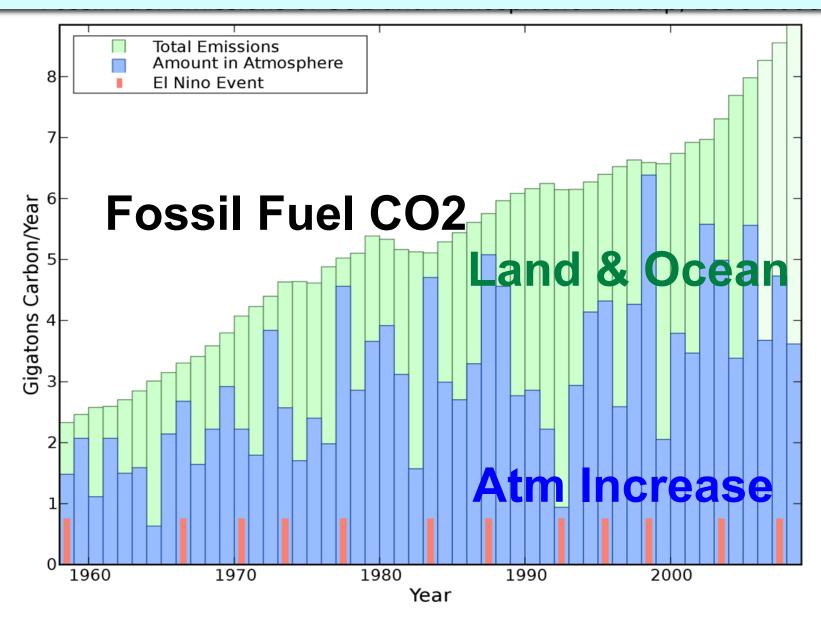


Challenge: Co-variation of Climate and Carbon Sinks



Focus of Coupled Carbon Climate Model

- Changes in CO2 and $\frac{\partial CO2}{\partial t}$ that are radiatively important
- Time scale of investigation:
 - Eqm: e.g. Last Glacial Max; 2xCO2
 - Transient: modern (now +/- several centuries). Leave out geology (and interactions with geology) if using complex Earth System Model.

Units: $1 Pg = 1 Gt = 10^{15} gram$

 $C(kgC / m^{3}) = \rho(kgAir / m^{3}) \times X(moleC / moleAir) \times (MWt_{C} / MWt_{air})$ $MWt_{C} = 12gm / mole; MWt_{air} = 29gm / mole$

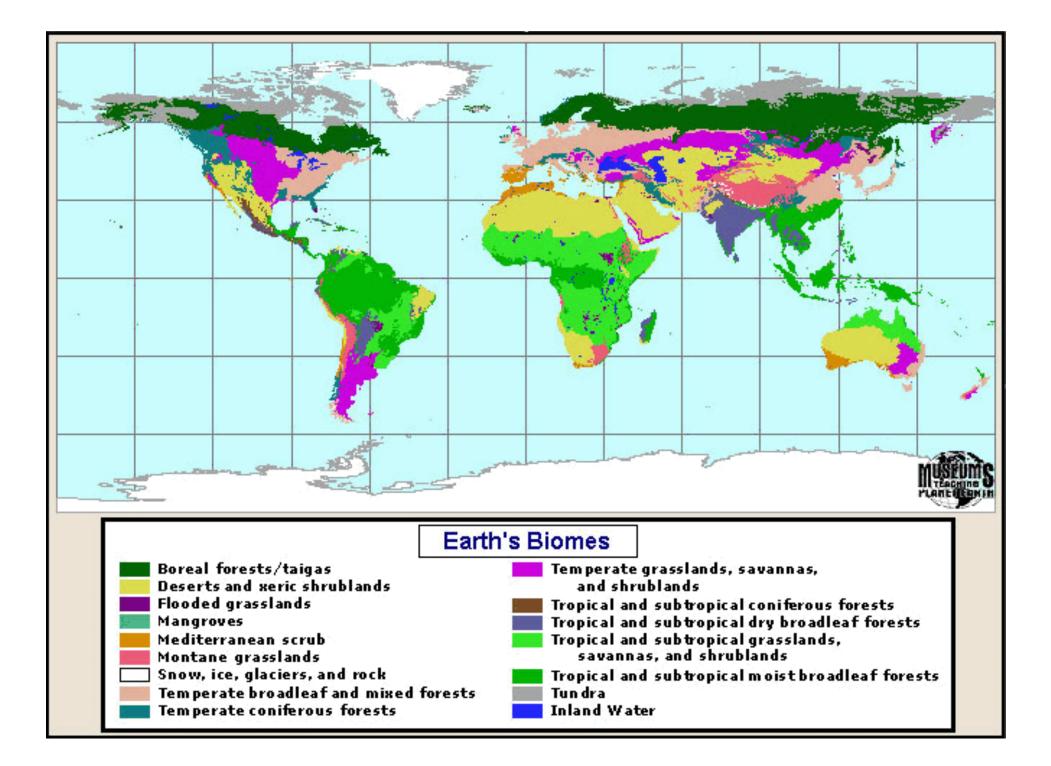
Area =
$$\int dx \, dy \, dz = 5 \times 10^{14} (m^2)$$

