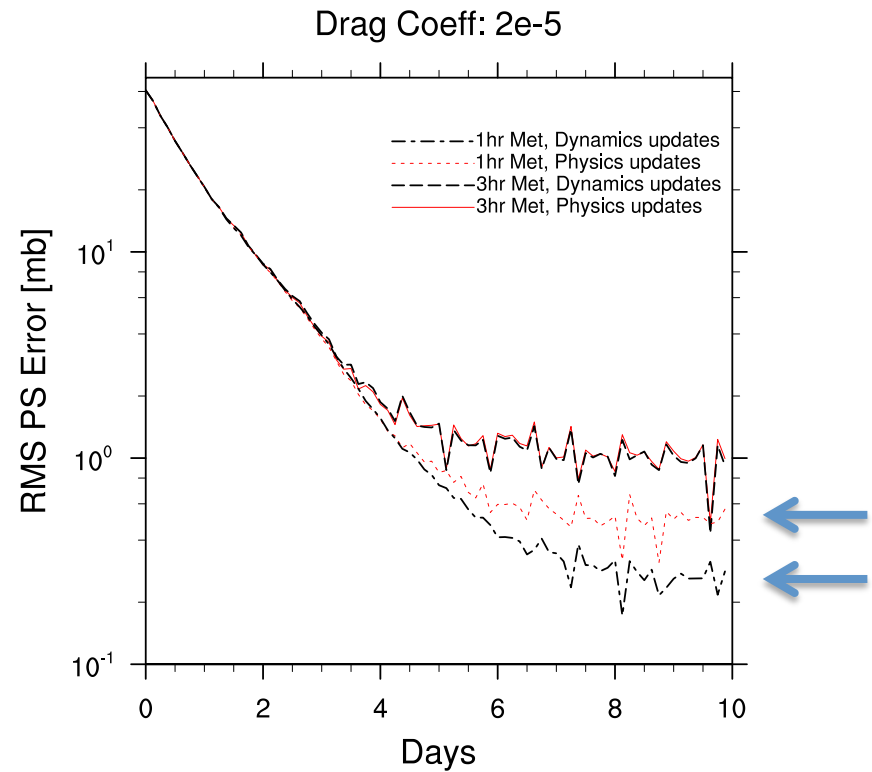
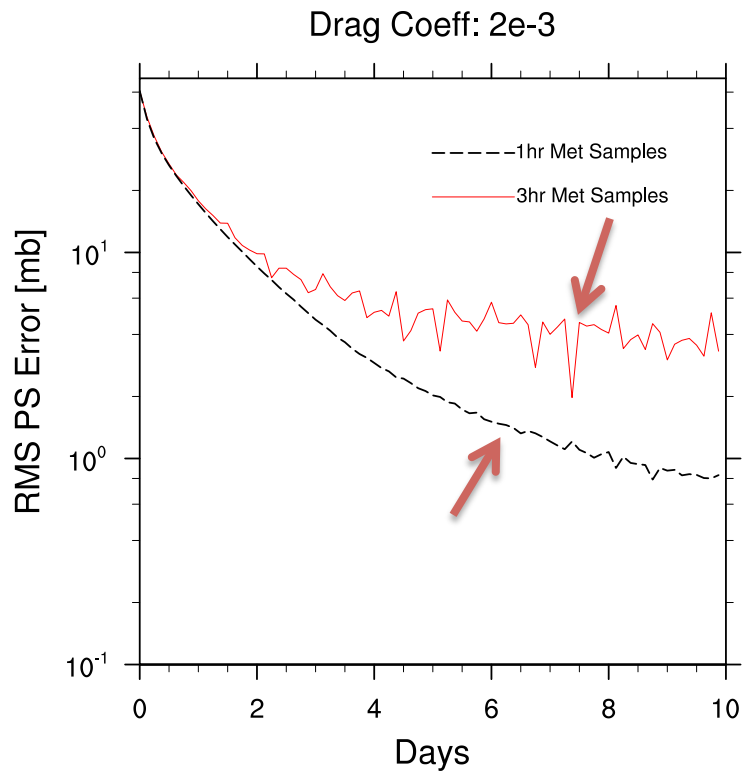
- Temporal resolution of meteorology is the largest error
- If you nudge weaker you get better results (time evolution is not linear).
- Updating nudging term every dynamics time-step (linear temporal interpolation between met field updates) only improves results if met fields are updated hourly







