



SciDAC
Scientific Discovery through
Advanced Computing



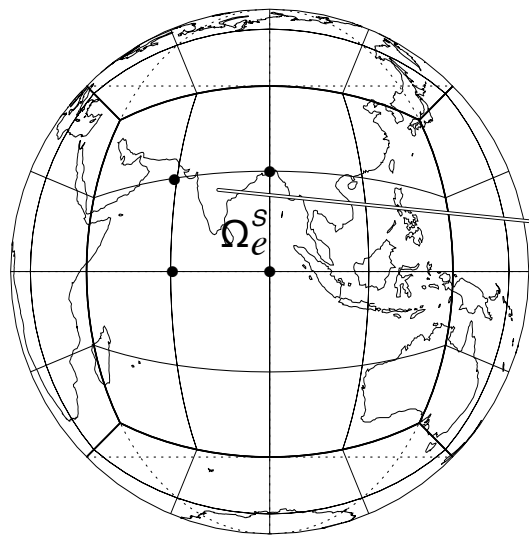
Conservation and coupling in CAM-SE (CAM=Community Atmosphere Model; SE=Spectral Elements)

Peter Hjort Lauritzen

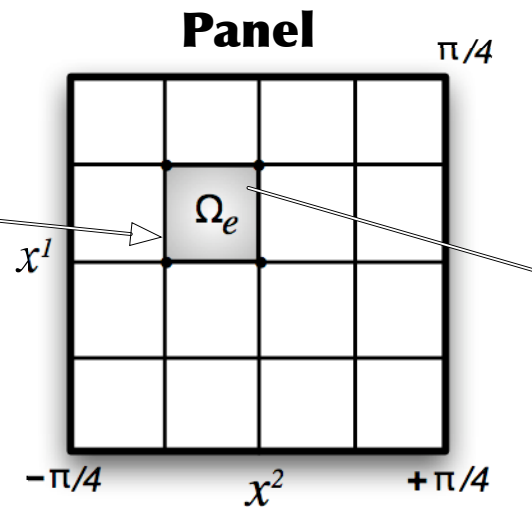
**Atmospheric Modeling and Predictability Section
Climate and Global Dynamics Laboratory
National Center for Atmospheric Research**

**Workshop on Physics Dynamics Coupling in Weather and Climate Models
September 20-22, 2016
Pacific Northwest National Laboratory Richland, WA, USA**

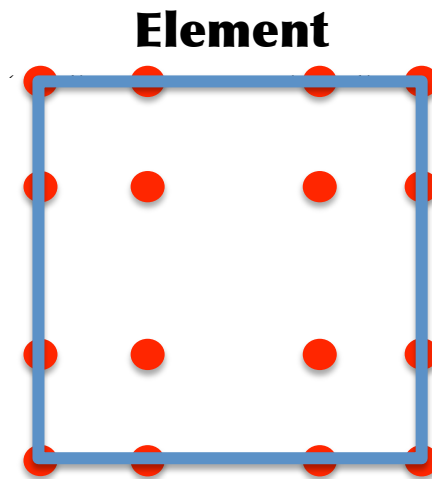
The spectral-element (SE) method: discretization grid



Physical Domain

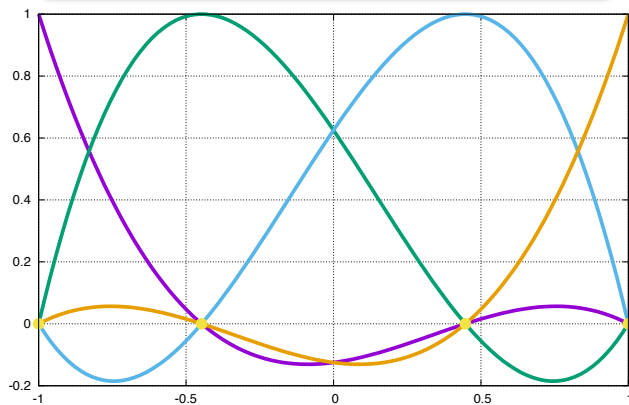


Computational Domain



GLL Quadrature Grid

Nodal 1D polynomial basis functions



GLL=Gauss-Lobatto-Legendre

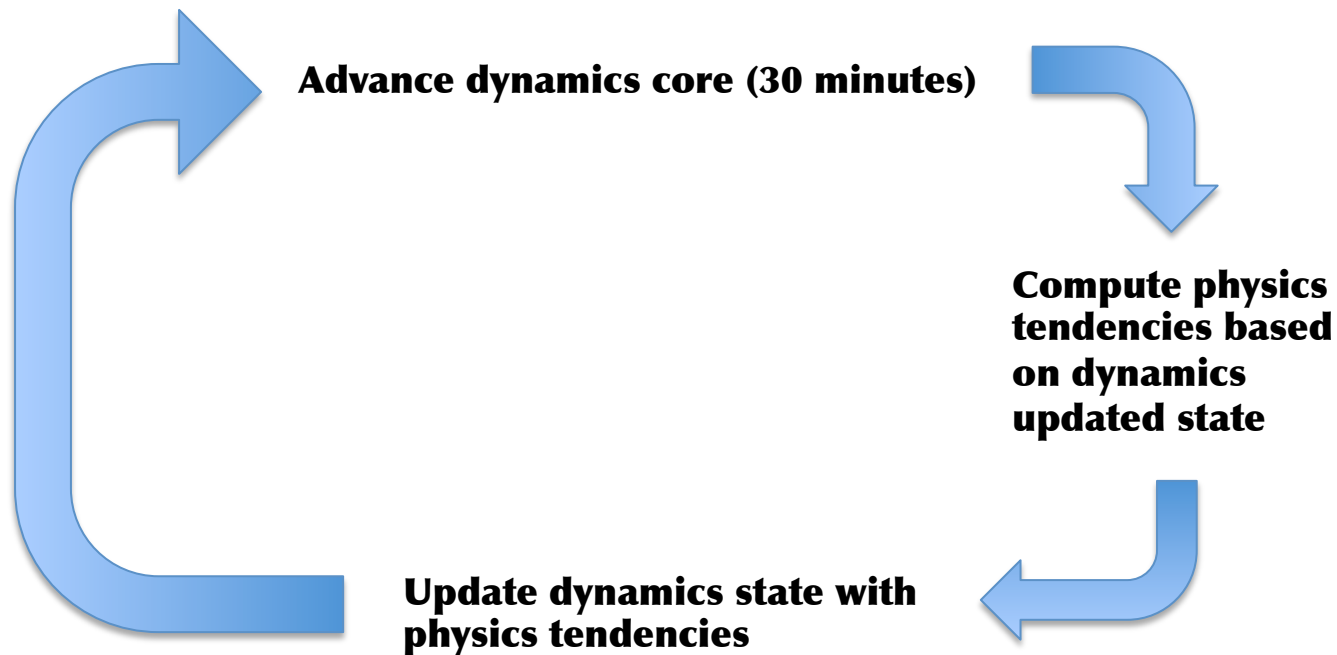
Conservation and coupling

- **Mass-conservation in dynamical core:** In the absence of sources and sinks, dry air mass and tracer mass should be conserved
- **Mass-conservation in physics-dynamics coupling:** When adding physics tendencies to tracers in the dynamical core, the mass budget should be closed (not necessarily strictly enforced!)
- **Closed energy budget:** Atmospheric component as a whole should have a closed energy budget (no spurious sources and sinks)
- **Closed energy budget in physics-dynamics coupling:** When adding physics tendencies to the state, the energy change in tendencies should match the energy change in the atmospheric state when tendencies have been added (not necessarily strictly enforced!)

Conservation and coupling: default CAM-SE

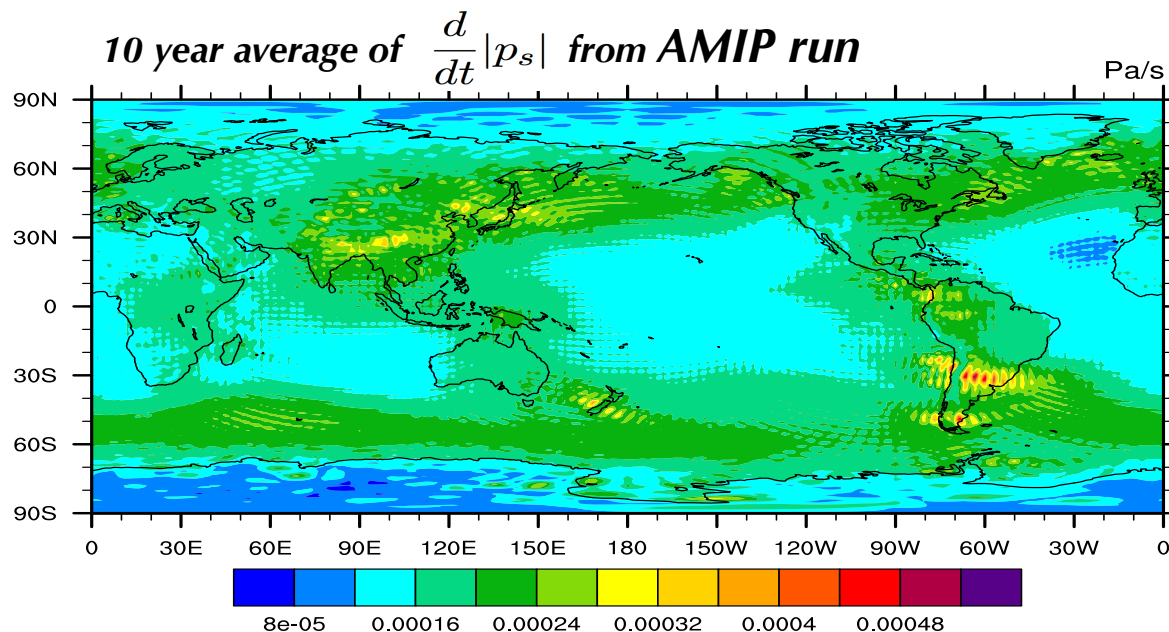
- **Mass-conservation in dynamical core: YES!**
- **Mass-conservation in physics-dynamics coupling: YES if “ftype=1”**