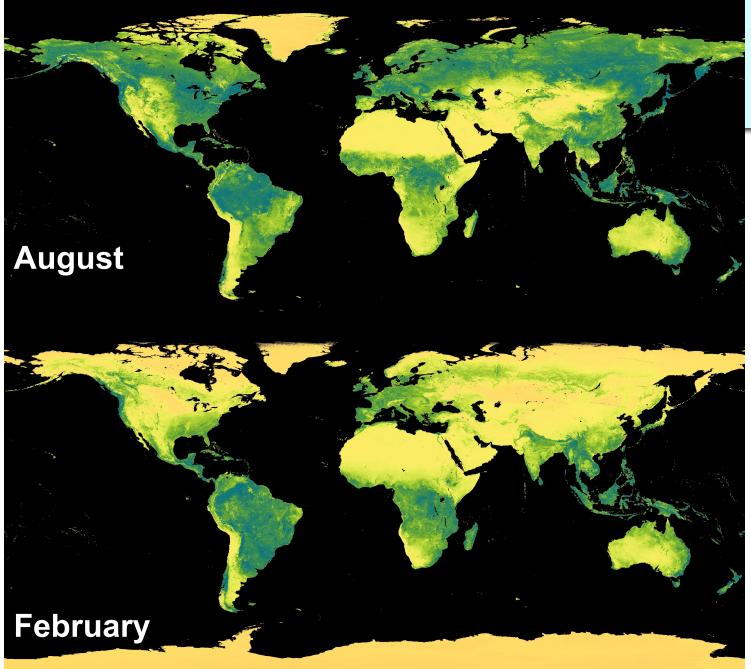
MassAtm = $\int \rho \, dx \, dy \, dz = \frac{100 P_{mb}}{g} (kgAir / m^2) \times Area \sim 5 \times 10^{18} kgAir$

 $MassC(300 \, ppmv) = MassAtm \times (300 \, x10^{-6}) \times (\frac{12}{29})$

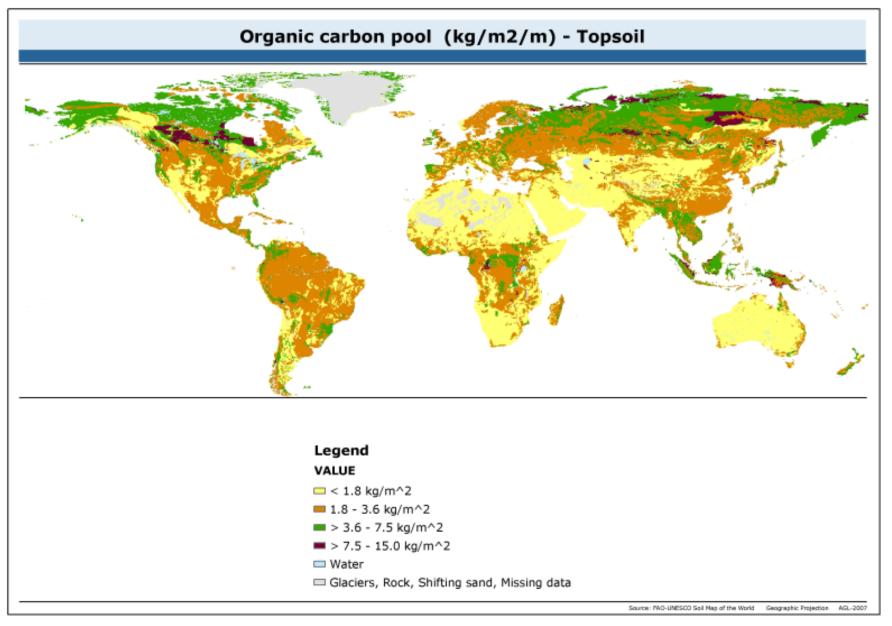
 $\sim 600 \times 10^{12} kg = 600 PgC = 600 GtC$ 1PgC $\rightarrow 0.5 ppmv(mixed _entireAtm)$