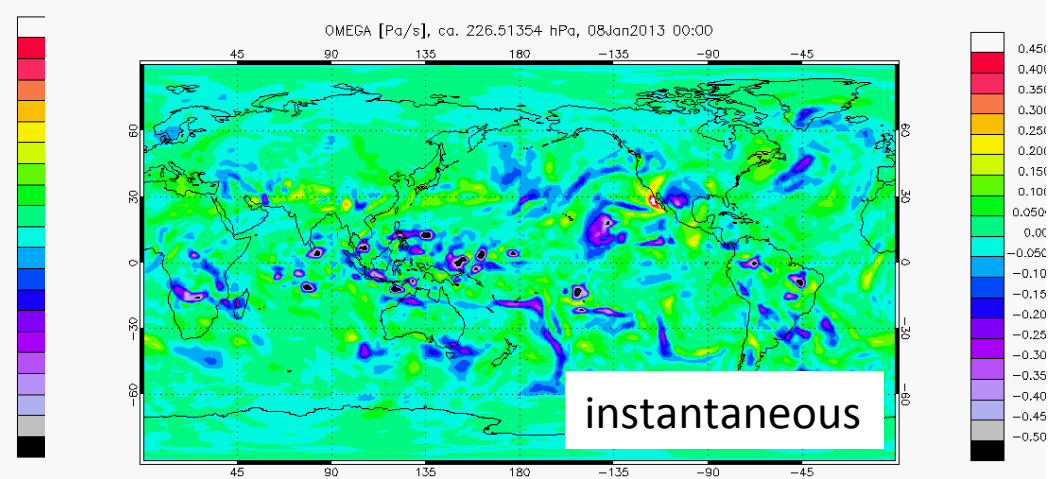
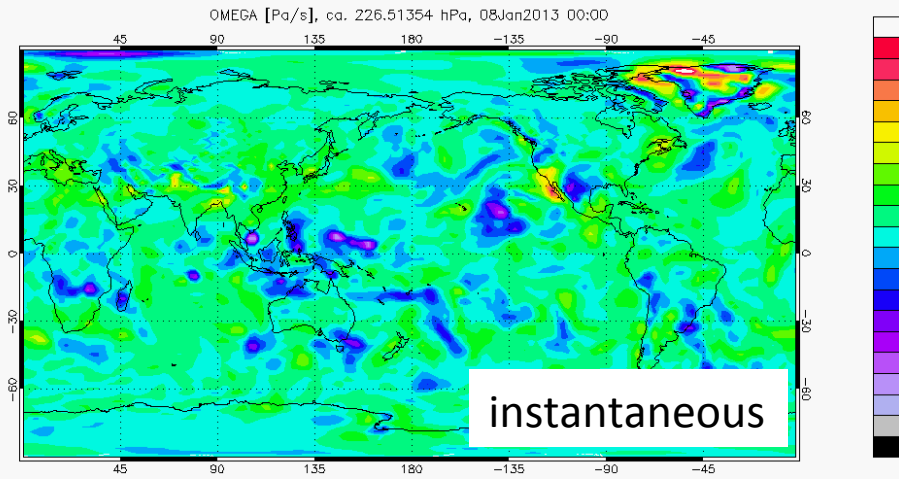
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FV:1.9x2.5

