



Desirable properties of transport schemes

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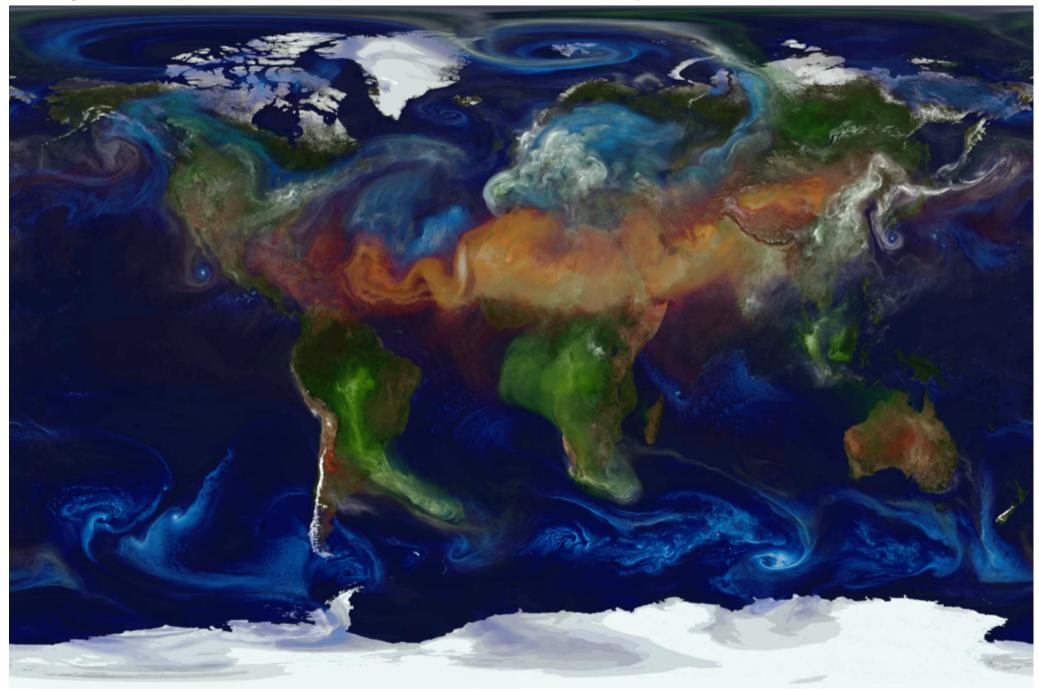


Desirable properties of transport schemes

What physical properties of the continuous equation of motion are important to respect in discretization schemes?

Here I'll focus on the continuity equation ...

GEOS-5 simulation: winds transporting aerosols (5/2005-5/2007) In general, dust appears in shades of orange, sea salt blue, sulfates white, and carbon green



The most important continuity equation in modeling

Consider the continuity equation for dry air

$$\frac{\partial \rho_d}{\partial t} + \nabla \cdot (\rho_d \mathbf{v}) = 0, \tag{1}$$

where ρ_d is the density of dry air (mass per unit volume of Earth's atmosphere) and **v** is a 3D velocity vector.

The most important continuity equation in modeling

Consider the conic ity equation for dry air approximately 0.01hPa globally $\frac{\partial \rho_d}{\partial t} + \nabla \cdot (\rho_d \mathbf{v}) = 0,$

where ρ_d is atmosphere

dry air (mass per unit volume of Earth's velocity vector.

Dry air makes up 99.75% of the mass of the atmosphere:

mean mass of dry air = $5.1352 \pm 0.0003 \times 10^{18} \text{ kg}$

mean mass of atmosphere = $5.1480 \times 10^{18} \text{ kg}$

Trenberth and Smith (2005)

Accurate to

(1)

The most important continuity equation in modeling

Consider the continuity equation for dry air

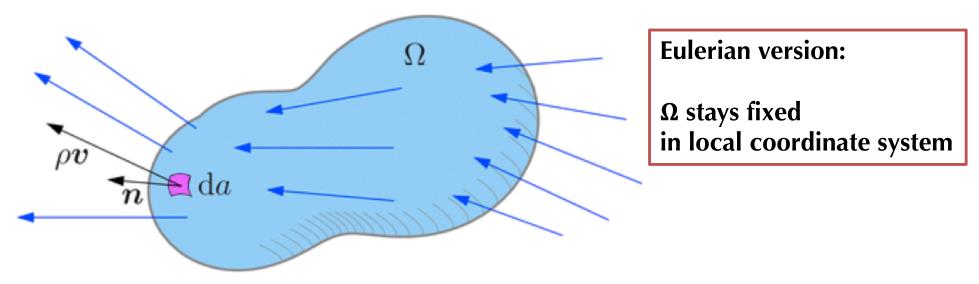
Accurate to approximately 0.01hPa globally

$$\frac{\partial \rho_d}{\partial t} + \nabla \cdot (\rho_d \mathbf{v}) = 0, \qquad (1)$$

where ρ_d is the density of dry air (mass per unit volume of Earth's atmosphere) and **v** is a 3D velocity vector.

Note that the continuity equation for air is "tightly" coupled with momentum and thermodynamic equations

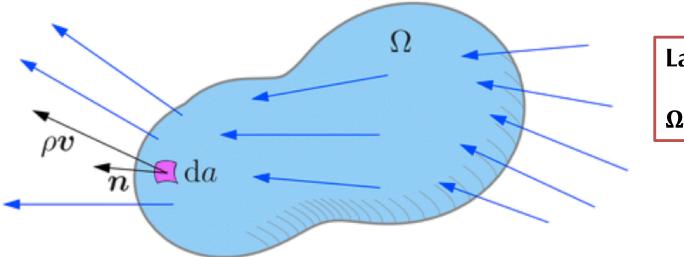
To solve (1) we need to know the velocity field!



The continuity equation is a conservation law for mass:

$$\frac{\partial}{\partial t} \iiint_{\Omega} \rho_d dV = - \iiint_{\Omega} \nabla \cdot (\rho_d \mathbf{v}) \, dV,$$
$$= - \oiint_{\partial \Omega} (\rho_d \mathbf{v}) \cdot \mathbf{n} \, dS,$$

where Ω is a fixed volume, $\partial \Omega$ the surface of Ω and **n** is outward pointing unit vector normal to the local surface. \Rightarrow The flux of mass through the area *da* is *da* times $\rho_d \mathbf{v} \cdot \mathbf{n}$.



Lagrangian version:

 Ω moves with the flow

The continuity equation is a conservation law for mass:

$$\frac{\partial}{\partial t} \iiint_{\Omega} \rho_d dV = - \iiint_{\Omega} \sqrt{(\rho_d v)} dV,$$
$$= - \oiint_{\partial \Omega} \sqrt{(\sigma_d v)} n dS,$$

Conservation of mass

Consider the continuity equation for X (e.g., water vapor, cloud ice, cloud liquid, chemical species, ...)

$$\frac{\partial}{\partial t} (m_{\chi} \rho_d) + \nabla \cdot (m_{\chi} \rho_d \mathbf{v}) = \rho_d S^{m_{\chi}}, \qquad (1)$$

where S^{m_X} is the source of X and/or sub-grid-scale transport term.

Integrate (1) over entire atmosphere Ω_{tot}

$$\frac{\partial}{\partial t} \iiint_{\Omega_{tot}} (m_x \rho_d) \, dV = \iiint_{\Omega_{tot}} \rho_d \, S^{m_X} \, dV.$$

Note: sub-grid-scale transport integrates to zero! Global mass only changes due to sources/sinks S^{m_X} .

Global conservation of mass

Globally the change in mass is exactly balanced by the source/sink terms!

The resolved-scale tracer transport must not be a spurious source or sink of mass

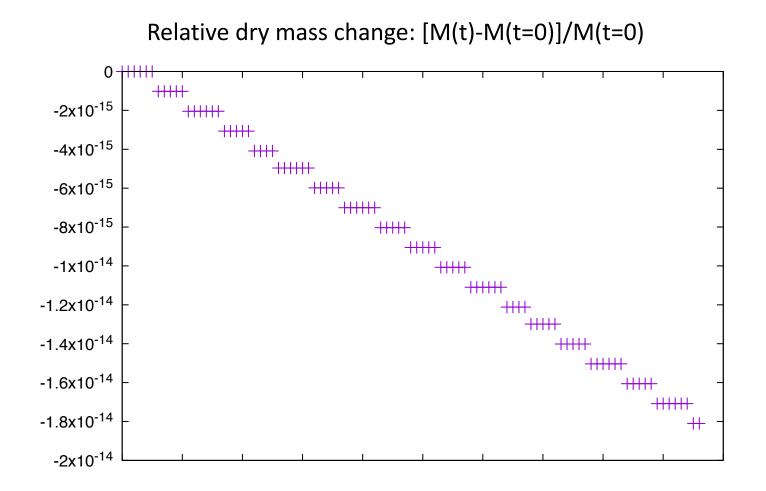
Why is that a problem?

Integrate (1) over entire atmosphere Ω_{tot}

$$\frac{\partial}{\partial t} \iiint_{\Omega_{tot}} (m_x \rho_d) \, dV = \iiint_{\Omega_{tot}} \rho_d \, S^{m_X} \, dV.$$

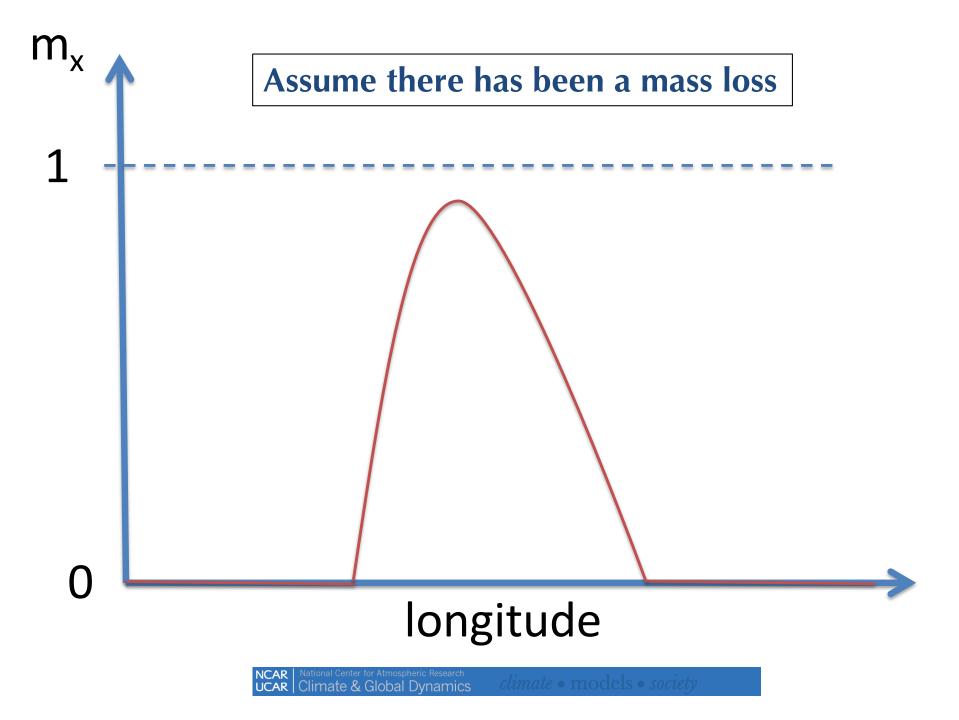
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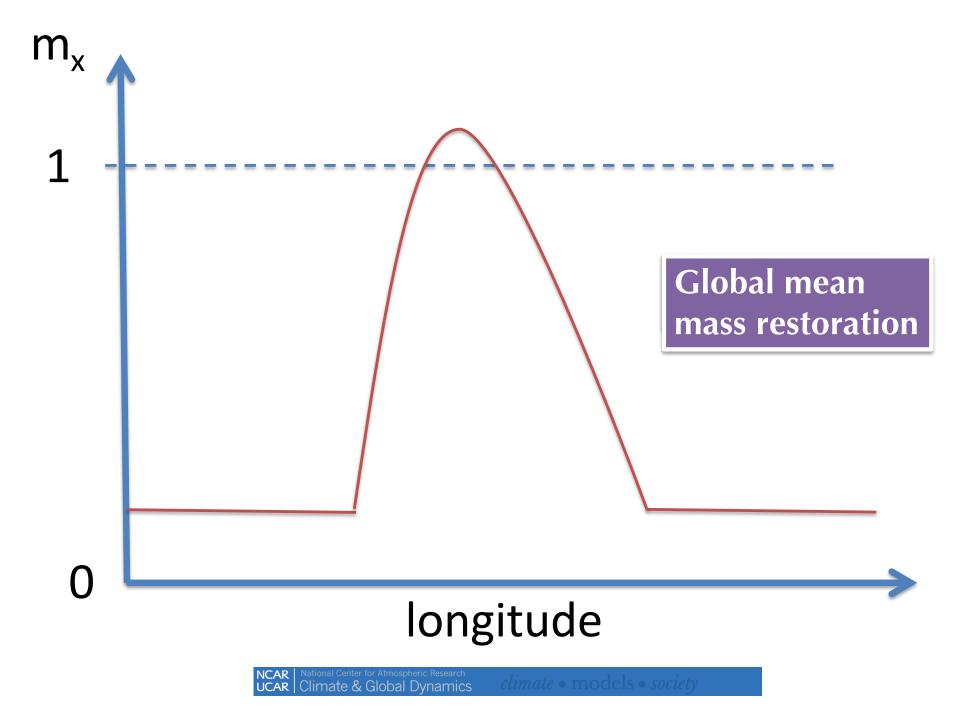
Accumulation of error