LAND BIOSPHERE





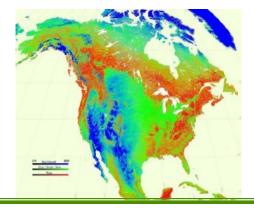
Satellite Greenness index: NDVI

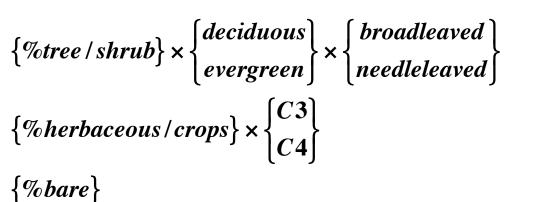


High soil carbon found when there is abundant organic matter input, and slow decomposition (cold, not well aerated).

(1) Who: Name Calling

- 1. Taxa/Species
- 2. Physionomic/structural: Biomes e.g. desert, rainforest albedo, roughness
- 3. Bioclimatic + CO2 +...
- 4. **Functional:** based on satellite photosynthesis index:

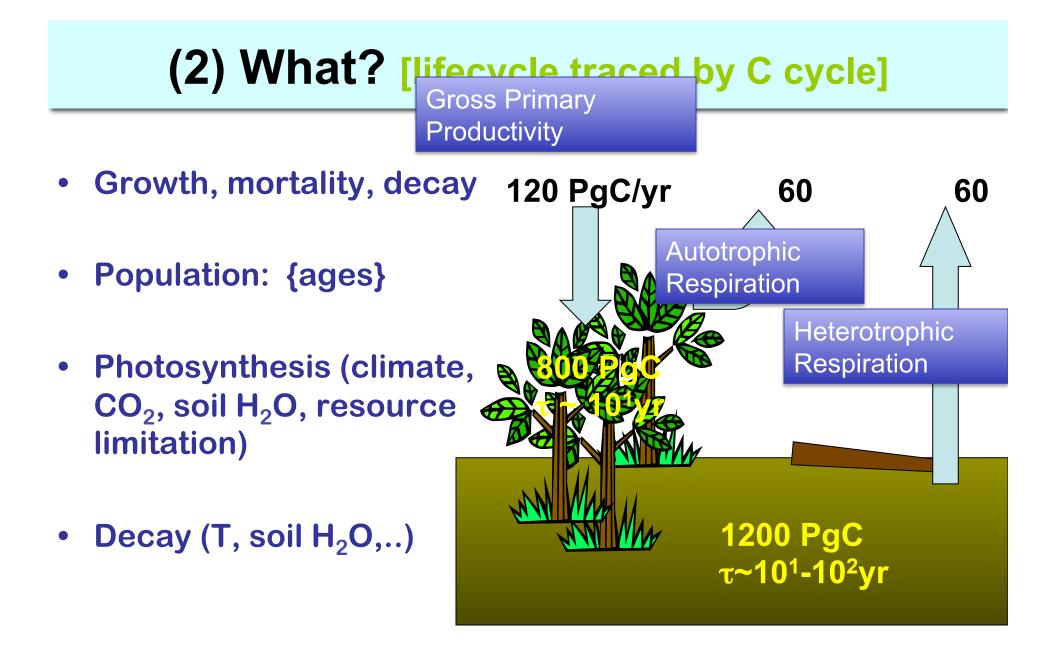




Mean annual temperature (°C)

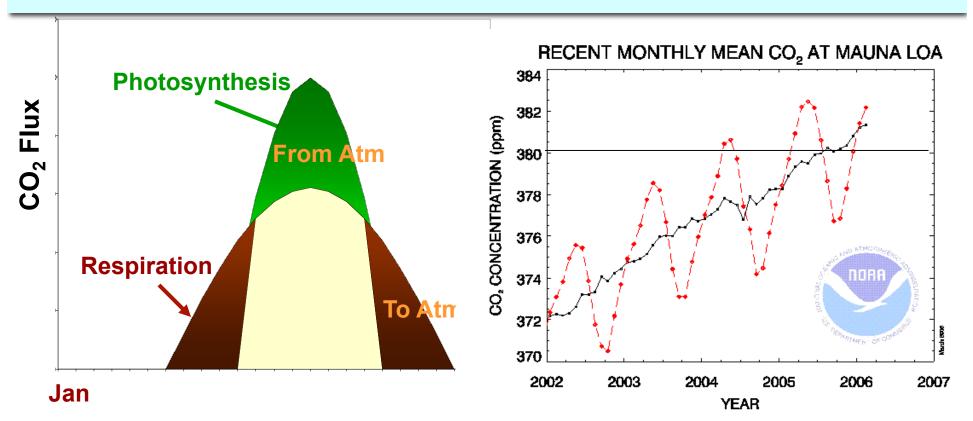
Subtropical Warm temperate Cold temperate Arctic-alpine