1. Physics updates state (“ftype=1”): mass budget closed but ...



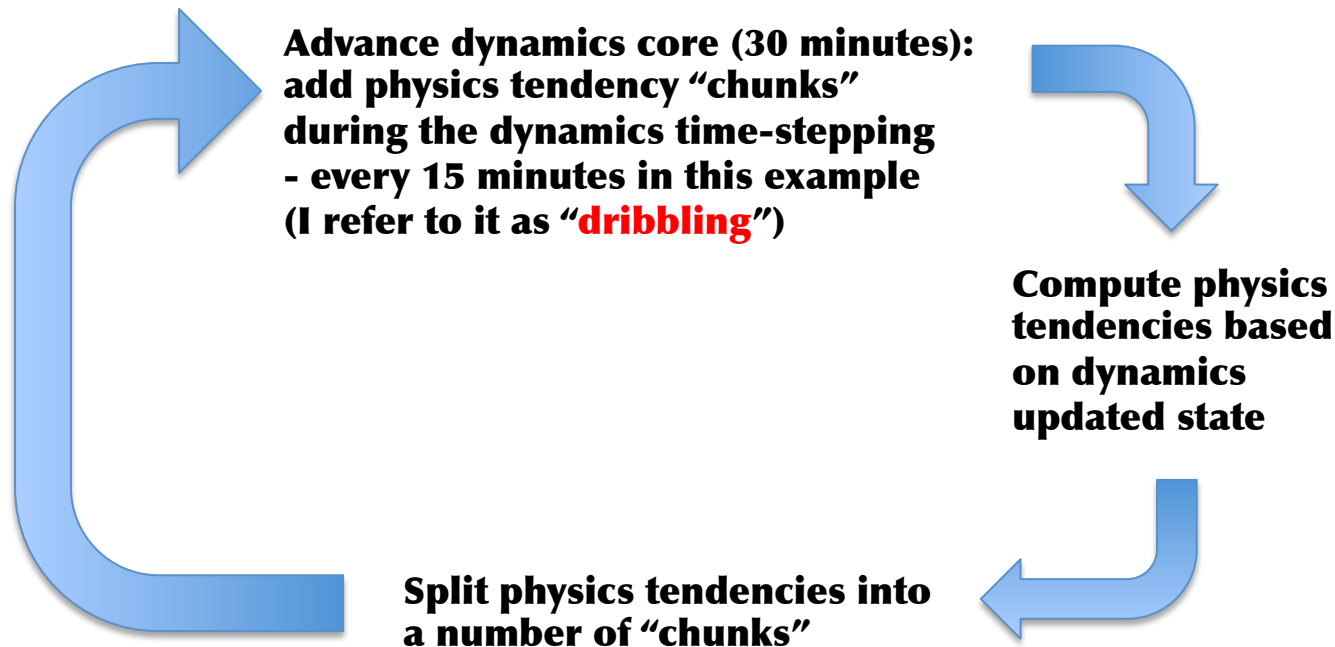
Conservation and coupling: default CAM-SE

- **Mass-conservation in dynamical core: YES!**
- **Mass-conservation in physics-dynamics coupling: YES if “ftype=1”**
- **1. Physics updates state (“ftype=1”): mass budget closed but ...**



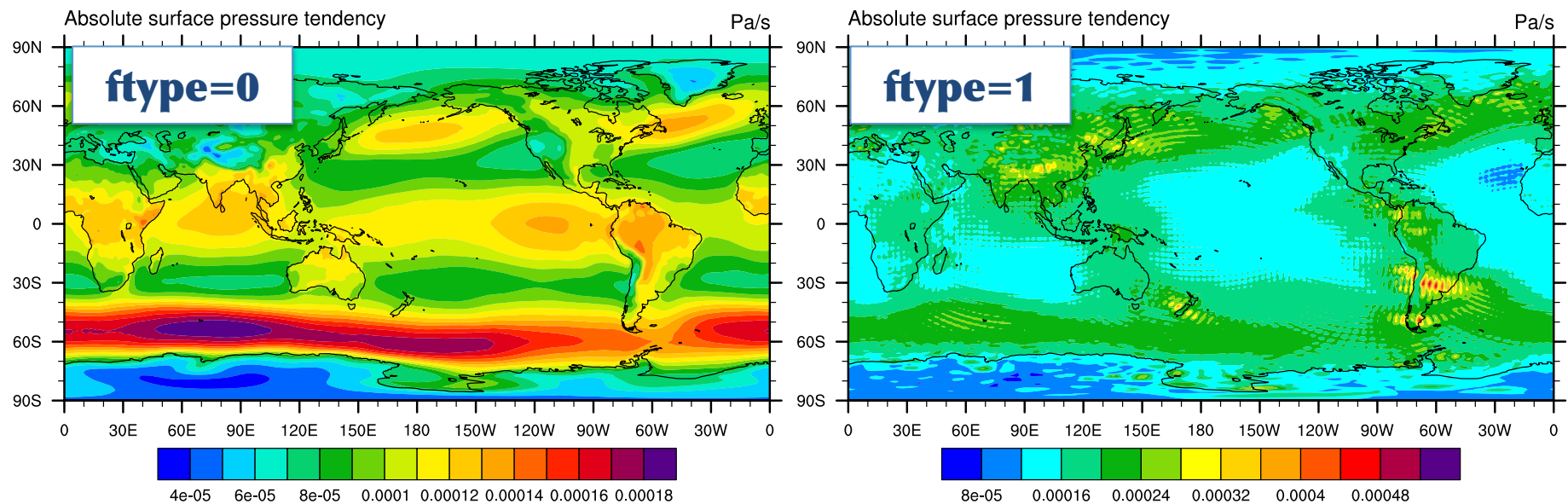
Conservation and coupling: default CAM-SE

- **Mass-conservation in dynamical core: YES!**
- **Mass-conservation in physics-dynamics coupling:**



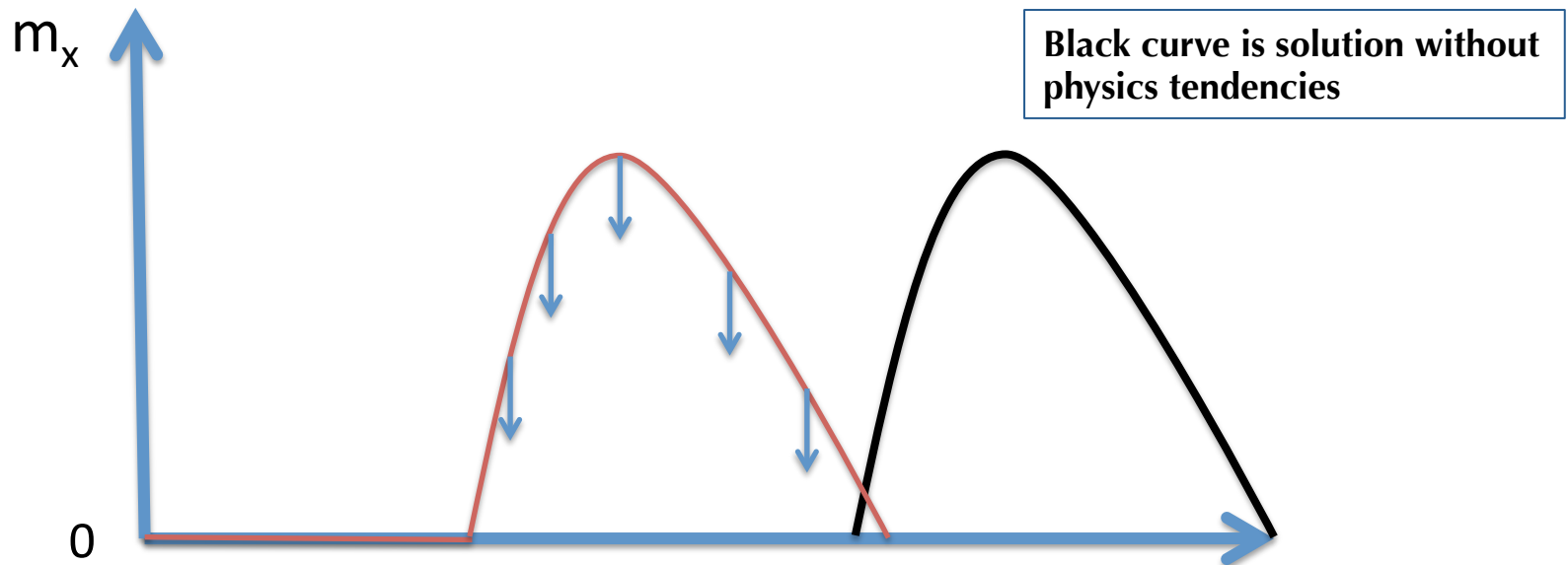
Conservation and coupling: default CAM-SE

