

NCAR CESM2.0 release of CAM-SE:

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A reformulation of the spectral-element dynamical core in dry-mass vertical coordinates with comprehensive treatment of condensates and energy

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CAM6 Aqua-planet simulation



Figure 4.5 year average zonally averaged total precipitation rate as a function of latitude for CAM6 Aqua-planet

CAM-SE	= new SE dynamical core (CESM2)
CAM-HOMME	= old SE dynamical core (CESM1)
-oldvisc	= CAM-SE with CAM-HOMME
-cpcnst	viscosity coefficients = CAM-SE with $c_p = c_p^{(d)}$ (CAM-HOMME setting)
-ppmlimter	= CAM-SE with limiter on the vertical remapping of wind components
-all	= CAM-SE with the 3 above changes

and finite-volume physics grid



volumes (manuscript in preparation).

Reference

Lauritzen, P.H. and co-authors (2017): NCAR CESM2.0 release of CAM-SE: A reformulation of the spectral-element dynamical core in dry-mass vertical coordinates with comprehensive treatment of condensates and energy: J. Adv Model. Earth Syst., submitted.

Introduction

For some time the atmosphere component of the CESM (Community Earth System Model), called CAM (Community Atmosphere Model), has supported a Spectral-Element (SE) dynamical core option; the SE dynamical core is based on a continuous Galerkin finite-element method discretized on the gnomonic cubed-sphere and it supports static mesh-refinement (see Figure)



ubed-sphere grid that define the elements in CAM-SE. The rees show the cubed-sphere panel edges. (b) A low re ment grid (e referred to as the CONUS-grid (Contiguous United States) used in CAM-SE.

In CESM1 the SE dynamical core was imported from HOMME (High-Order Methods Modeling Environment) as an external code base. In CESM2 the SE dynamical core resides in the CESM repository and has undergone numerous software engineering changes, science changes and performance optimizations described in separate sections below. The new version is referred to as CAM-SE (=SE in CESM2) and its predecessor is referred to as CAM-HOMME (=SE in CESM1).



Dry-mass vertical coordinate

Consider a general terrain following vertical coordinate n(d) that is a function of dry air mass per unit area M(d)

 $\eta^{(d)} = h(M^{(d)}, M_s^{(d)})$ where Ms(d) is dry air mass in a column (per unit area) and

 $h(M_s^{(d)}, M_s^{(d)}) = 1$ and $h(M_t^{(d)}, M_s^{(d)}) = 0$.

where Mt(d) is dry air mass (per unit area) above model top

=> Model interface levels (index k+1/2) are defined as

$$M_{k+1/2}^{(d)} = A_{k+1/2} M_t^{(d)} + B_{k+1/2} M_s^{(d)}$$

where Ak+/12 and Bk+1/2 are the 'usual' hybrid coefficients.

Note that by removing superscript (d) from the equations above (so that the dry variables represent moist variables), then the vertical coordinate is the usual hybrid-pressure coordinate widely used in hydrostatic global modeling



 $\mathcal{L}_{all} = \{`d`, `wv`, `cl`, `ci`, `rn`, `sw`\}$

referring to dry air, water vapor, cloud liquid, cloud ice, rain, and snow. Including condensates in the equations of motion has the following consequences:

Hydrostatic equation includes condensates

 $p(z) = -g \int_{-\infty}^{z'=\infty} \rho \, dz',$ $= -g \int_{z'=z}^{z'=\infty} \rho^{(d)} \left(\sum_{e \in \mathcal{E}} m^{(\ell)} \right) dz'$

where g is gravitational acceleration, $\rho^{(d)}$ is the density of dry air and total density is

 $\rho = \rho^{(d)} \left(\sum_{i=1}^{d} m^{(\ell)} \right)$

where m^(l) is (dry) mixing ratio for component l of moist air.

Energy conversion term in the thermodynamic equation

 $\frac{\partial T}{\partial t} + \mathbf{v} \cdot \nabla_{\eta^{(d)}} T - \frac{1}{c_p \rho} \omega = 0$ where ω is vertical pressure velocity, *T* temperature and

 $c_p = \frac{\sum_{\ell \in \mathcal{L}_{all}} c_p^{(\ell)} m^{(\ell)}}{\sum_{\ell \in \mathcal{L}_{all}} m^{(\ell)}}$

where $c_n^{(l)}$ is the heat capacity of component *l* at constant *p*. includes condensates. Same for pressure-gradient force.

Total energy equation includes condensates

Integrate adiabatic and frictionless equations of motion over the entire domain

 $\frac{\partial}{\partial t} \int_{\eta=0}^{\eta=1} \iint_{S} \left(\frac{\partial M^{(d)}}{\partial \eta^{(d)}} \right) \sum_{\ell \in \mathcal{L}_{all}} \left[m^{(\ell)} \left(K + c_p^{(\ell)} T + \Phi_s \right) \right] dA d\eta^{(d)} = 0.$

where $K=0.5 \times v^2$ is kinetic energy (per unit mass), and Φ_s is surface geopotential.

Modified hyperviscosity

Hyperviscosity applied on approximate pressure surfaces



and for $\Delta M^{(d)}$ (=dry layer mass) only apply hyperviscosity to difference between $\Delta M^{(d)}$ and reference dry layer mass



Viscosity coefficients have been reduced significantly:

Reduced damping

Figure Total kinetic energy spectrum of the horizontal winds at the 200 hPa level in CAM-HOMME and CAM-SE at 1° horizontal resolution, computed as the mean spectra from 30 days of 6-hourly instantaneous spectra. Black line is the κ^{-3} reference scaling, where k is spherical wave-number.

mproved performance

CAM6 Aqua-Planet (incl. I/O)



Figure Throughput in terms of simulated years per day for CAM6 Aqua-planet including standard I/O as a function of number of nodes on NCAR's Chevenne supercomputer. Note that for the right-most data-points there is only 9 physics columns per processor.



simulations

Optional finite-volume tracer transport



Tracer transport with CSLAM (Conservative Semi-Lagrangian Muli-tracer) scheme consistently coupled to SE (Mon. Wea. Rev. paper: DOI:10.1175/MWR-D-16-0258.1).

State passed to physics is integrated over (CSLAM) finite-