# “Full” model results

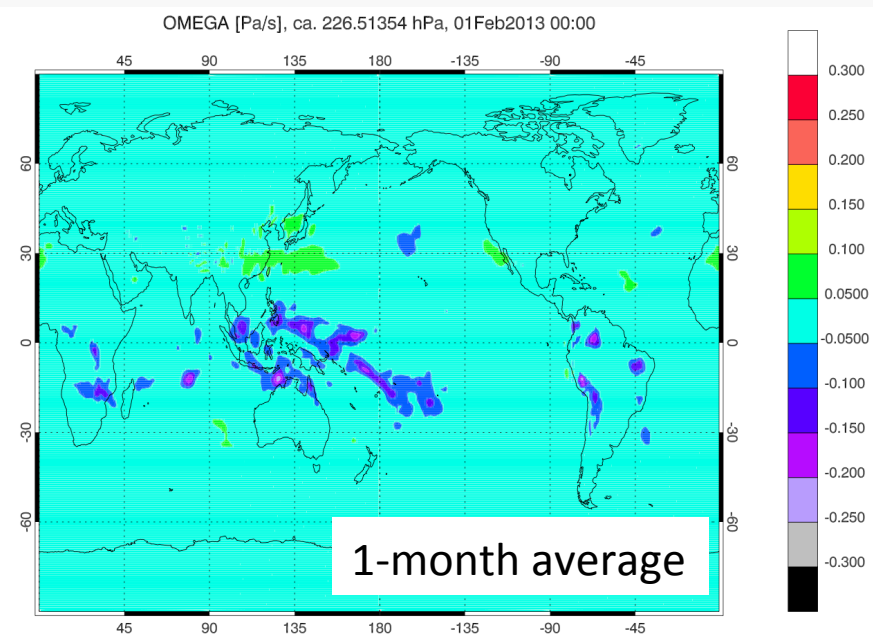
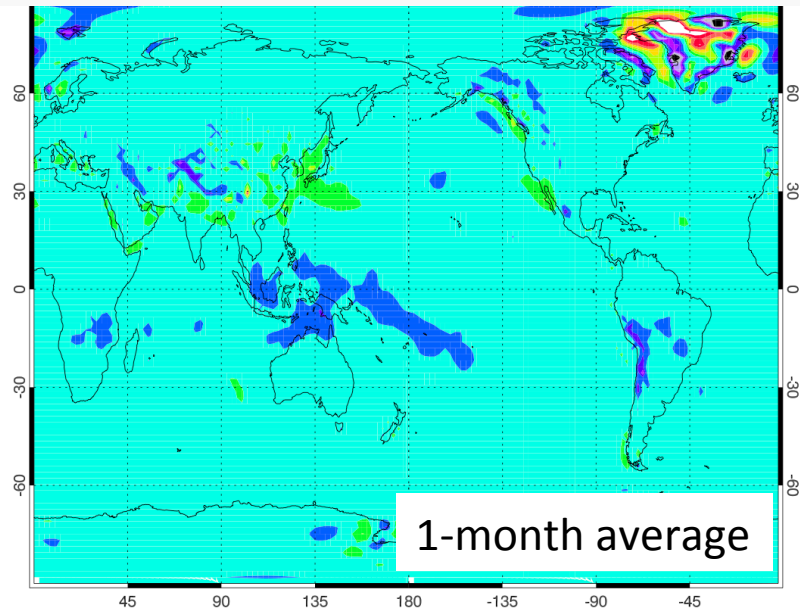
SE: NE30