- **Mass-conservation in dynamical core: YES!**
- **Mass-conservation in physics-dynamics coupling:**



Conservation and coupling: default CAM-SE

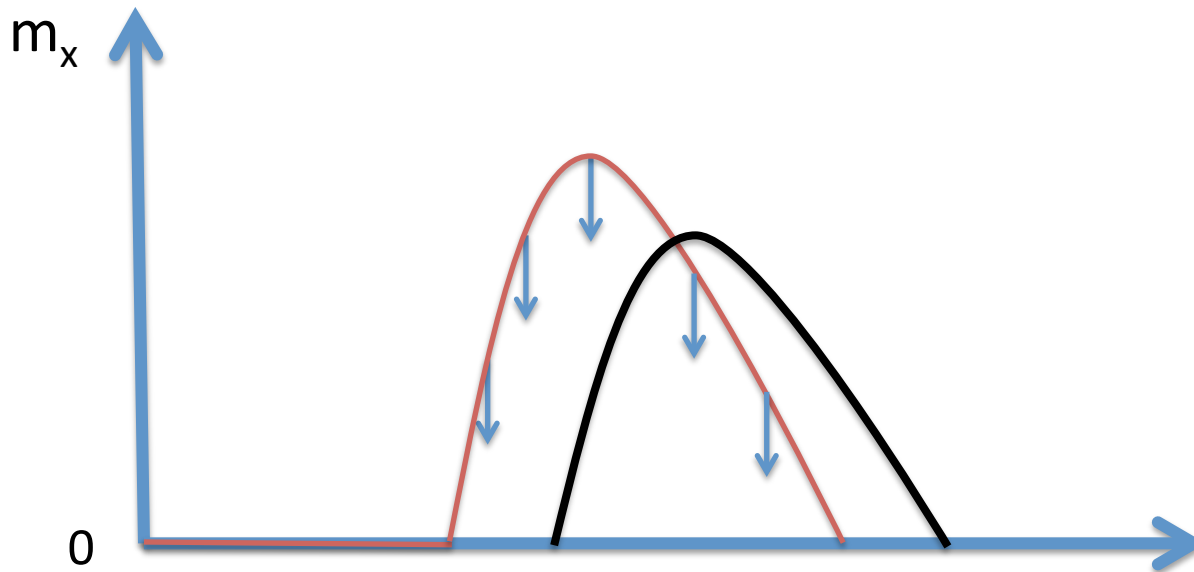
- **Mass-conservation in dynamical core: YES!**
- **Mass-conservation in physics-dynamics coupling: No!**
- **2. “Dribbling” (“ftype=0”): mass budget not always closed!**



Conservation and coupling: default CAM-SE

- **Mass-conservation in dynamical core: YES!**
- **Mass-conservation in physics-dynamics coupling: No!**

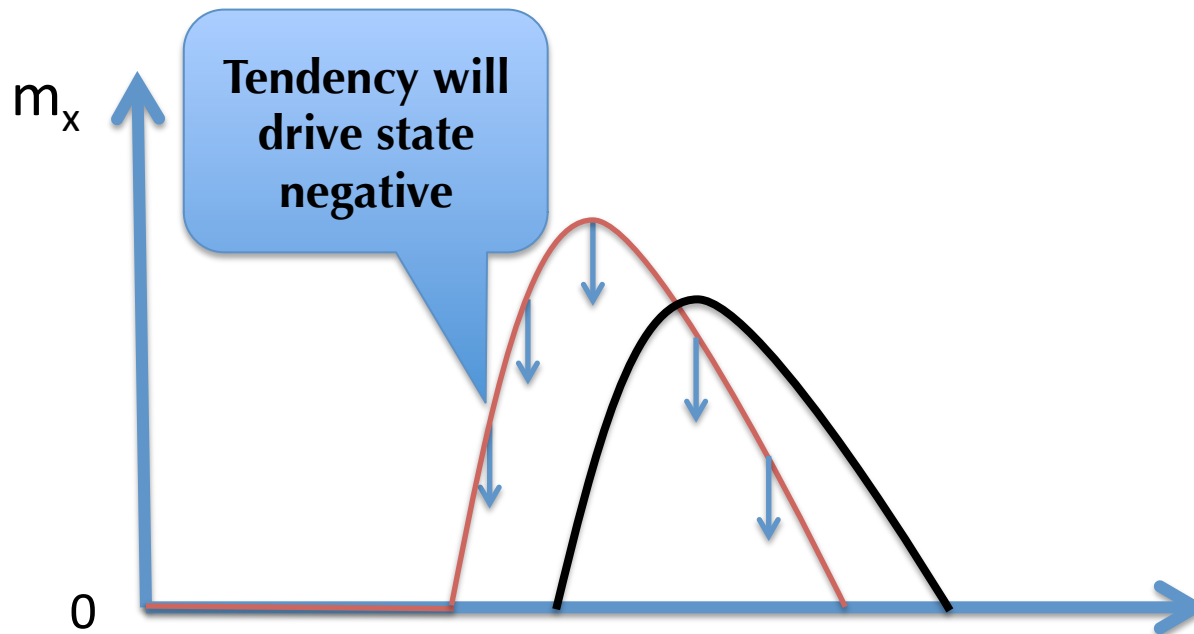
2. “Dribbling” (“ftype=0”): mass budget not always closed!



Conservation and coupling: default CAM-SE

- **Mass-conservation in dynamical core: YES!**
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2. “Dribbling” (“ftype=0”): mass budget not always closed!

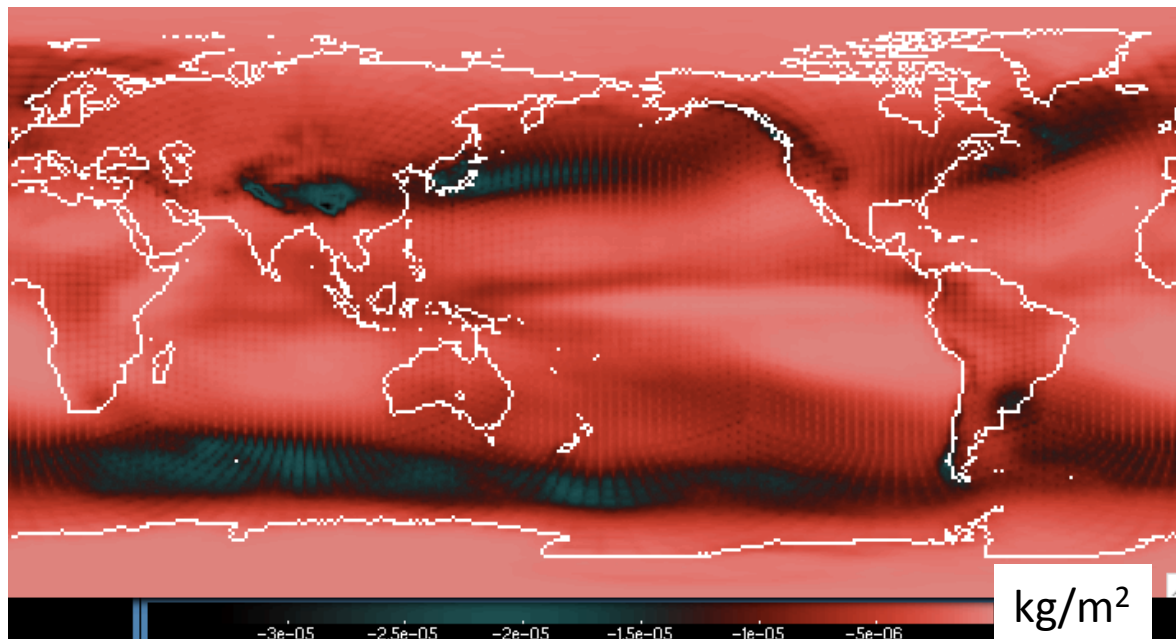


Conservation and coupling: default CAM-SE

- **Mass-conservation in dynamical core: YES!**
- **Mass-conservation in physics-dynamics coupling: No!**

2. “Dribbling” (“ftype=0”): mass budget not always closed!