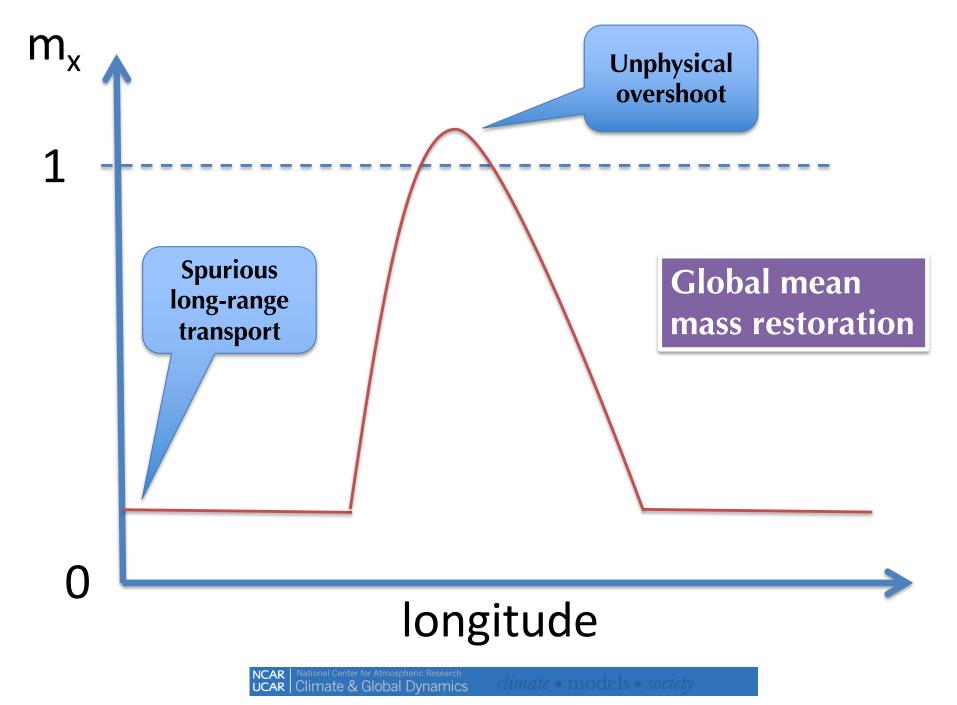


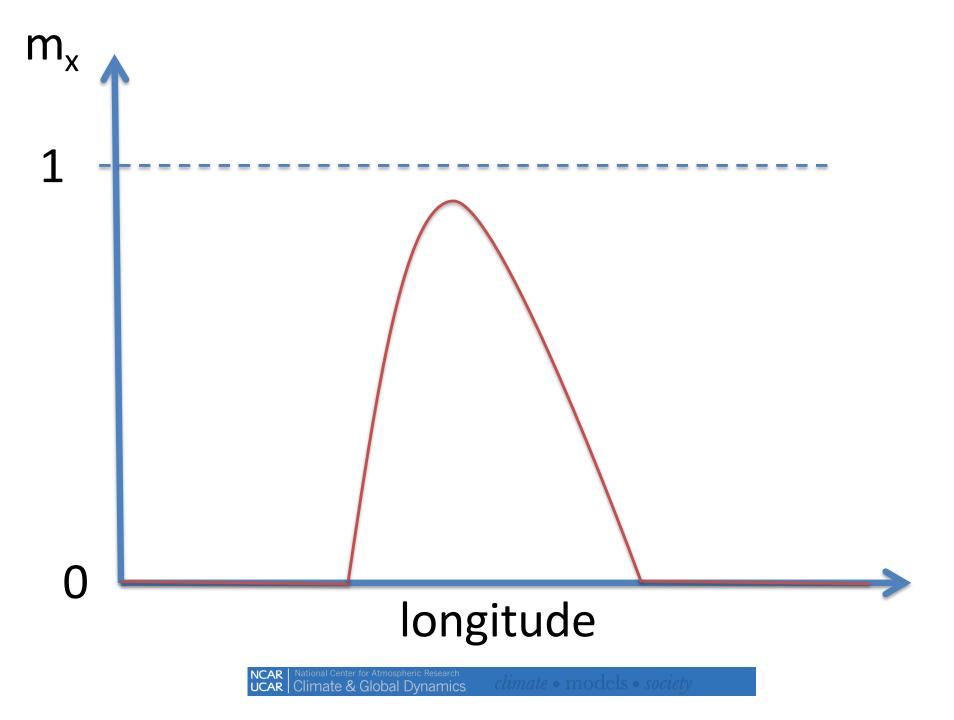
1000 year simulation \approx O(10⁷) 30 minute time-steps

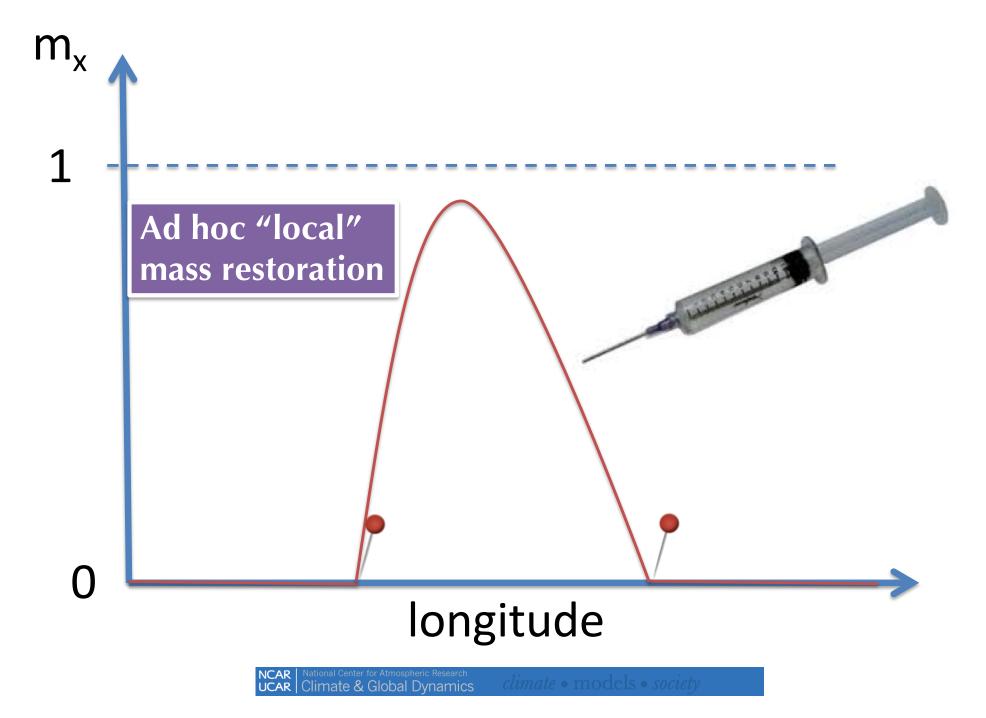
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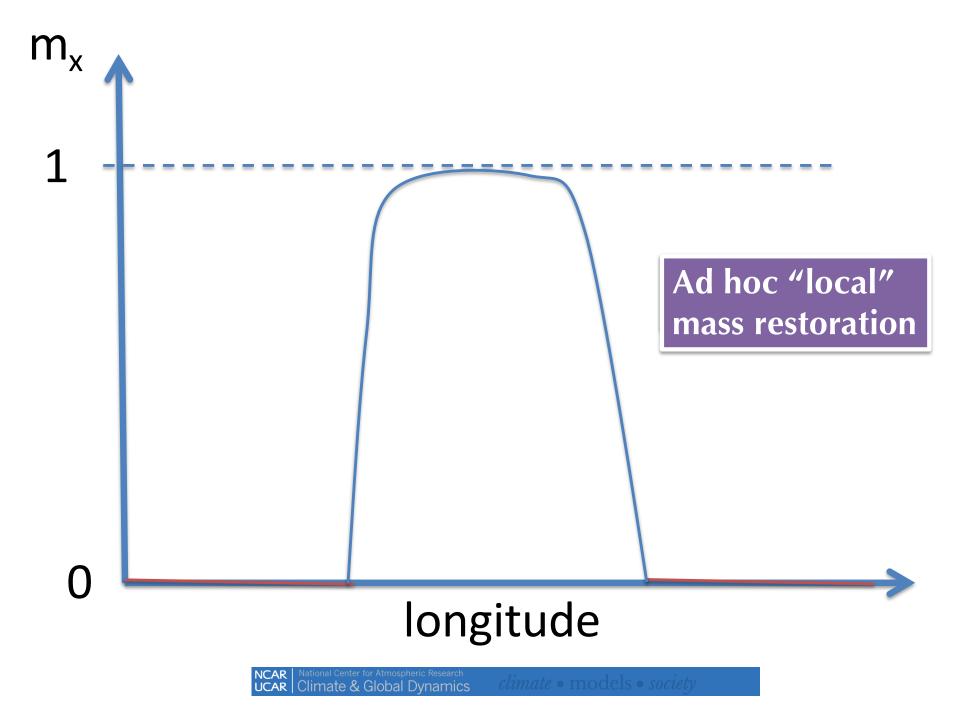


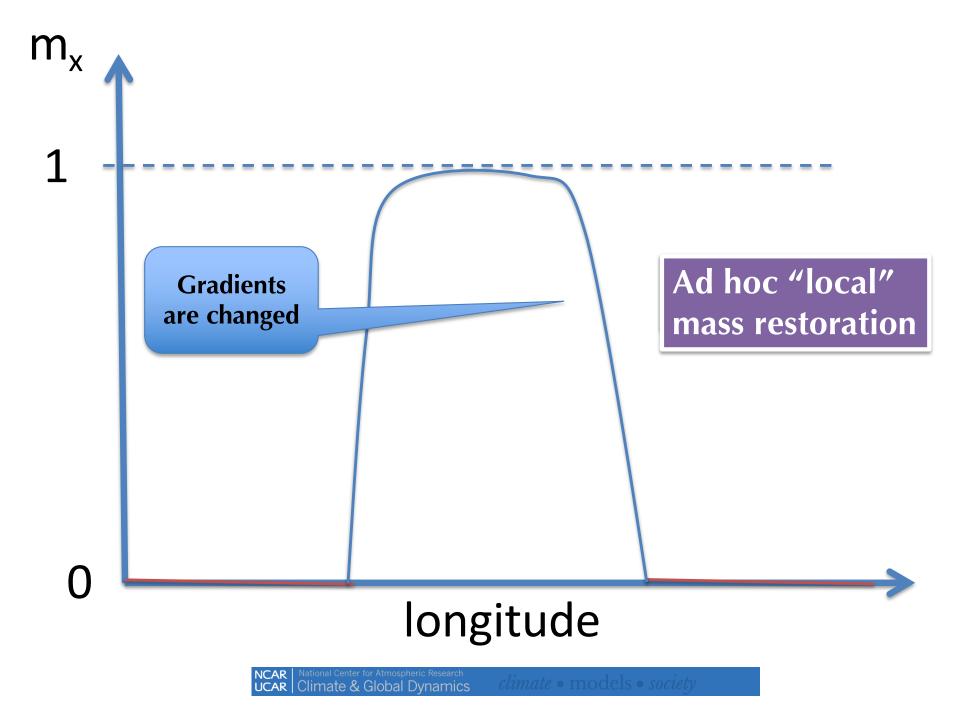




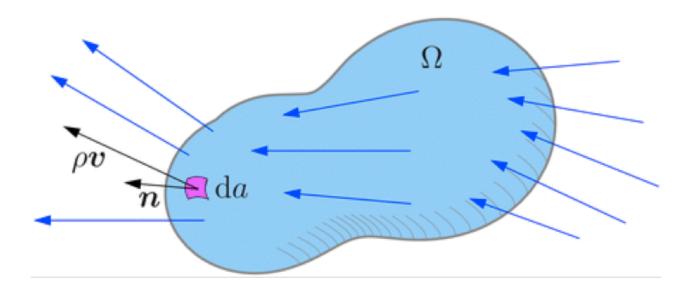








Inherent local mass-conservation is desirable



The continuity equation is a conservation law for mass:

$$\frac{\partial}{\partial t} \iiint_{\Omega} \rho_d dV = - \iiint_{\Omega} \nabla \cdot (\rho_d \mathbf{v}) \, dV,$$
$$= - \oiint_{\partial \Omega} (\rho_d \mathbf{v}) \cdot \mathbf{n} \, dS$$

Conservation of m_x along parcel trajectories

Consider the continuity equation for dry air and X

$$\frac{\partial \rho_d}{\partial t} + \nabla \cdot (\rho_d \mathbf{v}) = 0, \qquad (2)$$

$$\frac{\partial}{\partial t} (m_x \rho_d) + \nabla \cdot (m_X \rho_d \mathbf{v}) = \rho_d S^{m_X}, \qquad (3)$$

respectively. Applying the chain rule to (3), re-arranging and substituting (2) implies

$$\frac{Dm_X}{Dt}=S^{m_X},$$

where $D/Dt = \frac{\partial}{\partial t} + \vec{v}\nabla$ is the total (material) derivative.