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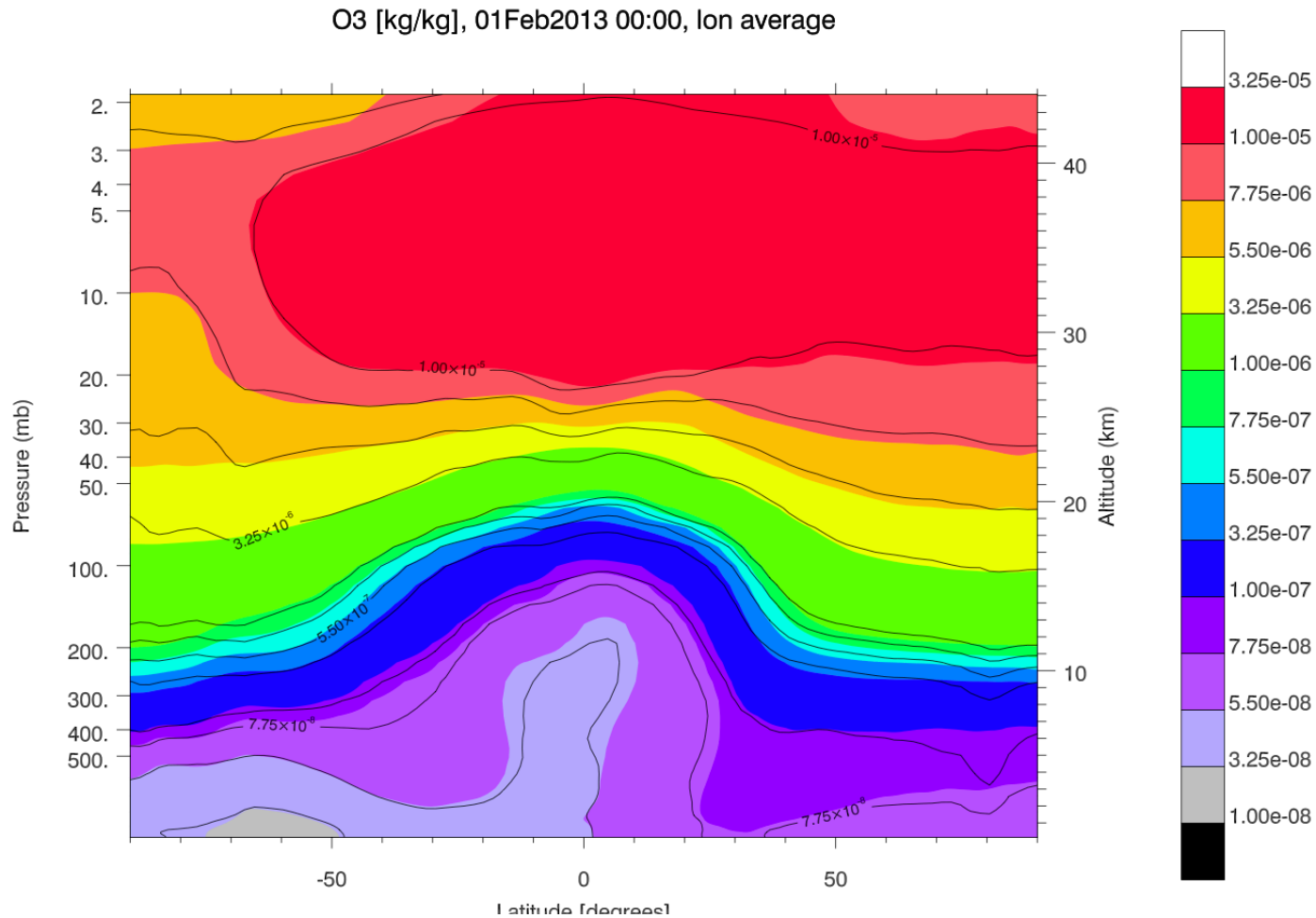
aconley 07.02.2



/glade/scratch/fvtt/archive/fcam4\_f19\_geos\_nudg01/atm/hist/fcam4\_f19\_geos\_nudg01.cam.h0.2013-01.nc

aconley 06.02.2014 17:44

# Filled contours: CAM-SE Black contours: CAM-FV



Logarithmic contour scale: factor of 3 too much mixing over Greenland/Equator

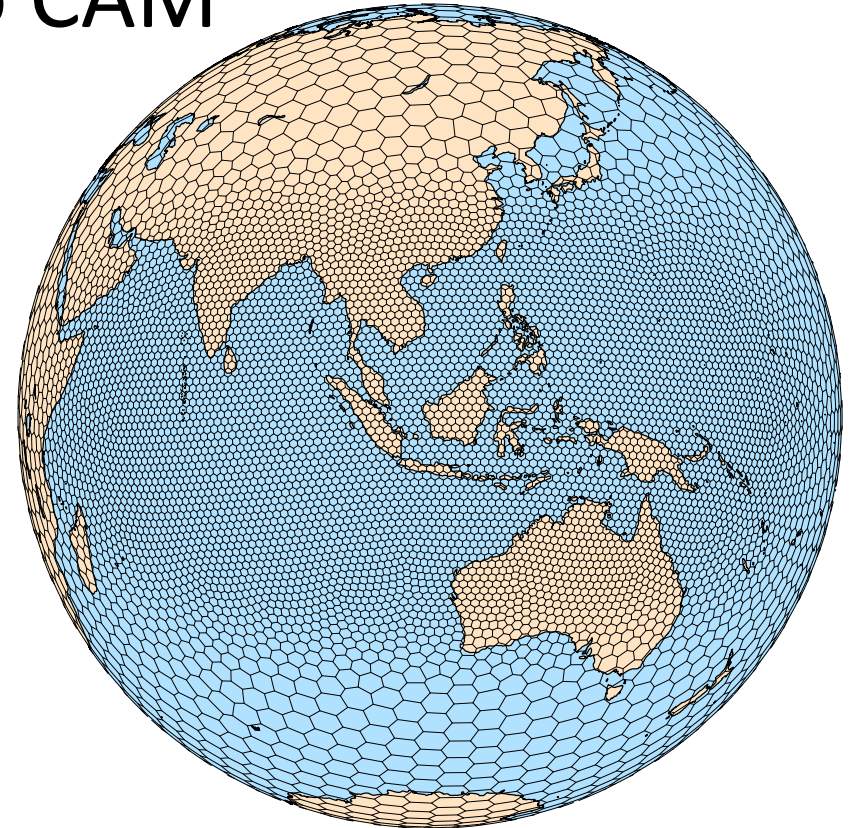
# Nonhydrostatic MPAS-Atmosphere dynamical core port to CAM