10-year average of column integrated cloud ice (CLDICE) gained in physics-dynamics coupling



$O(5E-7hPa)$ vertically integrated CLDICE gained in PDC

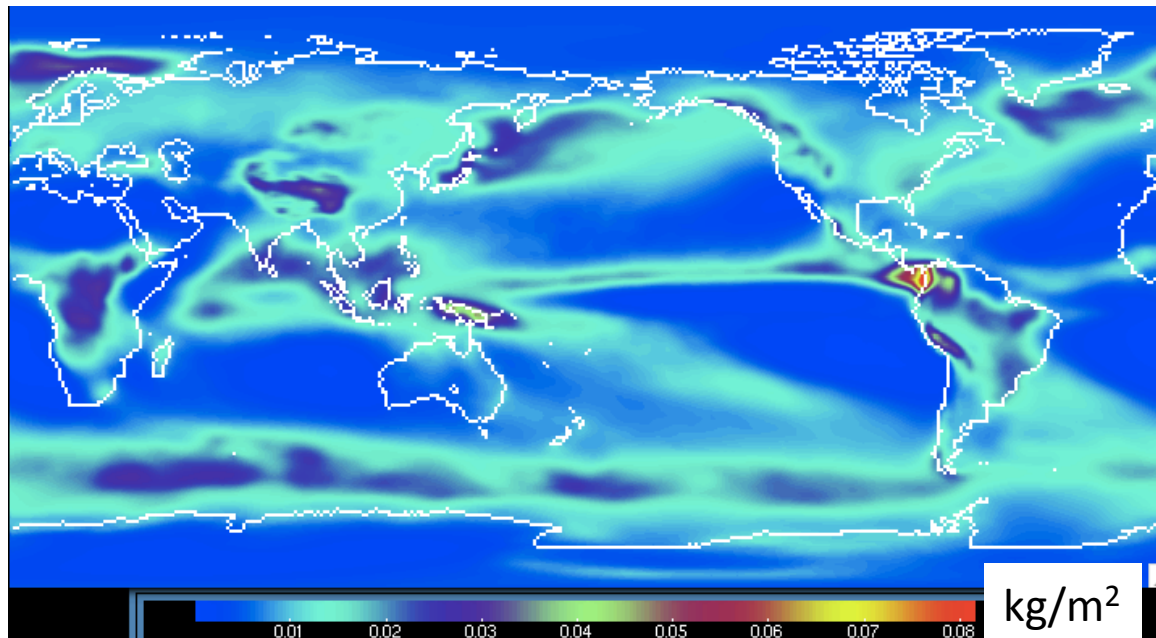
=> tendencies added every 15 minutes so in a 10 year simulation tendencies are added $O(3.5E5)$ times
=> accumulated CLDICE gained in PDC is $O(0.175hPa)$.

Conservation and coupling: default CAM-SE

- **Mass-conservation in dynamical core: YES!**
- **Mass-conservation in physics-dynamics coupling: No!**

2. “Dribbling” (“ftype=0”): mass budget not always closed!

10-year average of column integrated cloud ice (CLDICE)



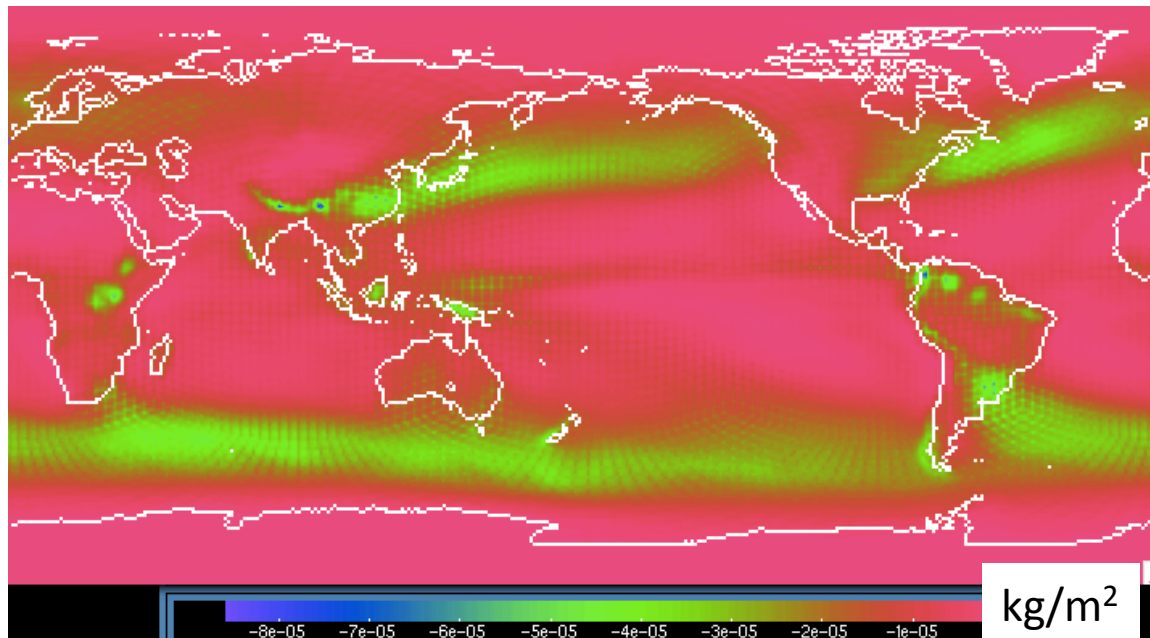
Global integral of CLDICE corresponds to approximately $1\text{E-}3\text{hPa}$

Conservation and coupling: default CAM-SE

- **Mass-conservation in dynamical core: YES!**
- **Mass-conservation in physics-dynamics coupling: No!**

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10-year average of column integrated cloud liquid (CLDLIQ) gained in physics-dynamics coupling



O(1E-6hPa) vertically integrated CLDLIQ lost in PDC

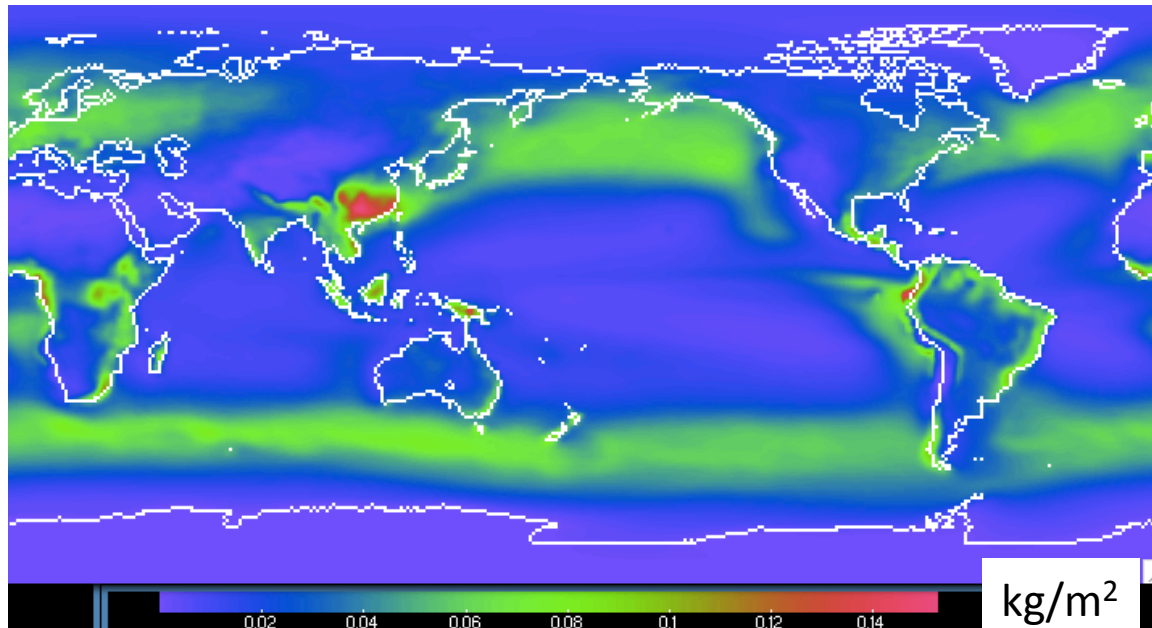
=> tendencies added every 15 minutes so in a 10 year simulation tendencies are added O(3.5E5) times
=> accumulated CLDLIQ gained in PDC is O(0.35hPa).

Conservation and coupling: default CAM-SE

- **Mass-conservation in dynamical core: YES!**
- **Mass-conservation in physics-dynamics coupling: No!**

2. “Dribbling” (“ftype=0”): mass budget not always closed!