Conservation of m_x along parcel trajectories

Consider the continuity equation for dry air and X

$$\frac{\partial \rho_d}{\partial t} + \nabla \frac{\partial \rho_d}{\partial t} + \nabla \cdot (m_x \rho_d) + \nabla \cdot (m_x \rho$$

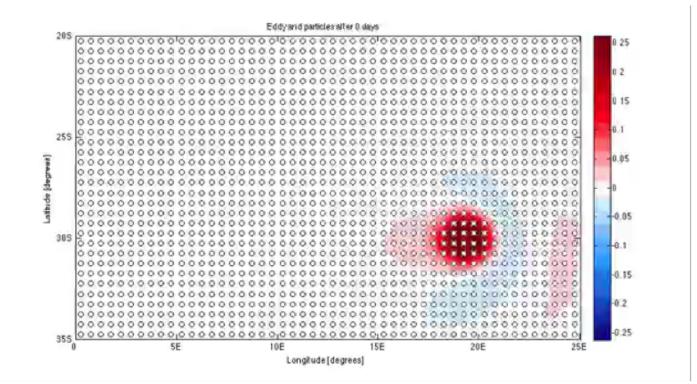
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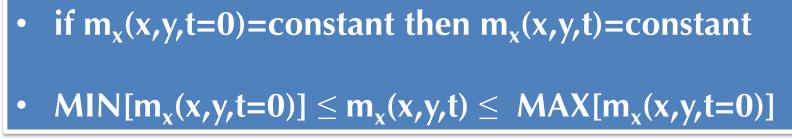
If the discretization scheme is based on the advective form of the continuity equation (.e.g, grid-point semi-Lagrangian schemes) then inherent massconservation is not guaranteed

$$\frac{Dm_X}{Dt} = S^{m_X},$$

where $D/Dt = \frac{\partial}{\partial t} + \vec{v}\nabla$ is the total (material) derivative.

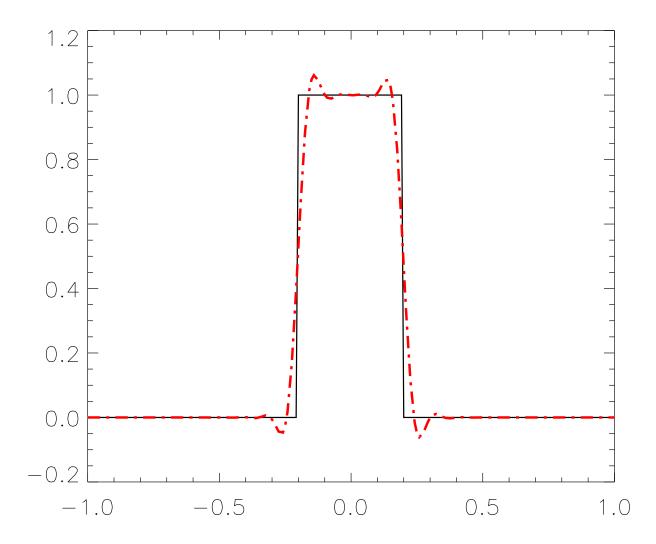
Conservation of m_x along parcel trajectories (if no sources/sinks of m_x)





Source: https://www.youtube.com/watch?v=tEHQH7Uly-8

Conservation of m_x along parcel trajectories (if no sources/sinks of m_x)



Nair et al., (2011)

Conservation of m_x along parcel trajectories

Atmospheric modelers tend to be a bit loose with the term `monotone'!

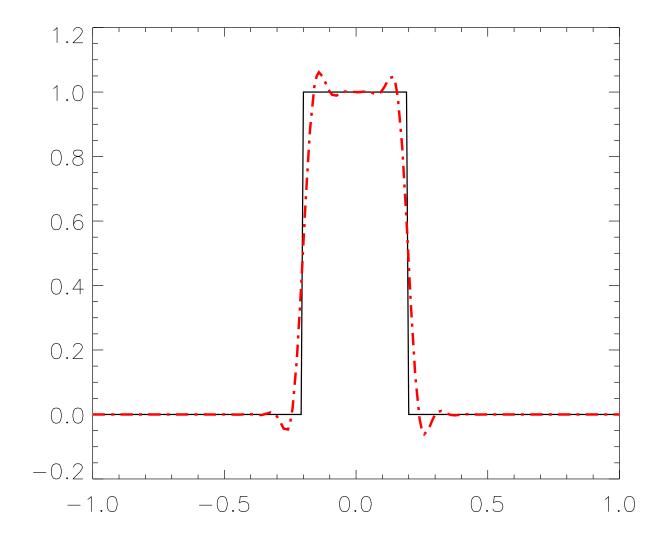
When modelers refer to "non-oscillatory", "shape-preserving", "physical realizable" or "monotone" they usually refer to the **monotonicity property** as defined by Harten (1983):

- 1. No new local extrema in m_x may be created
- 2. The value of a local minima/(maxima) is nondecreasing/(nonincreasing)

There are "stricter" characterizations such as total variation diminishing (TVD), however, they are probably too strong for our applications

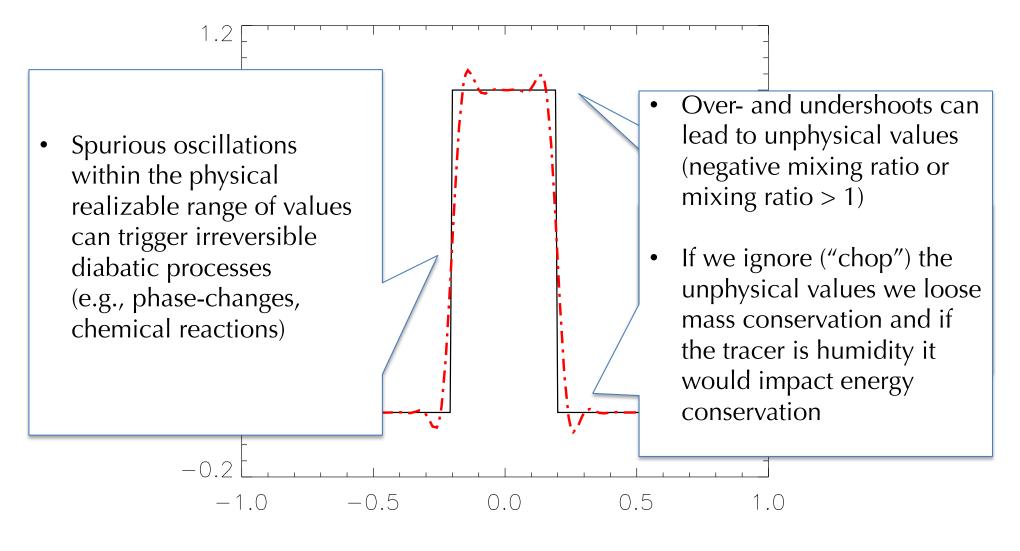
=> the monotonicity property applies to mixing ratio m_x and not tracer mass!

Why is the monotonicity property so important



Nair et al., (2011)

Why is the monotonicity property so important



Nair et al., (2011)

Conservation of mass along parcel trajectories

Note that

$$\frac{D\rho_d}{Dt} \neq 0,$$

but

$$\frac{D\rho_d}{Dt} = -\rho_d \nabla \cdot \vec{v}.$$

If we integrate ρ_d over a Lagrangian volume Ω_L then

$$\frac{\partial}{\partial t} \iiint_{\Omega_L} \rho_d \, dV = 0.$$

Lagrangian volumes are rapidly distorting

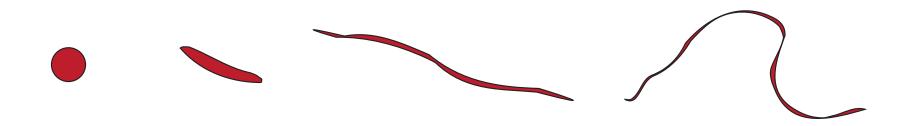
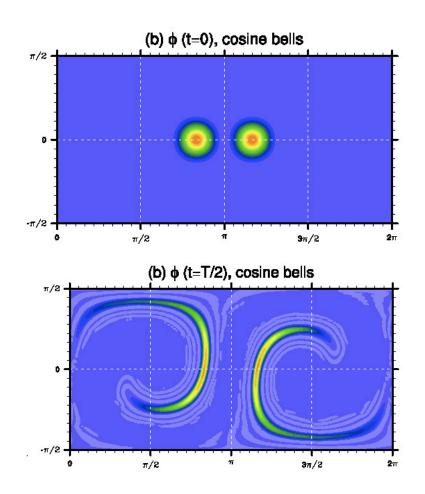


Fig. 2: In the highly nonlinear flows that characterize fluid motion in the atmosphere and ocean, Lagrangian control volumes are rapidly distorted due the presence of strong shear, rotation and dilation. The rapid distortion of Lagrangian control volumes makes the formulation of numerical models within the Lagrangian reference frame an extremely difficult challenge.

Ringler (2011)

Filament diagnostic (M. Prather, UCI)



The "filament" preservation diagnostic is formulated as follows. Define $A(\tau, t)$ as the spherical area for which the spatial distribution of the tracer $\phi(\lambda, \theta)$ satisfies

$$\phi(\lambda,\theta) \ge \tau, \tag{27}$$

at time *t*, where τ is the threshold value. For a non-divergent flow field and a passive and inert tracer ϕ , the area $A(\tau, t)$ is invariant in time.

The discrete definition of $A(\tau, t)$ is

$$A(\tau,t) = \sum_{k \in \mathcal{G}} \Delta A_k,$$
(28)

where ΔA_k is the spherical area for which ϕ_k is representative, *K* is the number of grid cells, and *G* is the set of indices

$$\mathcal{G} = \{k \in (1, \dots, K) | \phi_k \ge \tau\}.$$
⁽²⁹⁾

For Eulerian finite-volume schemes ΔA_k is the area of the *k*-th control volume. For Eulerian grid-point schemes a control volume for which the grid-point value is representative must be defined. Similarly for fully Lagrangian schemes based on point values (parcels) control volumes for which the point values are representative must be defined. Note that the "control volumes" should span the entire domain without overlaps or "cracks" between them.

Define the filament preservation diagnostic

$$\ell_{f}(\tau,t) = \begin{cases} 100.0 \times \frac{A(\tau,t)}{A(\tau,t=0)} & \text{if } A(\tau,t=0) \neq 0, \\ 0.0, & \text{otherwise.} \end{cases}$$
(30)

For infinite resolution (continuous case) and a non-divergent flow, $\ell_f(\tau, t)$ is invariant in time: $\ell_f(\tau, t = 0) = \ell_f(\tau, t) = 100$ for all τ . At finite resolution, however, the filament

This diagnostic does not rely on an analytical solution!