- Software engineering of port is complete; uses CAM5 physics.
- Held-Suarez and APE testing is underway.
- AMIP testing this spring.
- NWP testing in hydrostatic regime later this winter.

*Major concern: Scale-aware physics, physics (deep convection, microphysics) for nonhydrostatic resolutions in CAM*

*W.C. Skamarock, S.-H. Park, P.H. Lauritzen*

**MPAS**  
Model for Prediction Across Scales



Variable-resolution MPAS mesh,  
refinement over the Maritime Continent  
region





# Axial angular momentum

In the absence of any surface torque and zonal mechanical forcing, the hydrostatic primitive equations conserve the globally integrated AAM when assuming a constant pressure upper boundary [see, e.g., *Staniforth and Wood, 2003*]:

$$\frac{dM}{dt} = 0. \quad (2)$$

Typically numerical models are divided into a dynamical core (*dyn*) that, roughly speaking, solves the equations of motion on resolved scales and physical parameterizations that approximate sub-grid-scale processes (*phys*). There can therefore be two sources/sinks of AAM:

$$\frac{dM}{dt} = \left( \frac{dM}{dt} \right)_{dyn} + \left( \frac{dM}{dt} \right)_{phys}. \quad (3)$$

In the absence of mountain torque:

$$0 \sim \left( \frac{dM}{dt} \right)_{dyn} \ll \left( \frac{dM}{dt} \right)_{phys}.$$





# A simple way to assess axial angular momentum conservation

Held-Suarez forcing: flat-Earth (no mountain torque), physics replaced by simple boundary layer friction and relaxation of temperature toward reference profile

$$\frac{\partial v}{\partial t} = \dots - k_v(\sigma)v$$

$$\frac{\partial T}{\partial t} = \dots - k_T(\phi, \sigma)[T - T_{eq}(\phi, \rho)]$$

$$T_{eq} = \max \left\{ 200K, \left[ 315K - (\Delta T)_y \sin^2 \phi - (\Delta \theta)_z \log \left( \frac{\rho}{\rho_0} \right) \cos^2 \phi \right] \left( \frac{\rho}{\rho_0} \right)^\kappa \right\}$$

$$k_T = k_a + (k_s - k_a) \max \left( 0, \frac{\sigma - \sigma_b}{1 - \sigma_b} \right) \cos^4 \phi$$

$$k_v = k_f \max \left( 0, \frac{\sigma - \sigma_b}{1 - \sigma_b} \right)$$

$\sigma_b = 0.7$	$k_f = 1 \text{ day}^{-1}$ ,
$k_a = 1/40 \text{ day}^{-1}$	$k_s = 1/4 \text{ day}^{-1}$
$(\Delta T)_y = 60K$	$(\Delta \theta)_z = 10K$

$\rho_0 = 1000 \text{ mb}$	$\kappa = \frac{R}{c_p} = \frac{2}{7}$	$c_p = 1004 \text{ J kg}^{-1} \text{ K}^{-1}$
$\Omega = 7.292 \times 10^{-5} \text{ s}^{-1}$	$g = 9.8 \text{ m s}^{-2}$	$a_g = 6.371 \times 10^6 \text{ m}$