10-year average of column integrated cloud liquid (CLDLIQ)



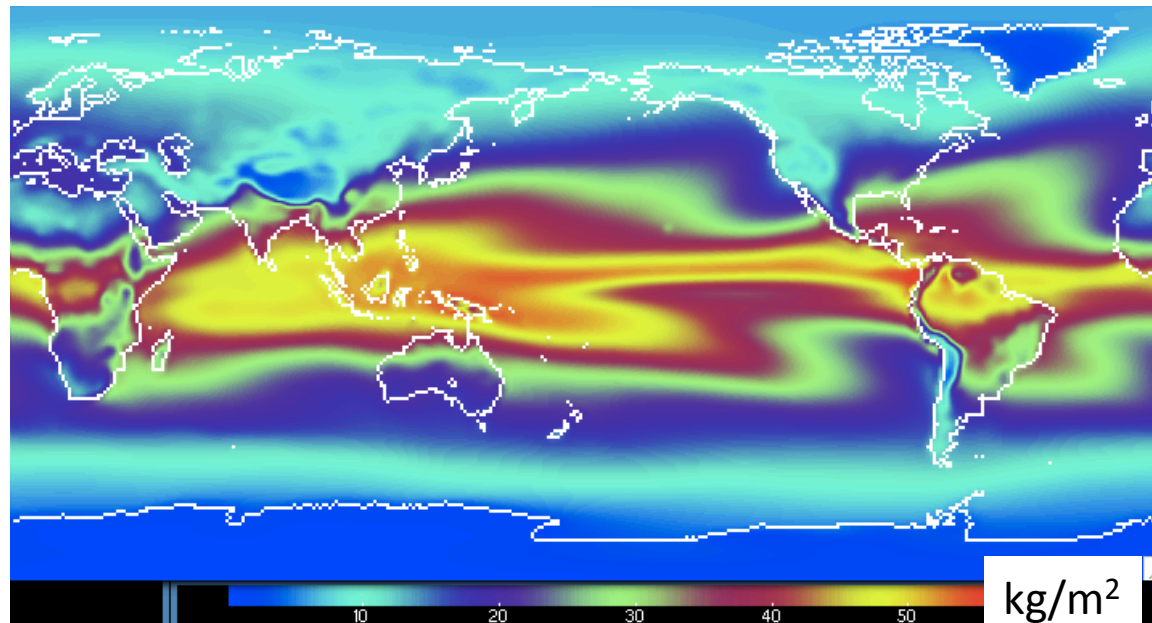
Global integral of CLDLIQ corresponds to approximately $2.6E-4$ hPa

Conservation and coupling: default CAM-SE

- **Mass-conservation in dynamical core: YES!**
- **Mass-conservation in physics-dynamics coupling: No!**

2. “Dribbling” (“ftype=0”): mass budget not always closed!

10-year average of column integrated water vapor (WV)



$O(1E-8hPa)$ vertically integrated CLDLIQ gained in PDC

=> tendencies added every 15 minutes so in a 10 year simulation tendencies are added $O(3.5E5)$ times
=> accumulated WV gained in PDC is $O(1E-3hPa)$.

Conservation and coupling: default CAM-SE

- **Mass-conservation in dynamical core: YES!**
- **Mass-conservation in physics-dynamics coupling: No!**

2. “Dribbling” (“ftype=0”): mass budget not always closed!

Solution(s):

- **Advect tendencies (expensive....!)**
- **Other?**

Conservation and coupling: CAM-SE

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(not necessarily strictly enforced!)

Energy Conservation and coupling in CAM-SE

10 year averages from AMIP simulation (specified SSTs cycling over same year)

Dynamical core module

- Rate of energy change due to explicit dissipation (hyperviscosity)

$$dE/dt = 0.0729 \text{ W/m}^2$$

- Frictional heating rate is calculated from K tendency produced from momentum diffusion and added to T:

$$dE/dt = 0.6997 \text{ W/m}^2$$

- Vertical remapping

$$dE/dt = -0.1547 \text{ W/m}^2$$

Total loss of energy in dynamics

$$dE/dt = -0.0723 \text{ W/m}^2$$

Rate of energy change due to “dribbling” physics tendencies in the dynamics

$$dE/dt = 0.056 \text{ W/m}^2$$

Physics module

- “physical” changes in energy due to water change

$$dE/dt = -0.0016 \text{ W/m}^2$$

- Change in energy due to change in pressure due to water vapor change (“dme_adjust”)

$$dE/dt = 0.2667 \text{ W/m}^2$$

- Energy fixer

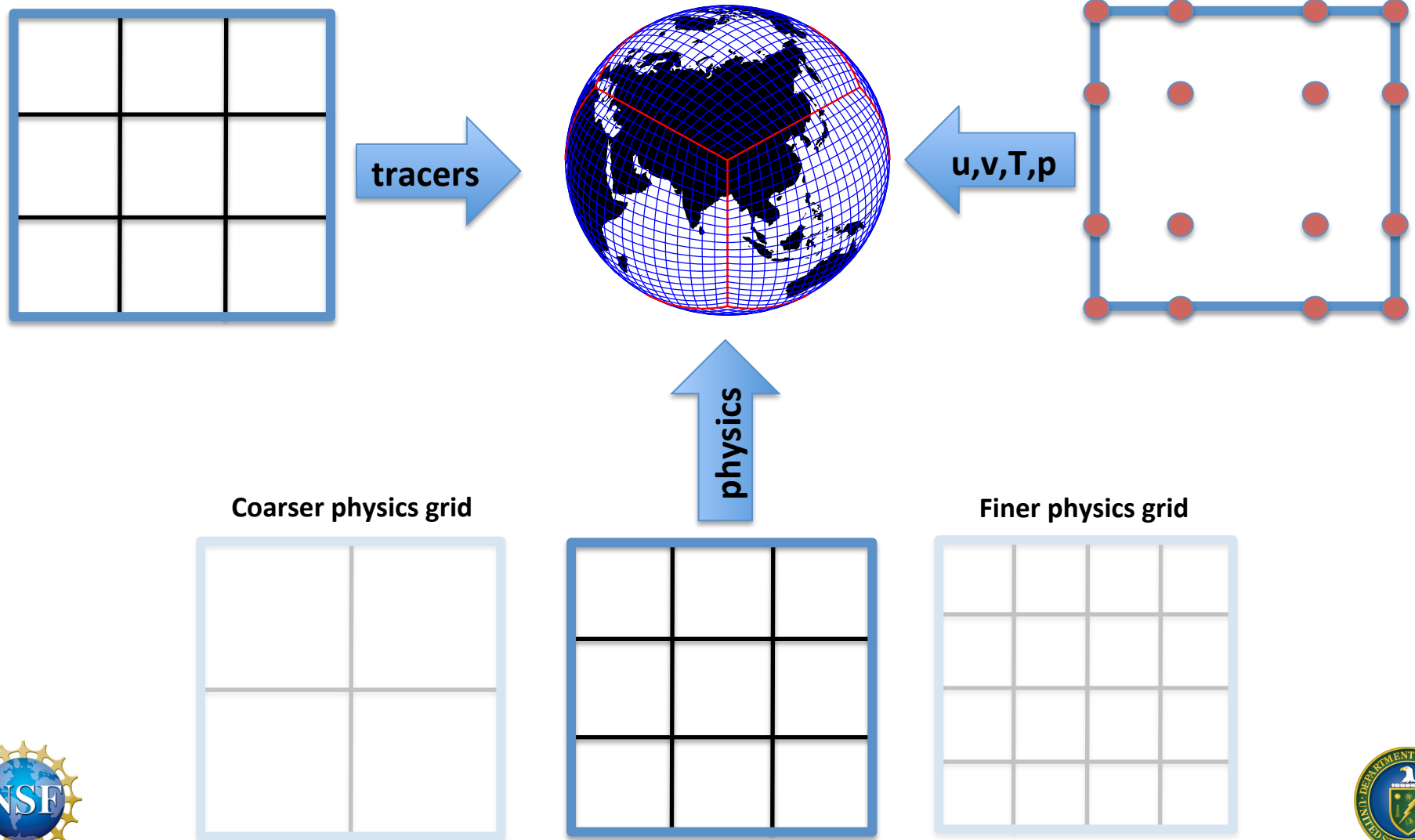
$$dE/dt = -0.1843$$

(= loss in dynamics + dme_adjust)