Lauritzen et al. (2012)

Filament diagnostic

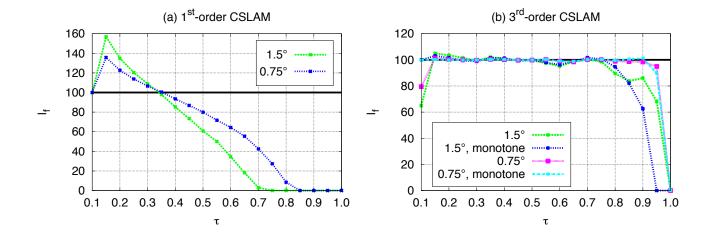


Fig. 6. Filament diagnostics $\ell_f(t=T/2)$ as a function of threshold value τ for different configurations of the CSLAM scheme with Courant number 5.5. (a) 1st-order version of CSLAM at $\Delta \lambda = 1.5^{\circ}$ and $\Delta \lambda = 0.75^{\circ}$, and (b) 3rd-order version of CSLAM with and without monotone/shape-preserving filter at resolutions $\Delta \lambda = 1.5^{\circ}$ and $\Delta \lambda = 0.75^{\circ}$.

Tracer mass and air mass consistency

Consider the continuity equation for dry air and X (no sources/sinks)

$$\frac{\partial \rho_d}{\partial t} + \nabla \cdot (\rho_d \mathbf{v}) = 0, \qquad (4)$$

$$\frac{\partial}{\partial t} (m_x \rho_d) + \nabla \cdot (m_X \rho_d \mathbf{v}) = 0, \qquad (5)$$

respectively.

Note that if m_x is 1 then (5) reduces to (4).

A scheme satisfying this is referred to as "free-stream preserving"

Examples of tracer mass and air mass consistency violation

Consider the continuity equation for dry air and X (no **S** Prescribed wind and mass fields from , e.g., re- $\frac{\partial \rho_d}{\partial t} + \nabla \cdot (\rho_d \mathbf{v}) = 0,$ $\frac{\partial}{\partial t} (m_x \rho_d) + \nabla \cdot (m_X \rho_d \mathbf{v}) = 0,$ analysis. (4)(5)Solve (4) and (5) with respectively. different numerical methods, on different grids Note that if m_x is 1 then (5) reduces to (4). and/or different time-steps

A scheme satisfying this is referred to as "free-stream preserving"

Examples of tracer mass and air mass consistency violation

Consider the continuity equation for dry air and X (no sources/sinks)

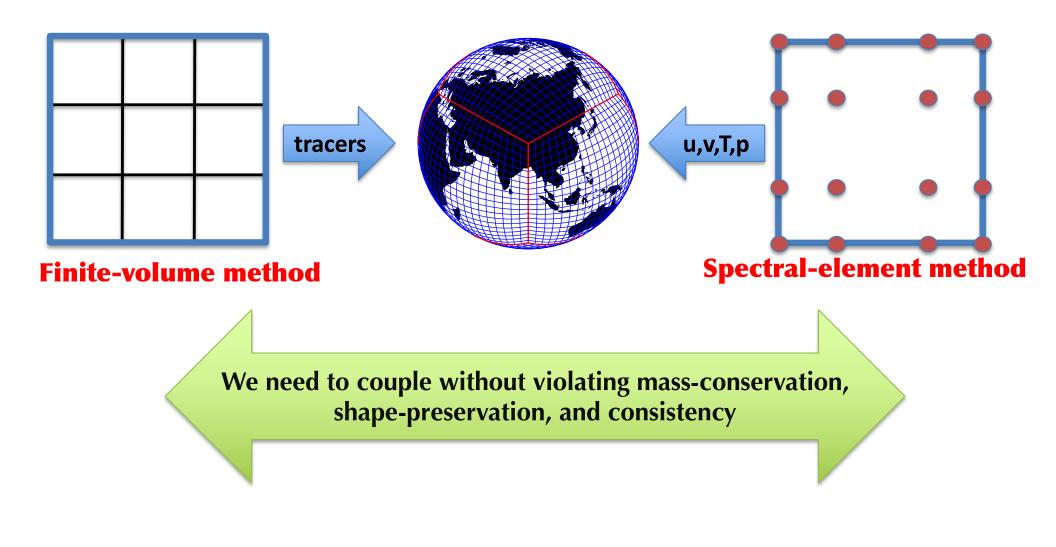
$$\frac{\partial \rho_d}{\partial t} + \nabla \cdot (\rho_d \mathbf{v}) = 0, \qquad (4)$$

$$\frac{\partial}{\partial t} (m_x \rho_d) + \nabla \cdot (m_X \rho_d \mathbf{v}) = 0, \qquad (5)$$

If consistency is violated:

- monotonicity preservation may be violated
- tracer mass-conservation may be violated
- (5) may start evolving independently of (4)

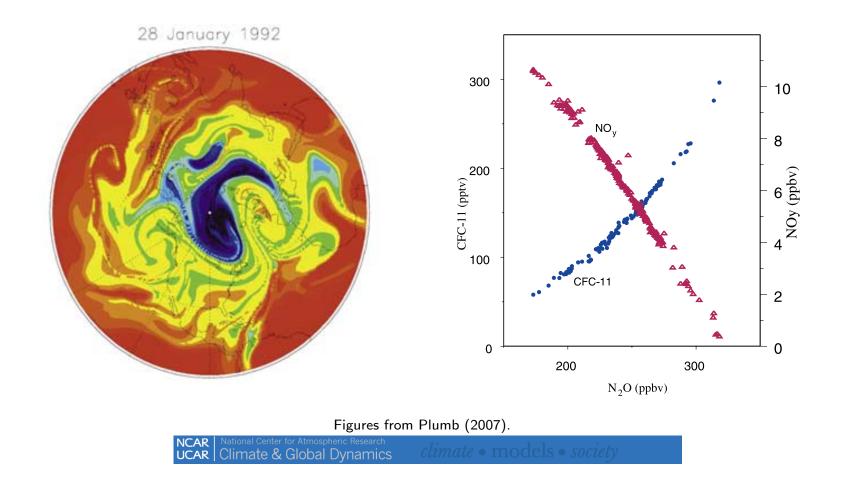
Example: Separating transport and dynamics grids/methods in CAM-SE



Correlations between long-lived species in the stratosphere

Relationships between long-lived stratospheric tracers, manifested in similar spatial structures on scales ranging from a few to several thousand kilometers, are displayed most strikingly if the mixing ratio of one is plotted against another, when the data collapse onto remarkably compact curves. - Plumb (2007)

E.g., nitrous oxide (N_2O) against 'total odd nitrogen' (NO_V) or chlorofluorocarbon (CFC's)



Correlations between long-lived species in the stratosphere

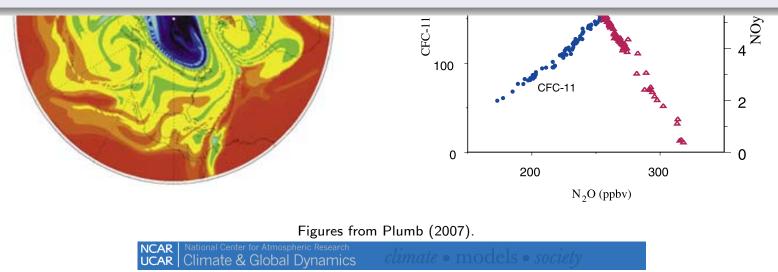
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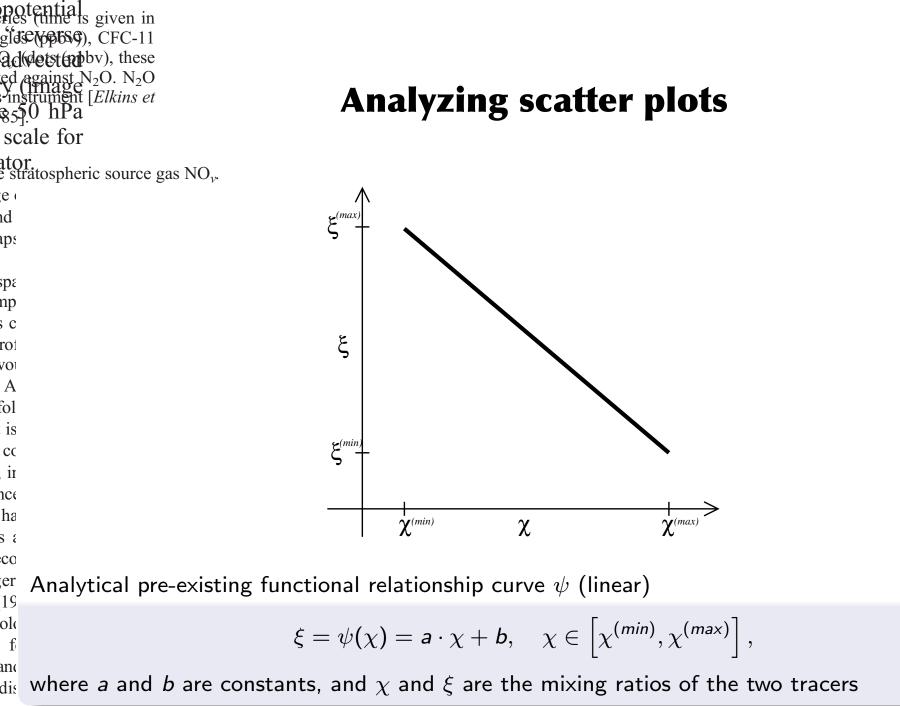
E.g., nitrous oxide $(N_2 O)$ against 'total odd nitrogen' (NO_V) or chlorofluorocarbon (CFC's)

Similarly:

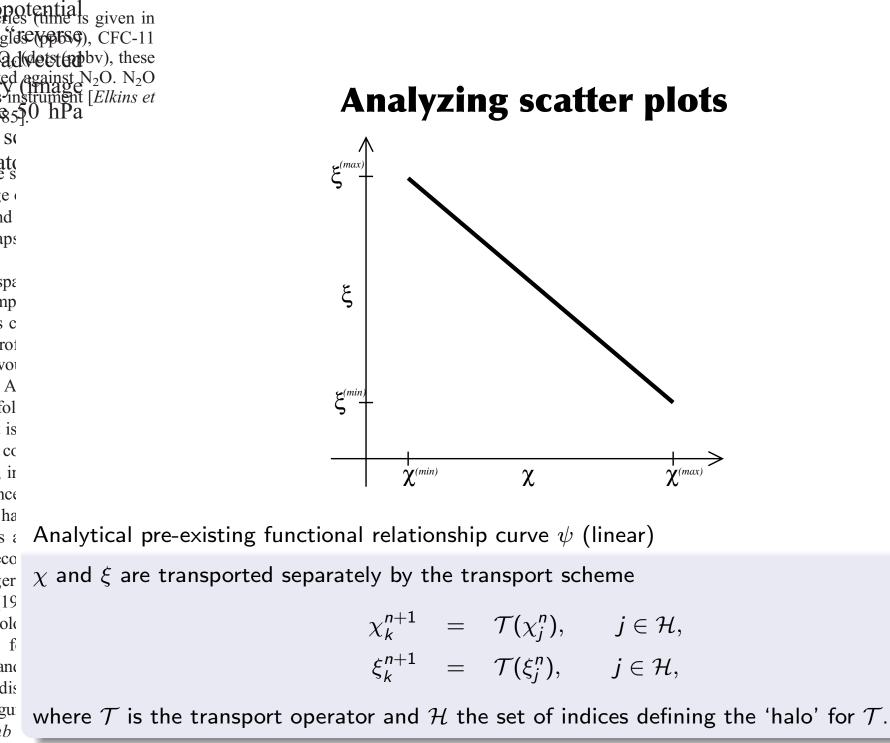
- The total of chemical species within some chemical family may be preserved following an air parcel although the individual species have a complicated relation to each other and may be transformed into each other through chemical reactions (e.g., total chlorine)
- Aerosol-cloud interactions (Ovtchinnikov and Easter, 2009)

The transport operator should ideally not perturb pre-existing functional relationships