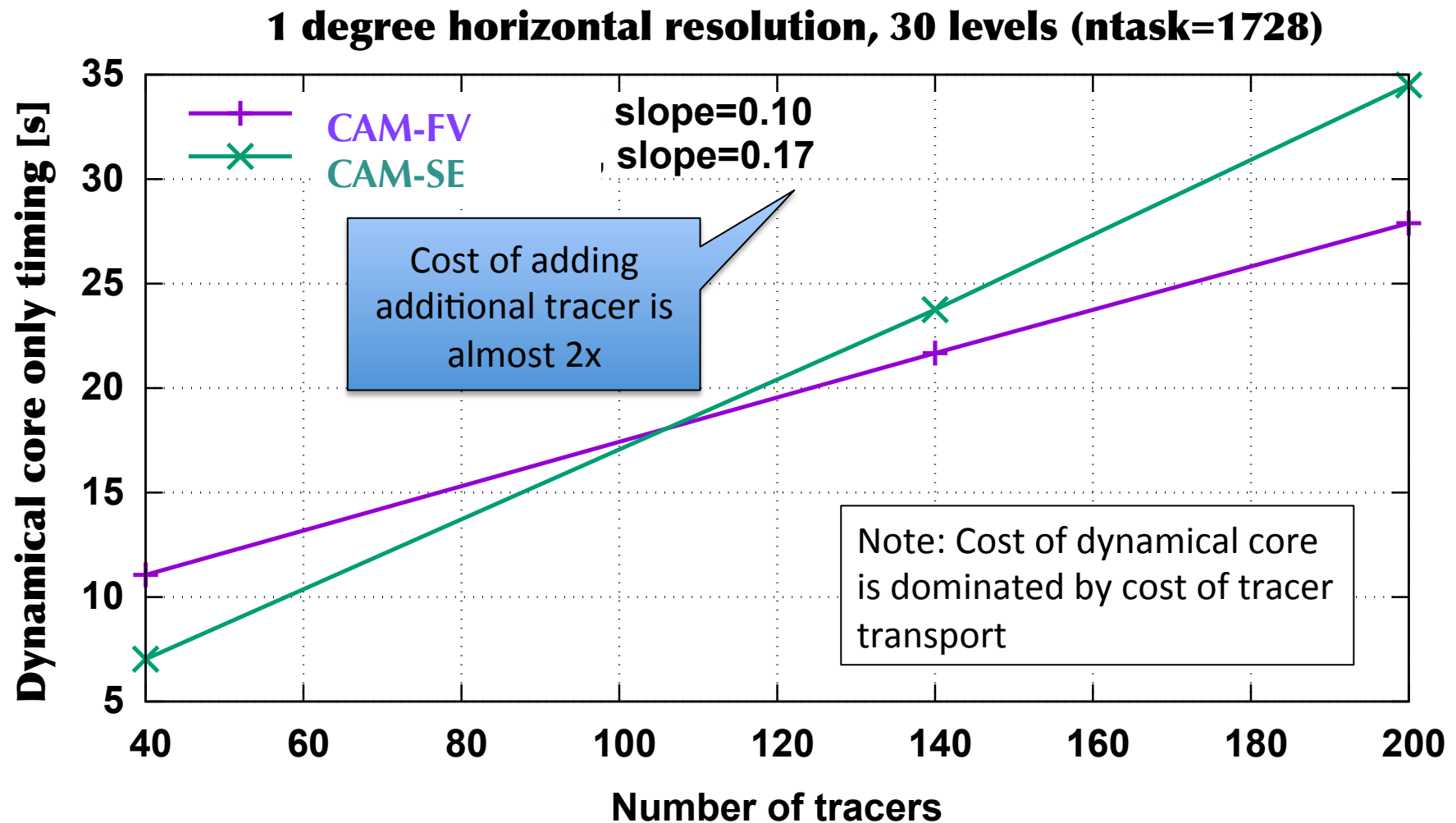
Physics-dynamics coupling layer



Separating dynamics, tracer and physics grids in CAM-SE



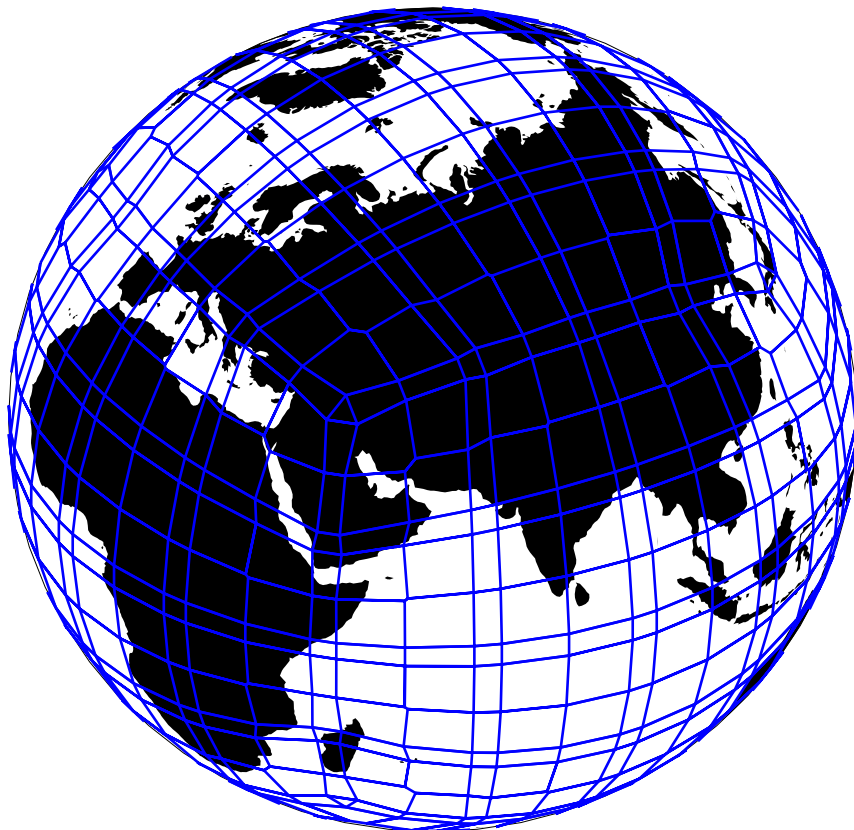
1. Why separate tracer grid? SE tracer transport cost



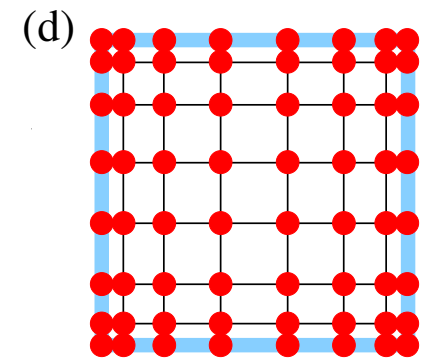
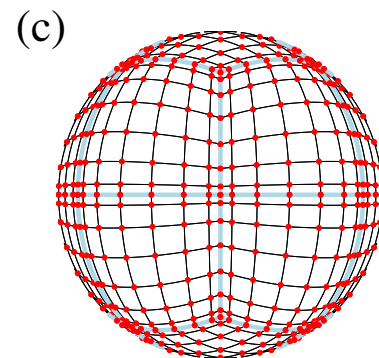
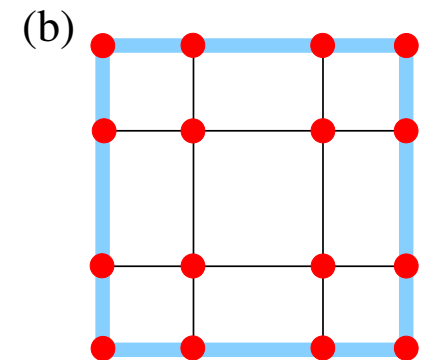
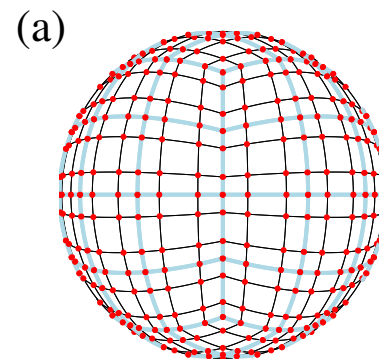
2. Why separate physics grid?

Non-uniform sampling of atmospheric state

Current physics/“coupler” grid

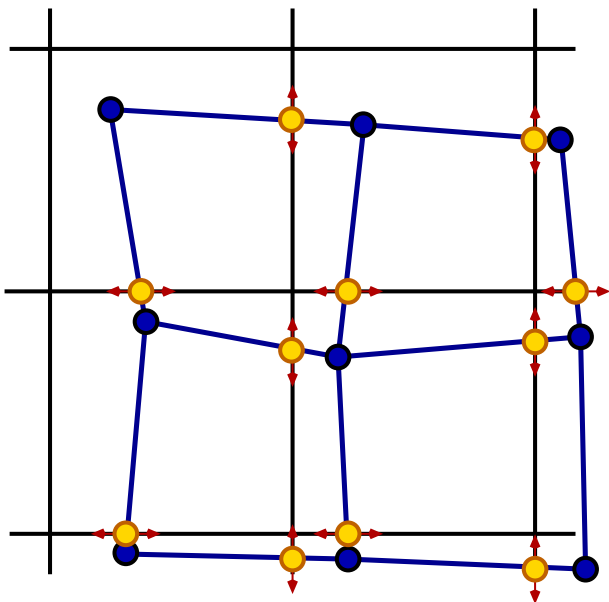
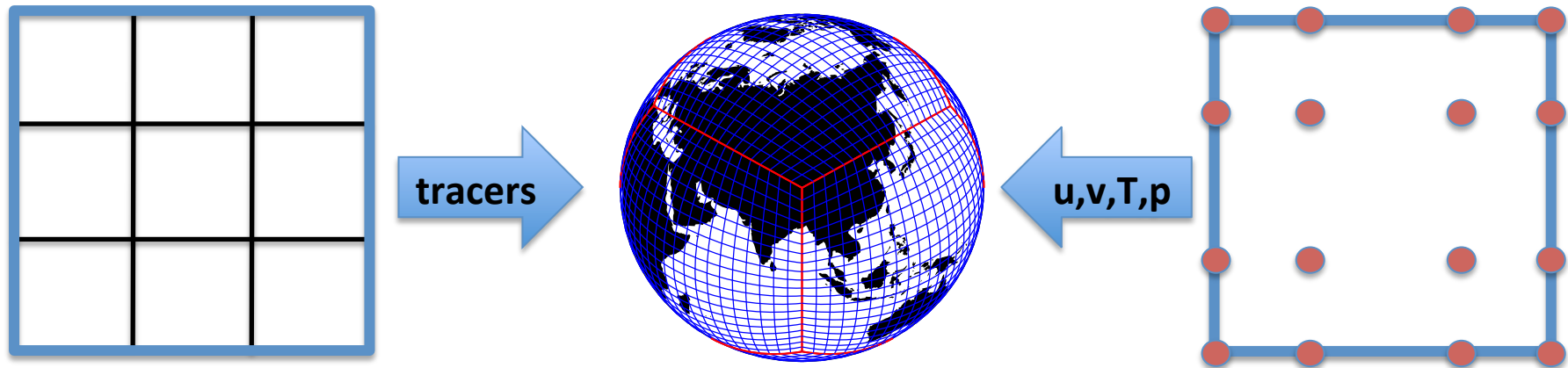


Gets worse with increasing order!





Separating transport and dynamics grids/ methods in CAM-SE



Coupling cell-integrated semi-Lagrangian scheme (CSLAM) with spectral-element dynamics

- Find upstream Lagrangian grid so that air mass-change in CSLAM cell exactly matches air mass change in the same cell but by integrating SE basis functions (see Lauritzen et al., 2016 for details)
- Inherent mass-conservation, consistency, shape-preservation as described in Lauritzen et al., (2010)

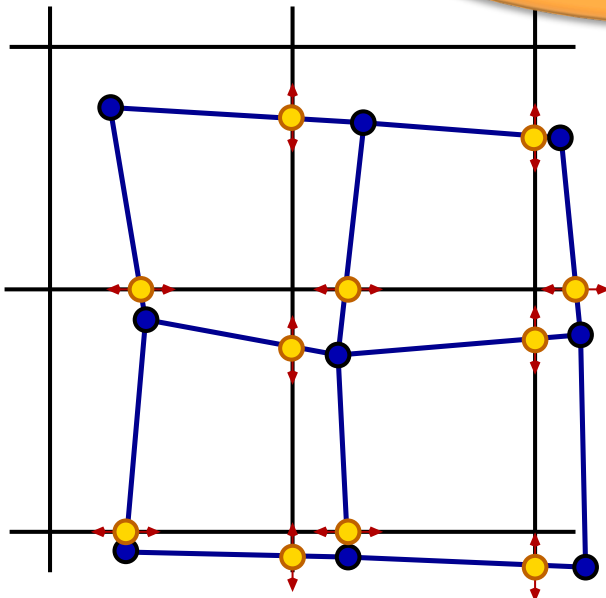
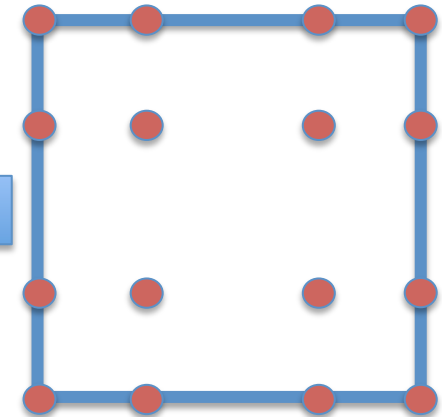


Separating transport and dynamics grids/ methods in CAM-SE



Note: coupling algorithm is general - it can be used with any dynamical core that can provide fluxes for CSLAM control volumes!

v, T, p



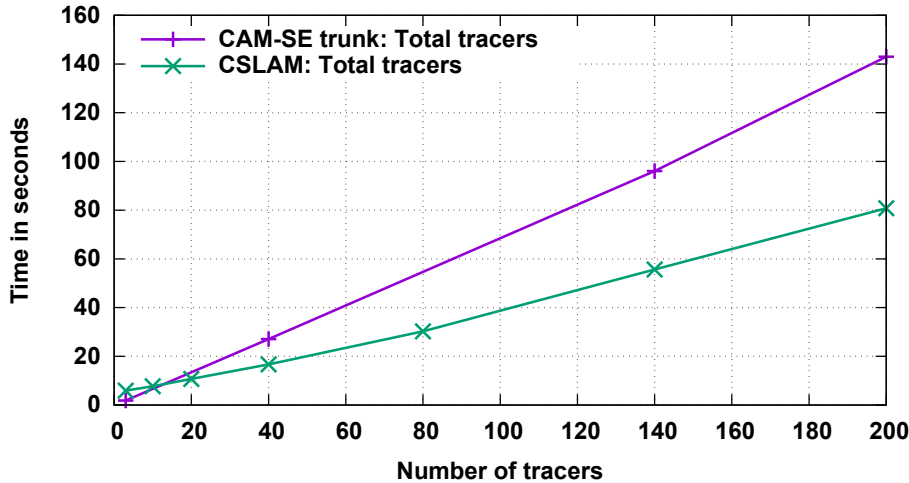
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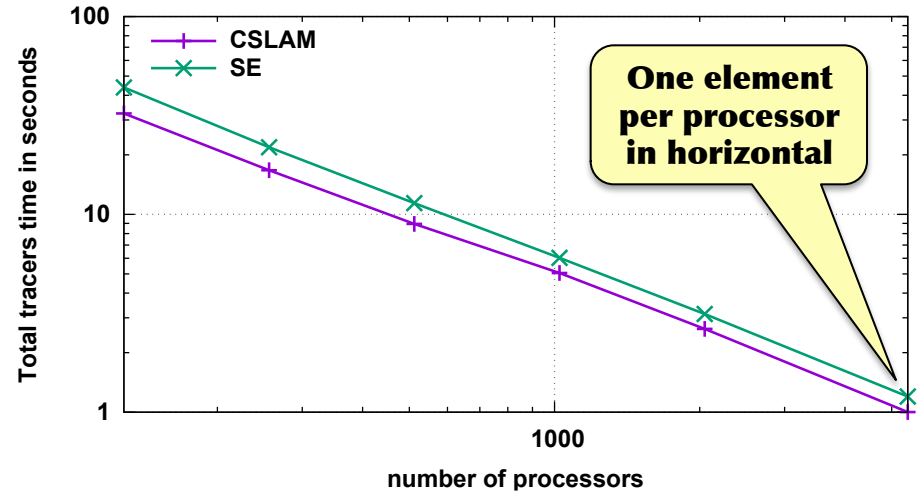


CSLAM transport is (a) faster than, (b) as scalable as and (c) more accurate than CAM-SE transport

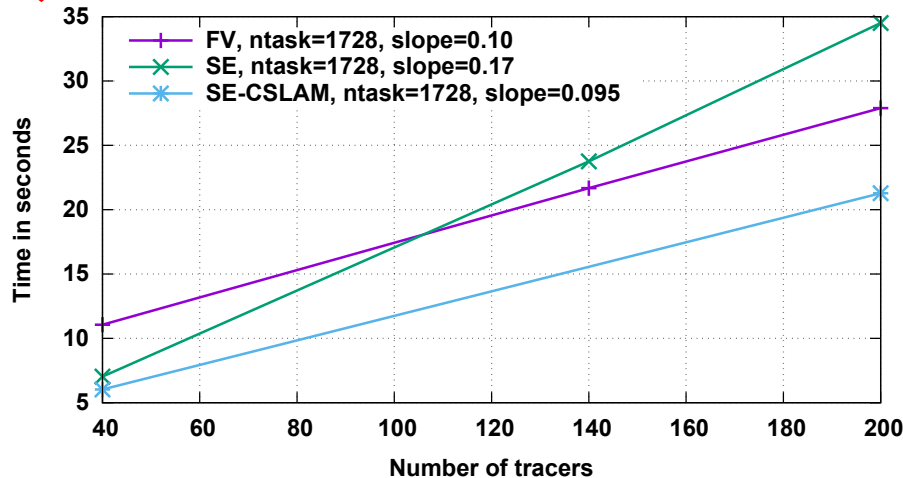
(a) 1 degree horizontal resolution, 30 levels, 256 tasks



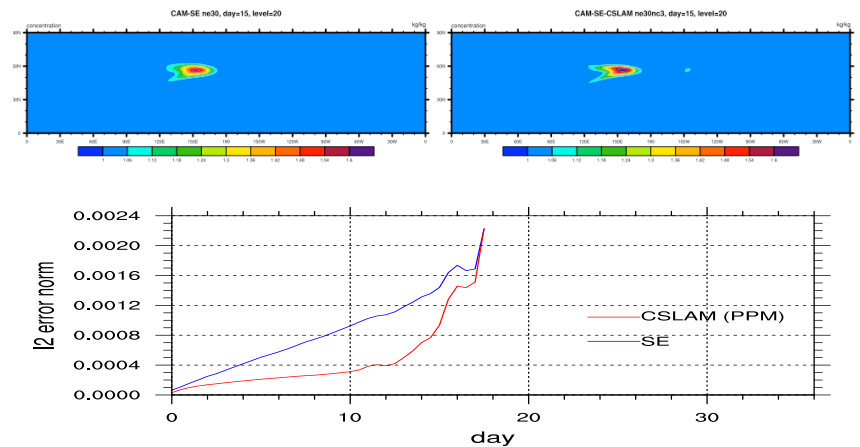
(b) 1 degree horizontal resolution, 30 levels, 40 tracers