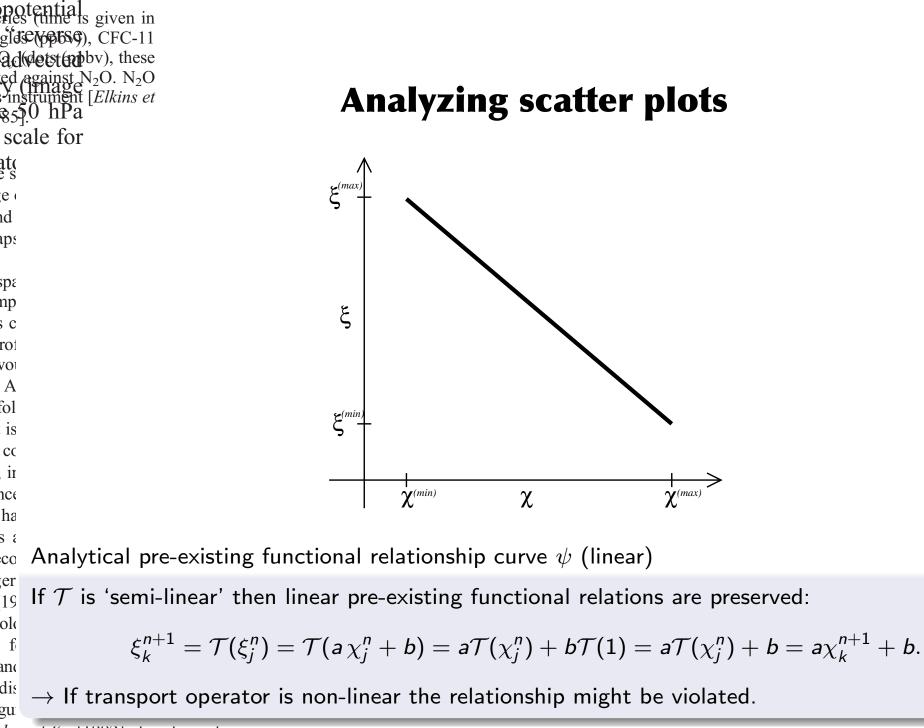


b and Ko [1992] also showed c mixing extends globally, the rtical flux of any species is e, the slope of the tracer-tracer



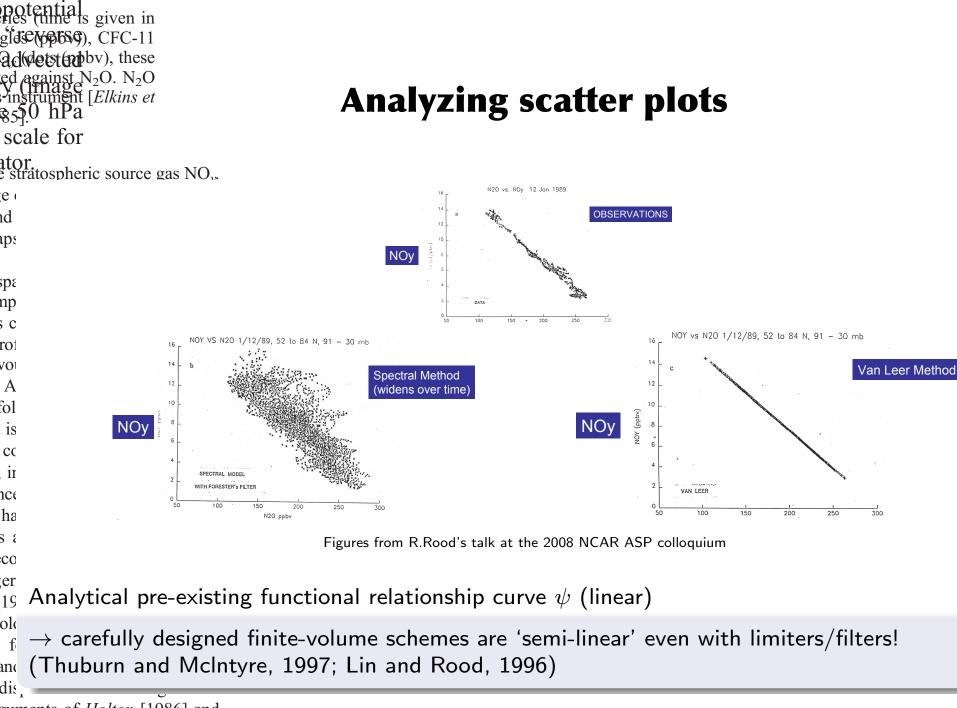
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guments of *Holton* [1986] and *b and Ko* [1992] also showed c mixing extends globally, the rtical flux of any species is e, the slope of the tracer-tracer is $d_{1}^{(2)}(d_{2}^{(1)})$ is accual to the

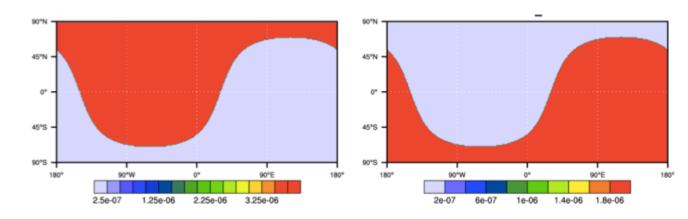
The terminator 'toy'-chemistry test: A simple tool to assess errors in transport schemes

(Lauritzen et al., 2015) See: <u>http://www.cgd.ucar.edu/cms/pel/terminator.html</u>

• Consider 2 reactive chemical species, Cl and Cl₂:

$$Cl_2 \rightarrow Cl + Cl : k_1$$
$$Cl + Cl \rightarrow Cl_2 : k_2$$

• Steady-state solution (no flow):



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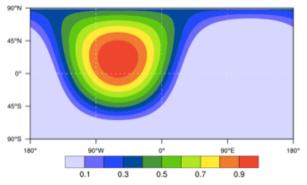
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• In any flow-field Cl_y=Cl+2*Cl₂ should be constant at all times (correlation preservation)





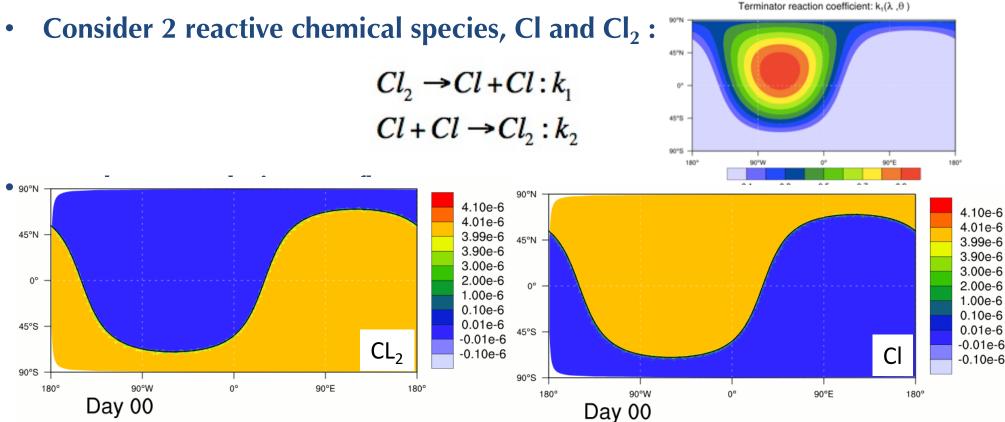
Terminator reaction coefficient: $k_1(\lambda, \theta)$



The terminator 'toy'-chemistry test: A simple tool to assess errors in transport schemes

(Lauritzen et al., 2015) See: http://www.cgd.ucar.edu/cms/pel/terminator.html





In any flow-field Cl_v=Cl+2*Cl₂ should be constant at all times (linear correlation preservation).



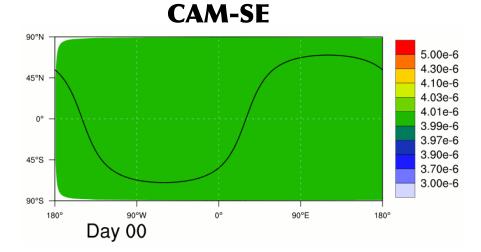
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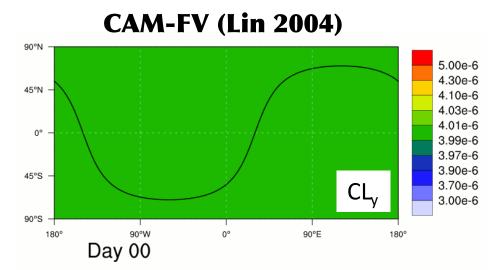
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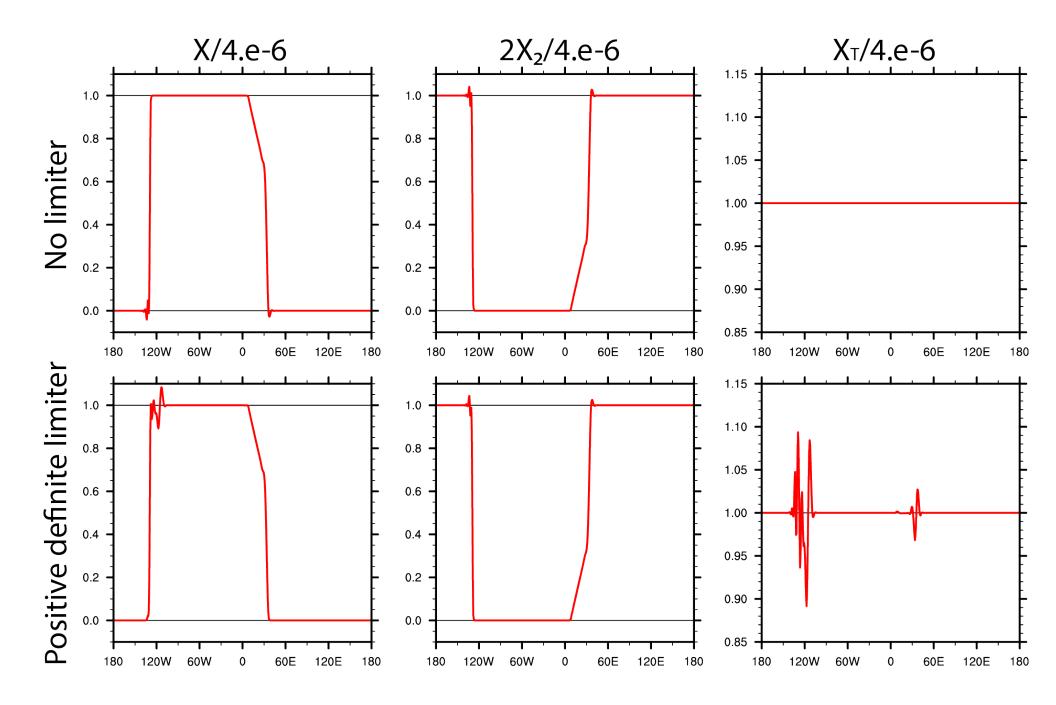


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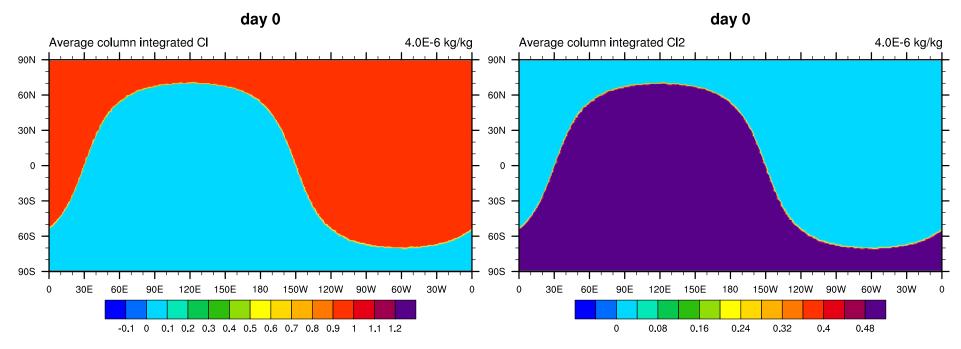


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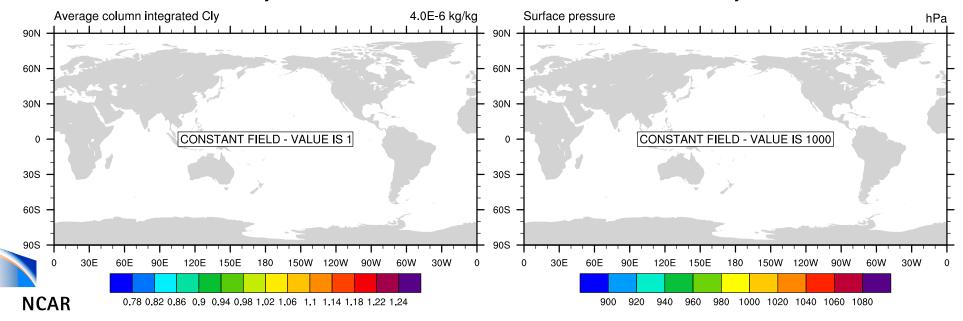