(a) 1 degree horizontal resolution, 30 levels, 1728 tasks

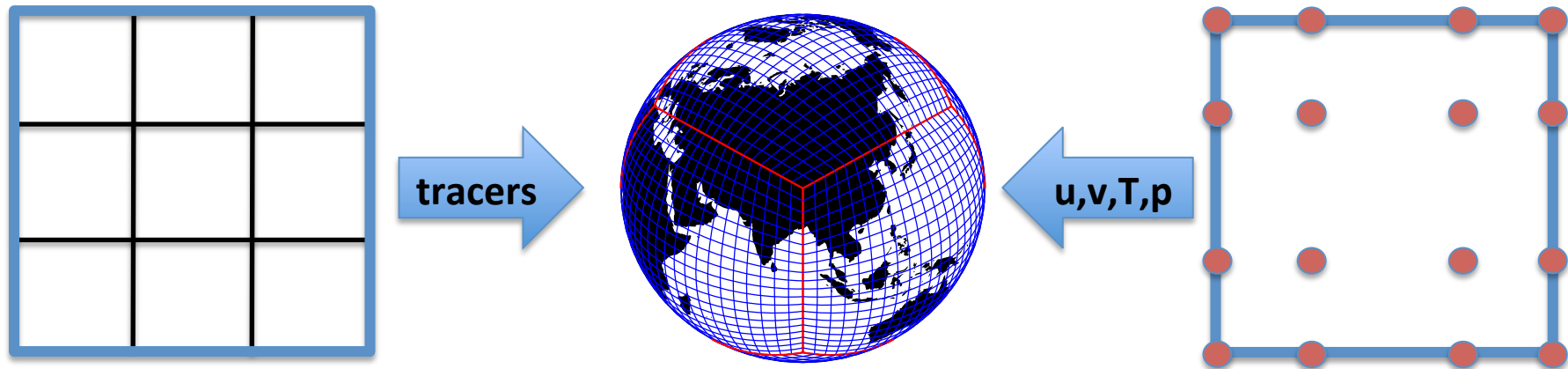


(c) Inert & passive tracer in dry baroclinic wave flow





Separating transport and dynamics grids/ methods in CAM-SE **with moisture**



Most global models used a moist pressure vertical coordinate

$$p(\eta) = A(\eta)p_0 + B(\eta)p_s \cdot$$

=> If there is a change in water vapor then levels move (problematic when coupling CSLAM and SE)

Solution: dry mass vertical coordinates (also makes it straight forward to add water loading in dynamical core)

Dry mass vertical coordinates

1.4. **Primitive equations.** The η_d -coordinate atmospheric primitive equations assuming floating Lagrangian vertical coordinates, neglecting dissipation and forcing terms can be written as

$$(17) \quad \frac{\partial \vec{u}}{\partial t} + (\zeta + f) \hat{k} \times \vec{u} + \nabla \left(\frac{1}{2} \vec{u}^2 + \Phi \right) + \frac{1}{\rho} \nabla p = 0$$

$$(18) \quad \frac{\partial T}{\partial t} + \vec{u} \cdot \nabla T - \frac{1}{c_p \rho} \omega = 0$$

$$(19) \quad \frac{\partial}{\partial t} \left(\frac{\partial p_d}{\partial \eta} \right) + \nabla \cdot \left(\frac{\partial p_d}{\partial \eta} \vec{u} \right) = 0$$

$$(20) \quad \frac{\partial}{\partial t} \left(\frac{\partial p_d}{\partial \eta} m_i \right) + \nabla \cdot \left(\frac{\partial p_d}{\partial \eta} m_i \vec{u} \right) = F_i, \quad i = v, cl, ci, \dots$$

Let m_i be the amount of water substance, chemical species, etc. of type i associated with unit mass of dry air:

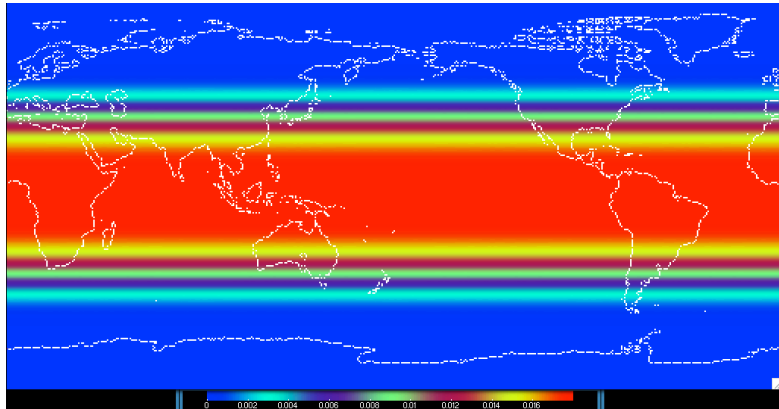
$$(3) \quad m_i \equiv \frac{\rho_i}{\rho_d},$$

where ρ_d is the mass of dry air per unit volume of moist air and ρ_i the mass of water substance of type i per unit volume of moist air. Here we consider the water substances water vapor m_v , cloud

Adding condensate loading is straight forward!

Initialization of idealized moist baroclinic wave (Ullrich et al., 2013) with dry-mass vertical coordinates

Relative humidity at level 30, day 0

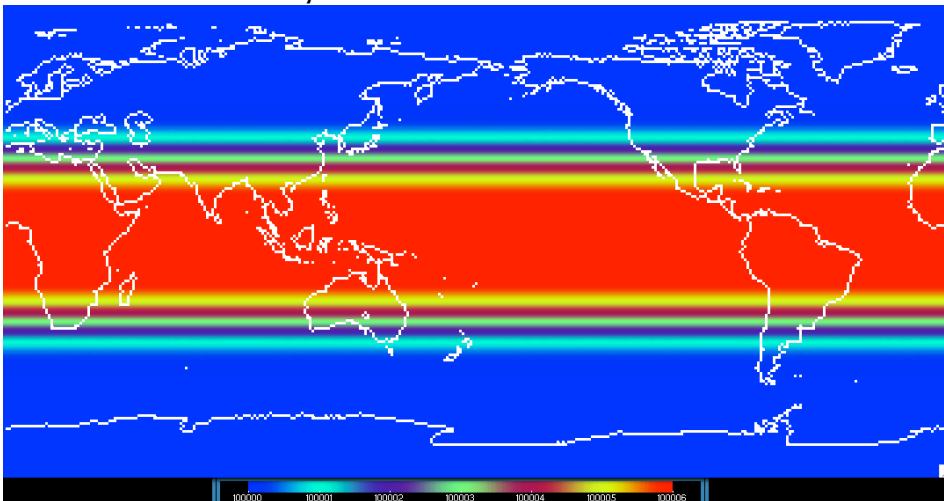


Dry surface pressure is obtained by high-order Gaussian quadrature integration of analytic moisture profile to get water vapor pressure (WVP):

$$p_{s_{dry}} = p_{s_{wet}} - wvp$$

$$p(\eta) = A(\eta)p_0 + B(\eta)p_{s_d},$$

PS_{dry} , day 0 ($PS_{wet} = 1000hPa$)

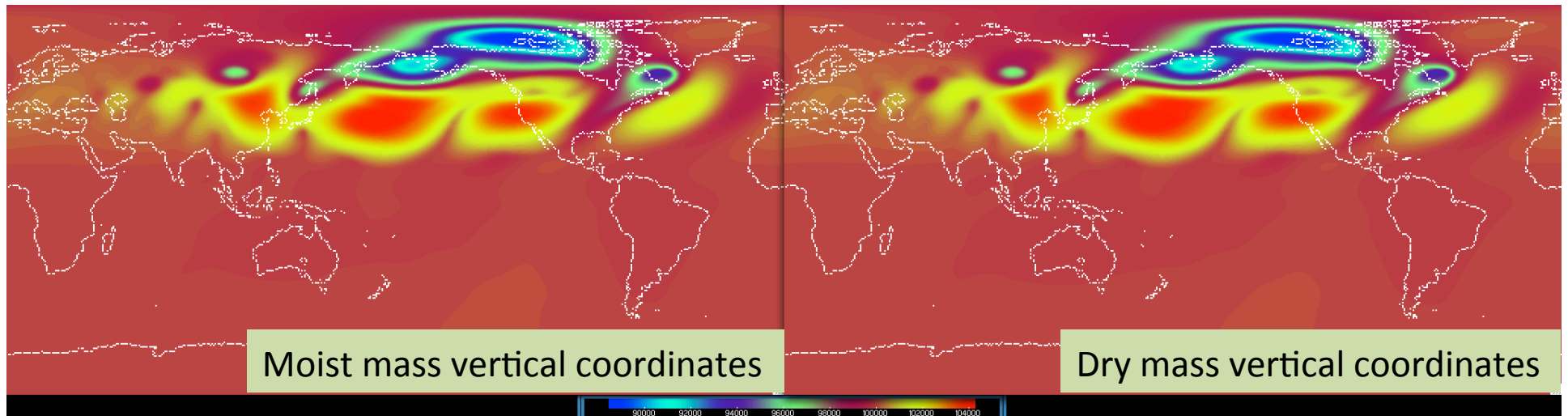


For the dry-mass vertical coordinate $p_{s_{dry}}$ and q (dry mixing ratio for water vapor) must be initialized carefully to get a balanced initial state:

PS_{wet} matches 1000hPa to $1E-10!$

Initialization of idealized moist baroclinic wave (Ullrich et al., 2013) with dry-mass vertical coordinates

Full (moist) surface pressure at day 15





Separating dynamics, tracer and physics grids in CAM-SE