3D version: Initial condition





day 0



CAM-SE



Conserving sum of "families" of species

Chlorine (in CAM-chemistry)

Total Organic Chlorine (set at the surface)

 $\mathsf{T}_{\mathsf{C}\ell}^{\mathsf{O}\mathsf{R}\mathsf{G}} = \mathsf{C}\mathsf{H}_3\mathsf{C}\ell + 3\mathsf{C}\mathsf{F}\,\mathsf{C}\ell_3 + 2\mathsf{C}\mathsf{F}_2\,\mathsf{C}\ell_2 + 3\mathsf{C}\ell\,\mathsf{C}\ell_2\mathsf{F}\mathsf{C}\,\mathsf{C}\ell\,\mathsf{F}_2 + \mathsf{H}\mathsf{C}\mathsf{F}_2\,\mathsf{C}\ell + 4\mathsf{C}\mathsf{C}\ell_4 + 3\mathsf{C}\mathsf{H}_3\mathsf{C}\,\mathsf{C}\ell_3.$

Total Inorganic Chlorine (created from break down of T_{Cl}^{ORG})

 $\mathsf{T}_{\mathsf{C}\ell}^{\mathsf{INORG}} = \mathsf{C}\ell + \mathsf{C}\ell\,\mathsf{O} + \mathsf{O}\,\mathsf{C}\ell\,\mathsf{O} + 2\mathsf{C}\ell_2 + 2\mathsf{C}\ell_2\,\mathsf{O}_2 + \mathsf{HO}\,\mathsf{C}\ell + \mathsf{C}\ell\,\mathsf{O}\,\mathsf{NO}_2 + \mathsf{H}\,\mathsf{C}\ell,$

Total Chlorine

$$TCLY = T_{C\ell}^{ORG} + T_{C\ell}^{INORG}$$

Total chlorine TCLY should be conserved in the upper troposphere and stratosphere (despite complex chemical reactions between the different chlorine species)!

Reactants		Products	Rate
PAN + M	\rightarrow	CH3CO3 + NO2 + M	k(CH3CO3+NO2+M)·1.111E28 ·exp(-14000/T)
CH3CO3 + CH3CO3	\rightarrow	2·CH3O2 + 2·{CO2}	2.50E-12.exp(500/T)
GLYALD + OH	\rightarrow	HO2 + .2·GLYOXAL + .8·CH2O + .8·{CO2}	1.00E-11
GLYOXAL + OH	\rightarrow	$HO2 + CO + \{CO2\}$	1.10E-11
CH3COOH + OH	\rightarrow	CH3O2 + {CO2} + H2O	7.00E-13
C2H5OH + OH	\rightarrow	HO2 + CH3CHO	6.90E-12·exp(-230/T)
C3H6 + OH + M	\rightarrow	PO2 + M	$ko=8.00E-27\cdot(300/T)^{3.50};$
			ki=3.00E-11; f=0.50

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Conserving sum of "families" of species TCLY [mol/mol], ca. 992.55608 hPa, lon avera TCLY [mol/mol], ca. 35.923249 hPa, lon aver-3.80e-09 3 75e-09 .75e-09 3.70e-09 5.70e-09 3.65e-09 50 -3.60e-09 5.60e-09 3.55e-09 3.55e-0 atitude [degrees] 3.50e-09 .50e-09 3.45e-09 145e-09 3.40e-09 3.35e-09 1.35e-09 3.30e-09 3.30e-09 -50 -3.25e-09 3.20e-09 3.20e-09 5.15e-09 3.15e-09 Jul 1998 Mar 1998 May 1998 Sep 1998 Nov 1998 Mar 1998 Jul 1998 Jan 1999 May 1998 Sep 1998 Nov 1998

(left) longitude-averaged surface TCLY as a function of time and latitude: Constant!(right) same as (left) but near tropopause: Spurious 7% deviations (near sharp gradients)!

Problem?

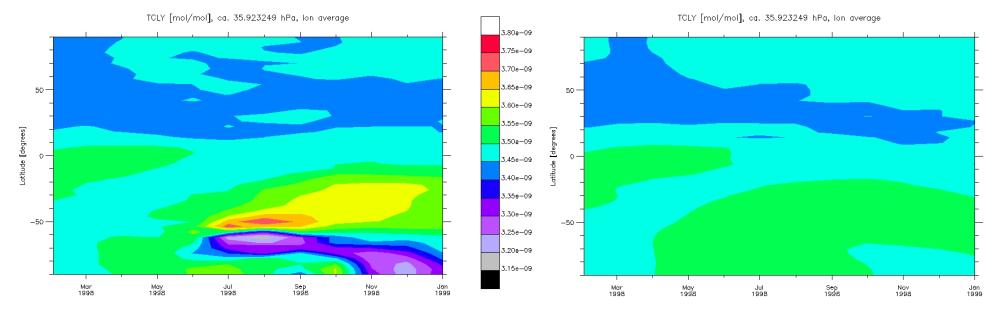
Transport scheme can not maintain the sum when transporting the species individually:

$$\sum_{i=1}^{N_{\chi}} \mathfrak{T}(\chi_{i}) \neq \mathfrak{T}\left(\sum_{i=1}^{N_{\chi}} \chi_{i}\right),$$

"Semi-linear" property is a necessary but not sufficient condition for conserving a sum of more than 2 tracers

where N_{χ} is the number of species χ_i .

Conserving sum of "families" of species

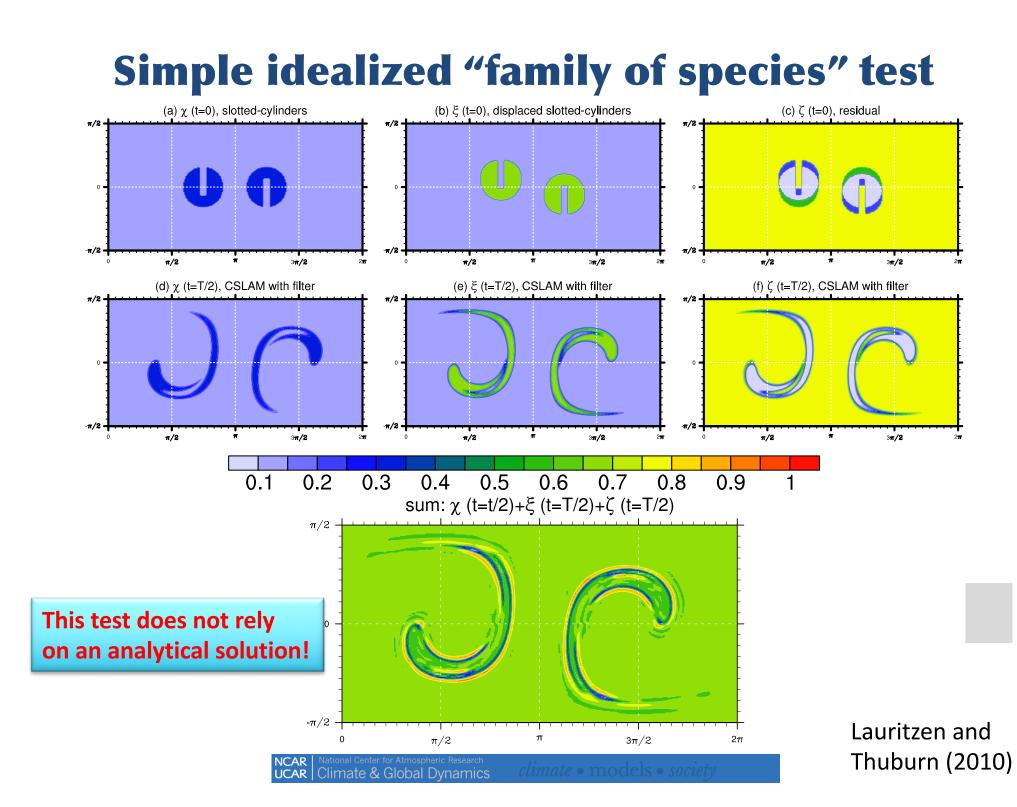


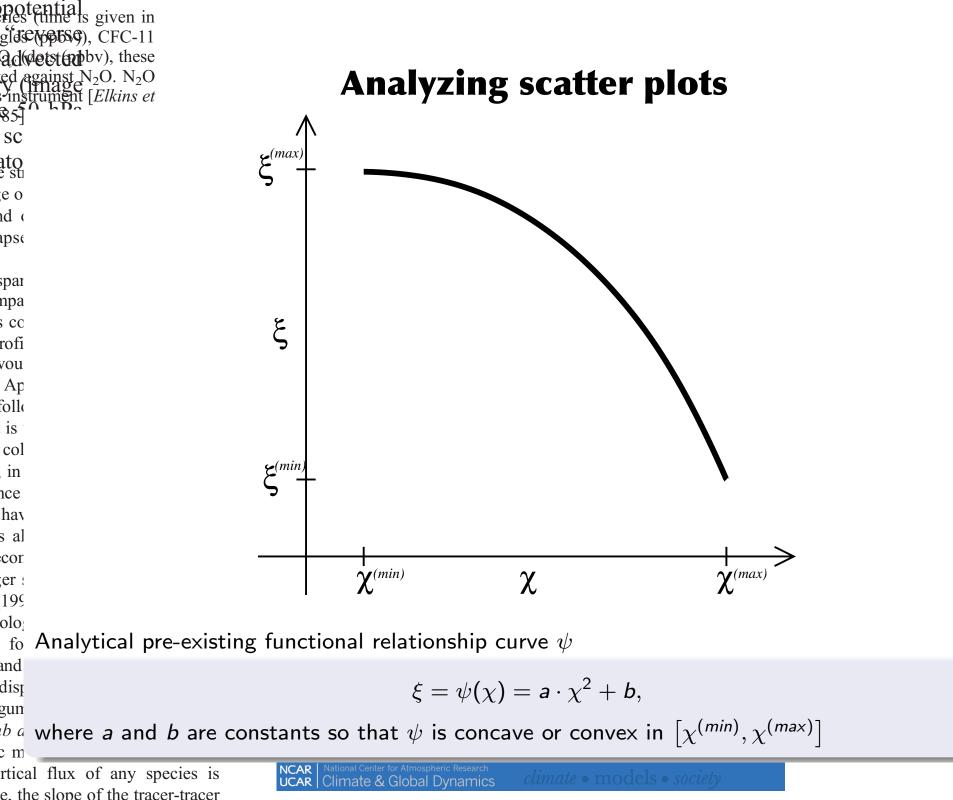
(left) same as previous slide:

• large unphysical deviations from constancy in TCLY near the edge of the polar stratospheric vortex \Rightarrow less TCLY over South pole \Rightarrow less ozone loss (error on the order of 10%).

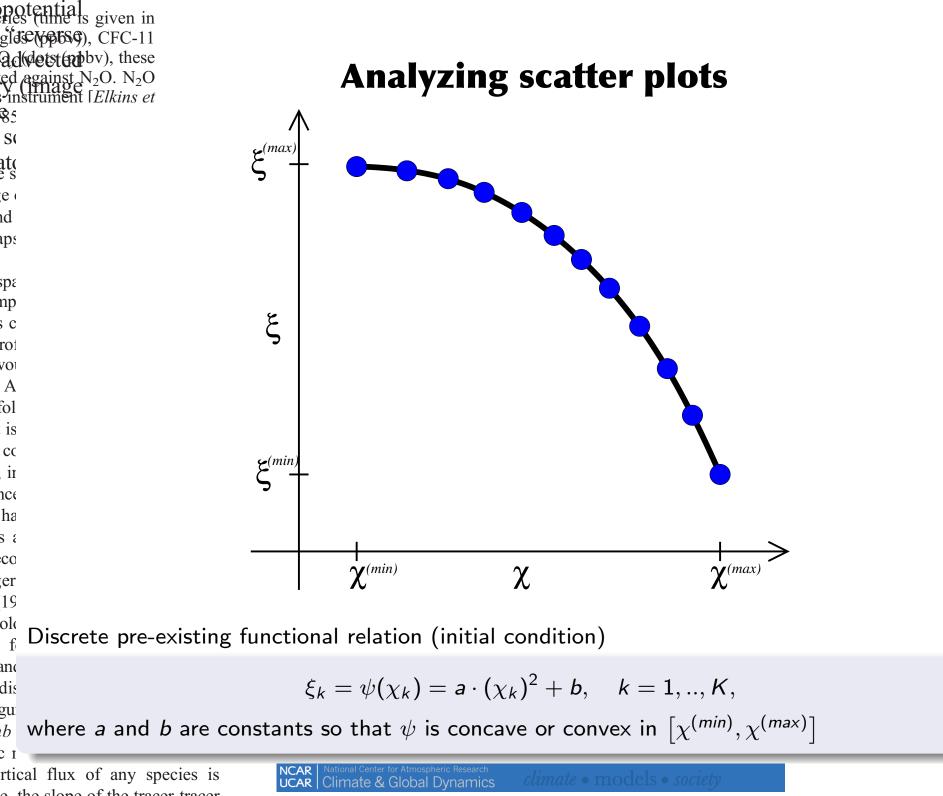
(right) same as (left) but using a fixer:

- (i) transport the individual species
- (ii) transport the total
- in each grid cell scale the individual species by the difference between (i) and (ii)

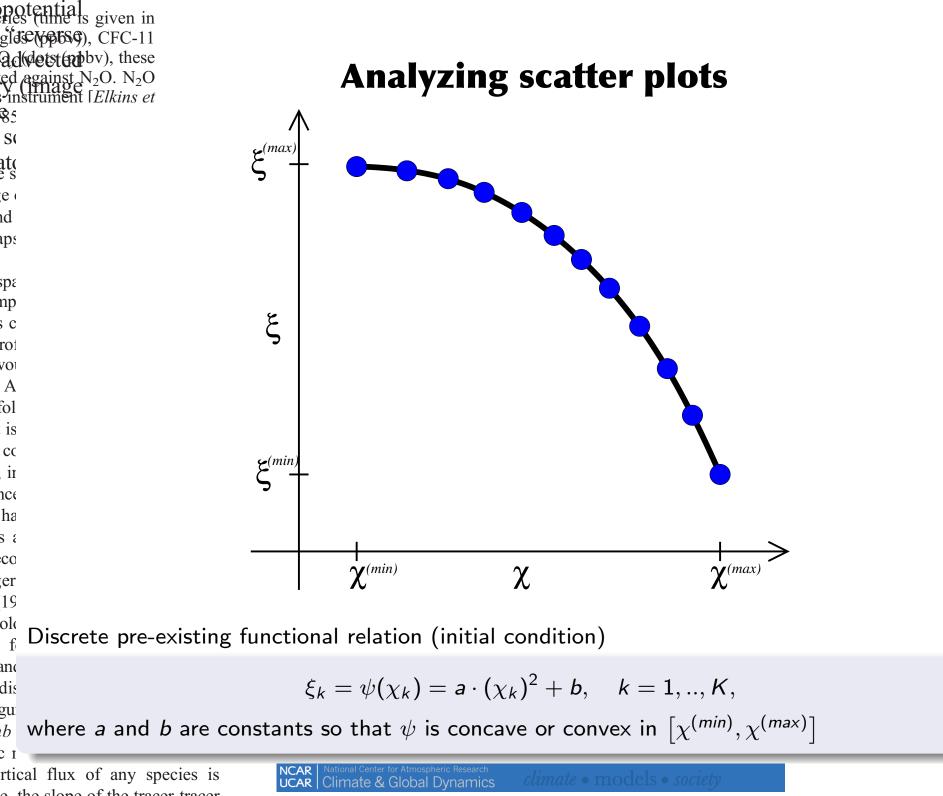




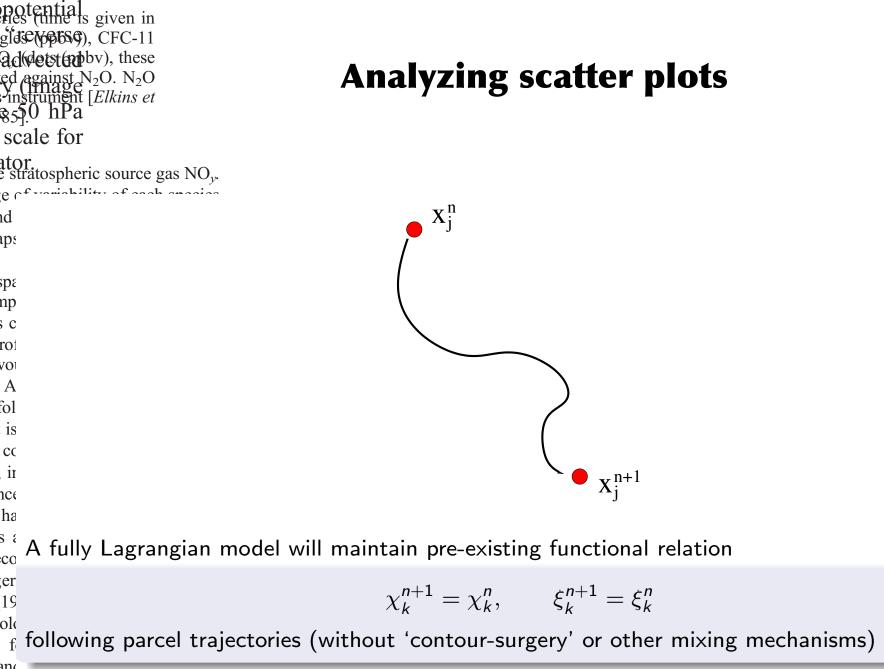
e, the slope of the tracer-tracer (1, 1, (2), (1), (1))



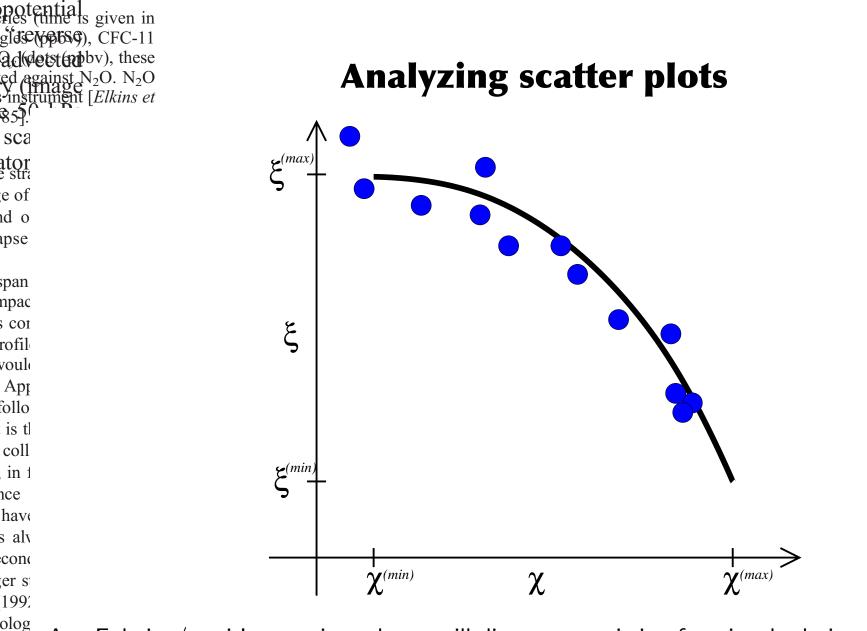
e, the slope of the tracer-tracer



e, the slope of the tracer-tracer



displacement. Building on the guments of *Holton* [1986] and *b and Ko* [1992] also showed c mixing extends globally, the rtical flux of any species is e, the slope of the tracer-tracer





$$\xi_k^{n+1} = \mathcal{T}(\xi_j^n) \neq \mathbf{a} \cdot \mathcal{T}\left(\chi_j^n\right)^2 + \mathbf{b}, \quad j \in \mathcal{H}$$

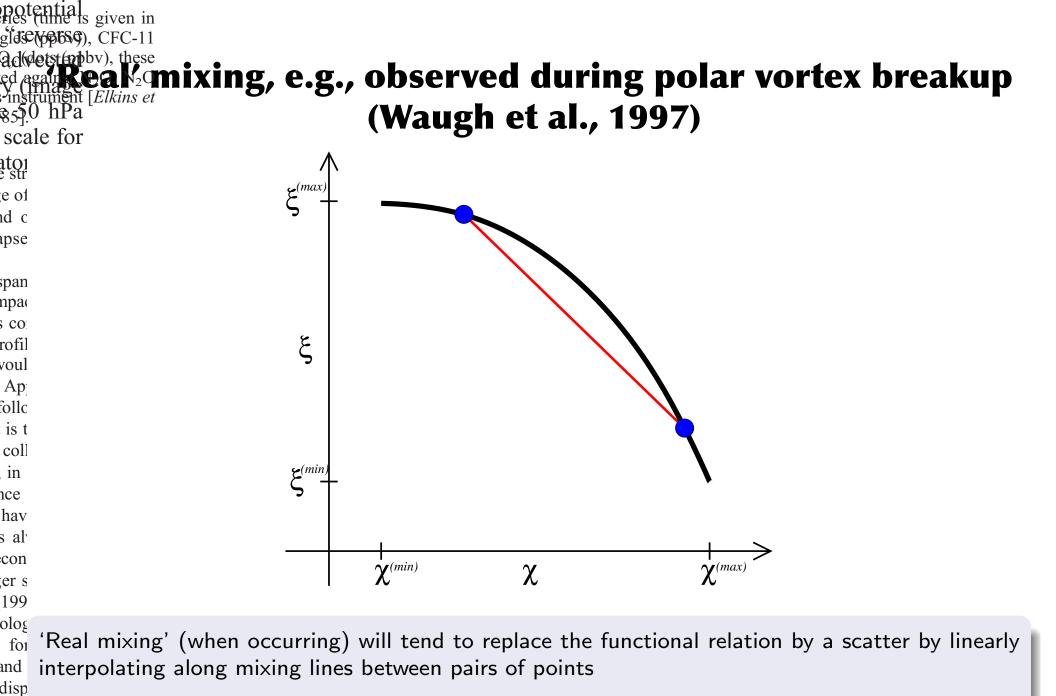
where \mathcal{T} is the transport operator and \mathcal{H} the set of indices defining the 'halo' for \mathcal{T} .

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rtical flux of any species is e, the slope of the tracer-tracer

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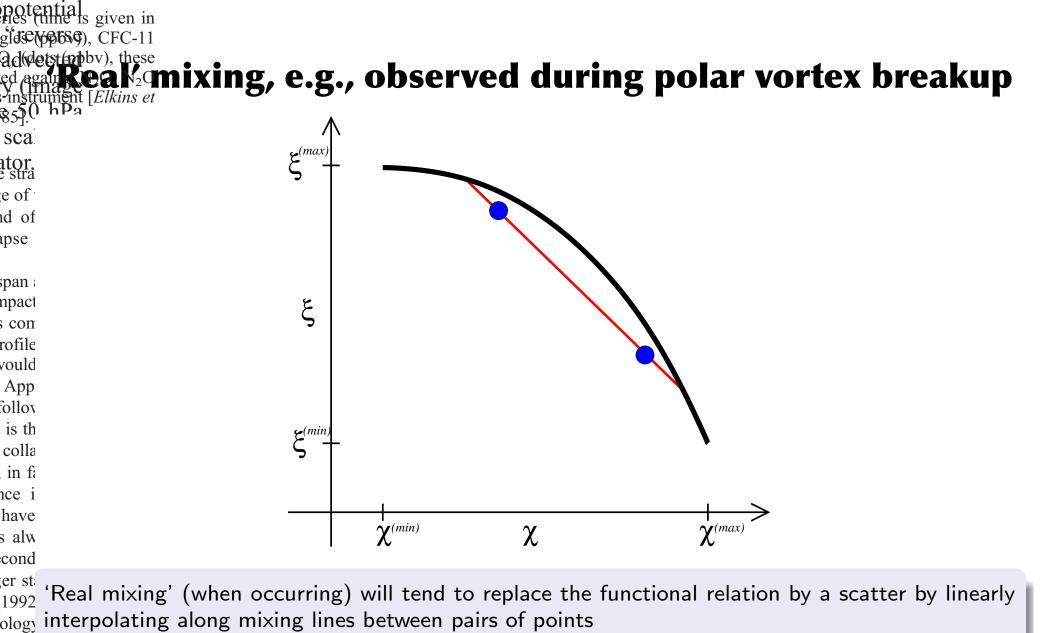
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b and Ko [1992] also showed c mixing extends globally, the rtical flux of any species is e, the slope of the tracer-tracer

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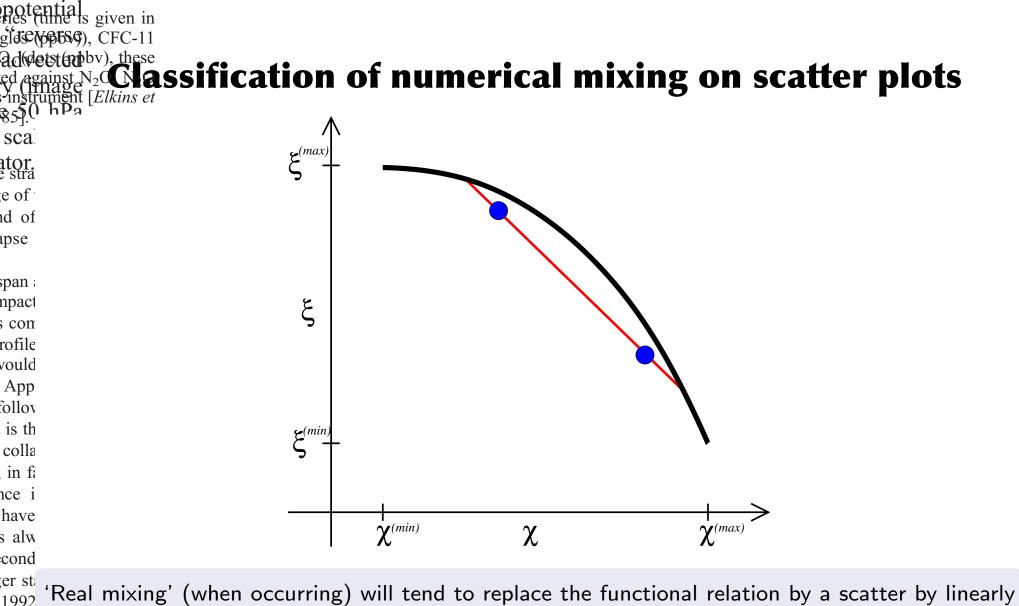
for \rightarrow Ideally numerical mixing should = 'real mixing'!

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gume However, it may be shown mathematically that schemes that exclusively introduce 'real ab an mixing' are 1st-order schemes (Thuburn and McIntyre, 1997).

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rtical flux of any species is e, the slope of the tracer-tracer



ology interpolating along mixing lines between pairs of points

```
for \rightarrow Ideally numerical mixing should = 'real mixing'!
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and t displa

gume However, it may be shown mathematically that schemes that exclusively introduce 'real ab an mixing' are 1st-order schemes (Thuburn and McIntyre, 1997).

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rtical flux of any species is e, the slope of the tracer-tracer

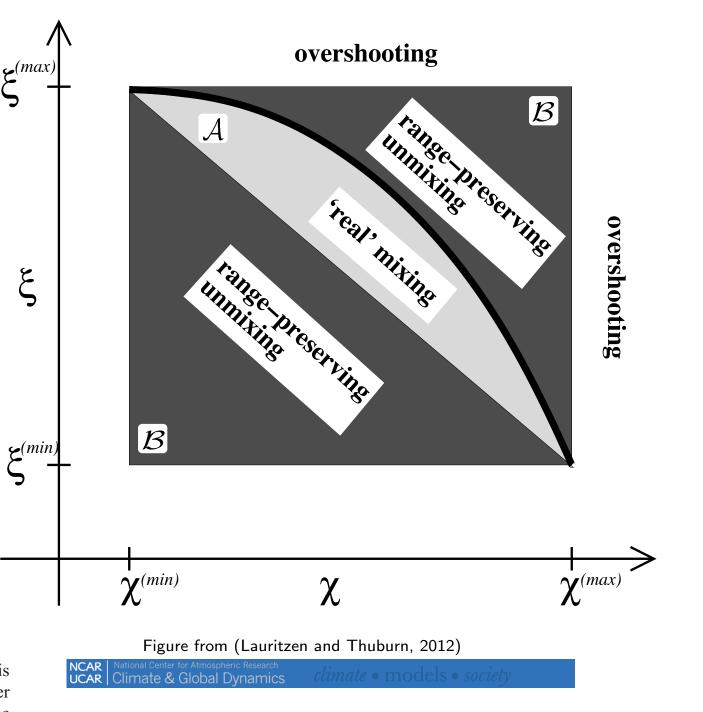
gias (under is given in gias (under is given in adversed against N₂ Classification of numerical mixing on scatter plots instrument [Elkins et 353.0 hPa

e stratospheric source gas N e of variability of each spec ad of latitude covered by apse to remarkably comp span a range of latitudes and

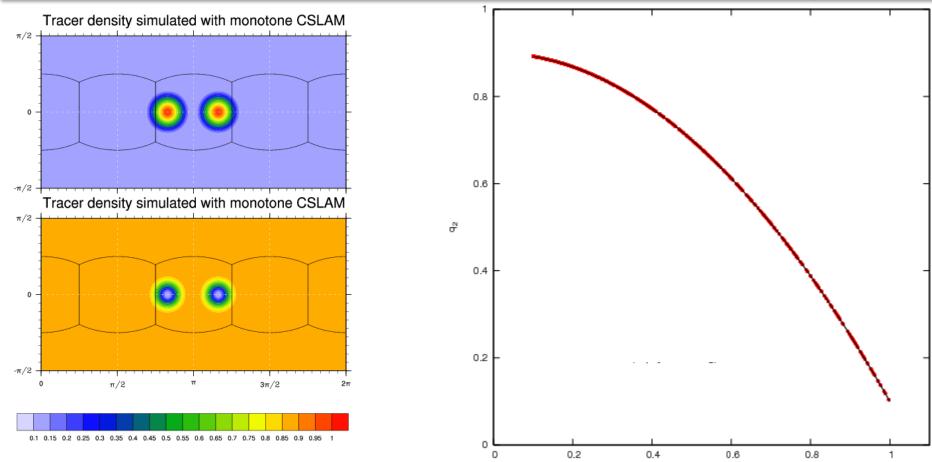
scale for

npactness would be much l s comprised vertical profile rofiles on a surface of const yould then be a function c Apparently compact function follow from a simple change is that data from near-vert collapse in tracer-tracer sp in fact, another manifestat nce if the isosurfaces of have the same shape, a gir s always accompanied by cond. In fact, the more lc er statement than the clima 1992] argued that if compa ology, it is present on sho for these long-lived trace and their relationship is the displacement. Building on guments of Holton [1986] : b and Ko [1992] also show c mixing extends globally, rtical flux of any species is

e, the slope of the tracer-tracer



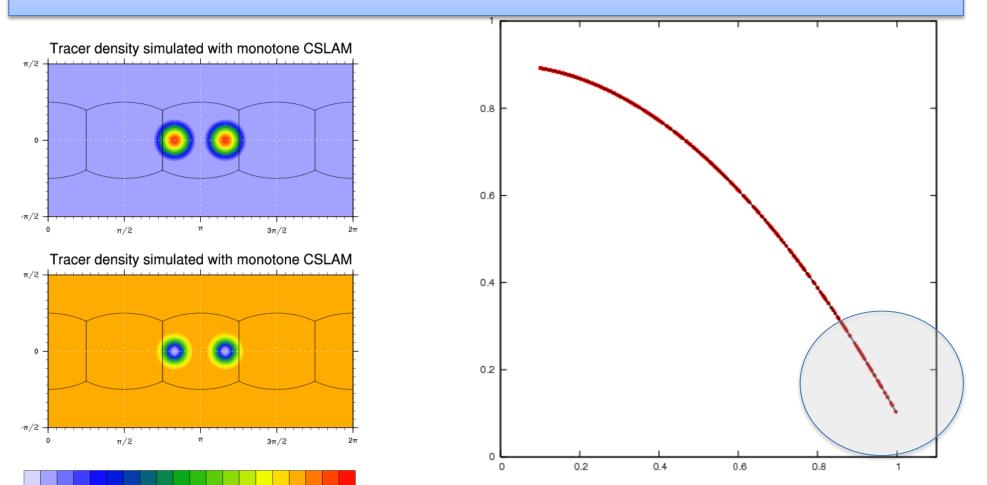
First-order scheme: only `real mixing'



Nair and Lauritzen (2010) flow field

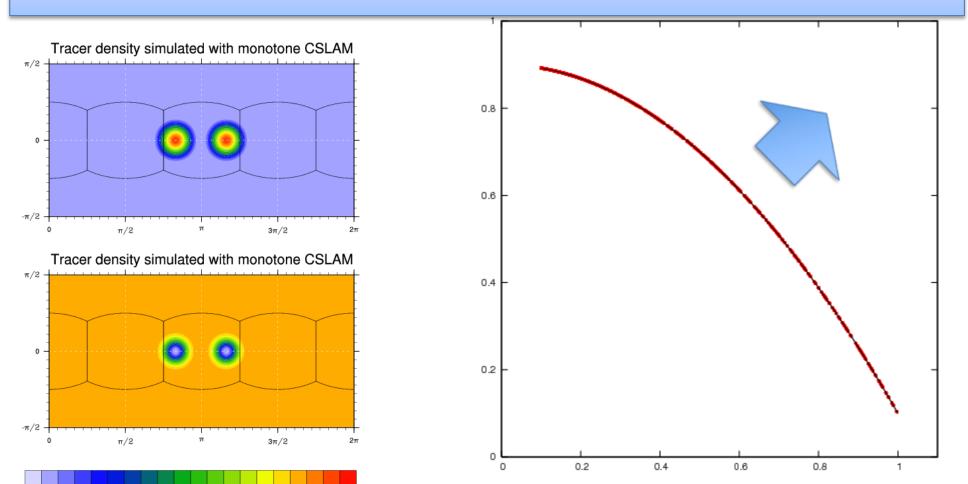
91

Note: 1. Max value decrease, 2. Unmixing even if scheme is shapepreserving, 3. No expanding range unmixing



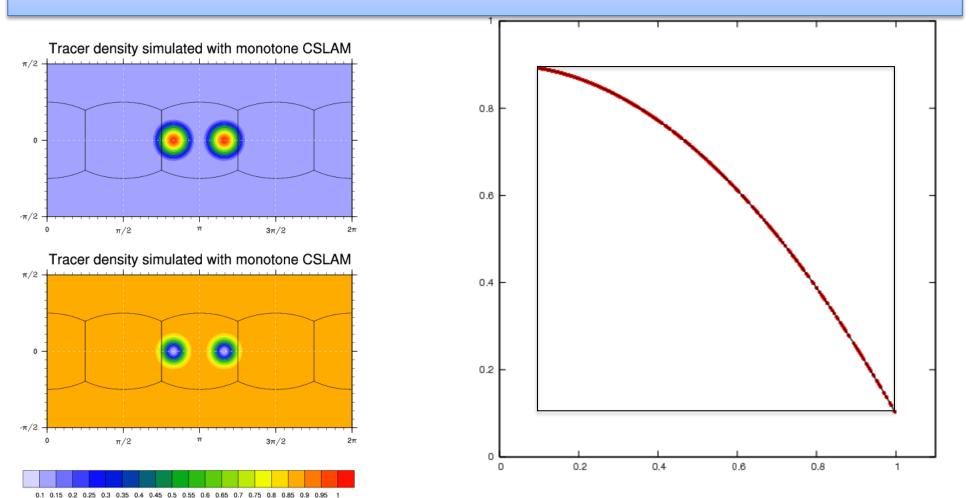
0.1 0.15 0.2 0.25 0.3 0.35 0.4 0.45 0.5 0.55 0.6 0.65 0.7 0.75 0.8 0.85 0.9 0.95

Note: 1. Max value decrease, 2. Unmixing even if scheme is shapepreserving, 3. No expanding range unmixing



0.1 0.15 0.2 0.25 0.3 0.35 0.4 0.45 0.5 0.55 0.6 0.65 0.7 0.75 0.8 0.85 0.9 0.95

Note: 1. Max value decrease, 2. Unmixing even if scheme is shapepreserving, 3. No expanding range unmixing



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Summary

- Inherent mass-conservation is important
- Shape-preservation is important
- Filament preservation is important
- Consistency (mass-wind) is important
- Correlation preservation (linear, non-linear) is important

I have shown you several idealized test cases to assess these aspects of transport schemes



0 More information: <u>http://www.cgd.ucar.edu/cms/pel</u> Email: pel@ucar.edu NCAR | National Center for Atmospheric Research UCAR | Climate & Global Dynamics