Static land cover: typically 4 --> lookup table of structural properties and BGC chacteristics. Prognostic land cover: 3+



Net Primary Production (NPP) = GPP – R_auto

Two major fluxes - seasonal cycle

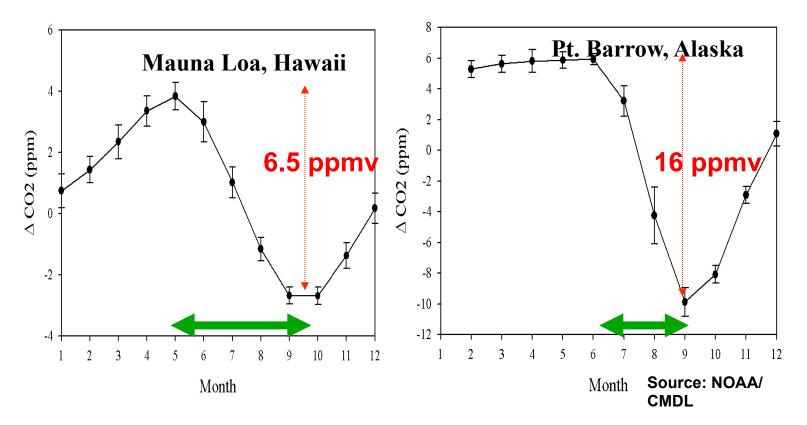


- Seasonal asynchrony between photosynthesis and decomposition
- \rightarrow net fluxes of CO₂ to and from atm
- \rightarrow seasonal cycle of CO₂ in atm
- Annual imbalance → carbon source/sink

Fung et al JGR 1983, 1987:

Simple prescriptions of seasonality. 1987 based on satellite obs + temperature

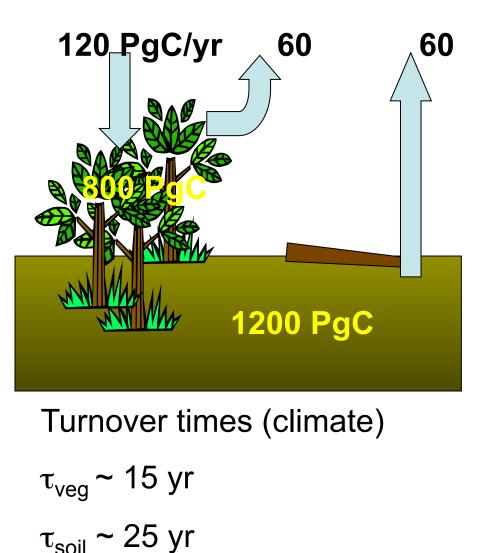
Atmospheric CO₂ Signature of Ecosystem C Exchange: Seasonal Cycle

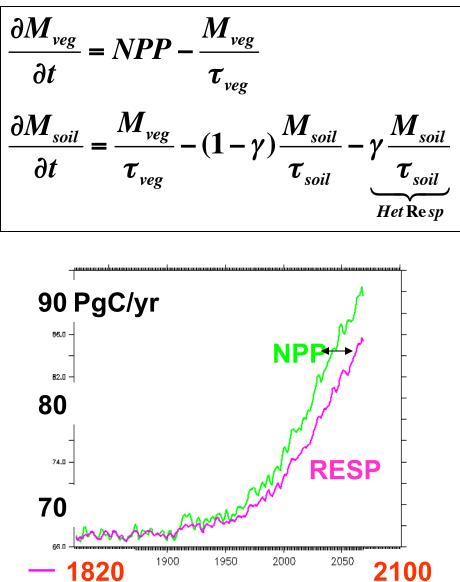


• Amplitude of atmospheric CO₂ seasonal cycle increases poleward: telecoping of growing season and greater asynchroneity bet' fluxes

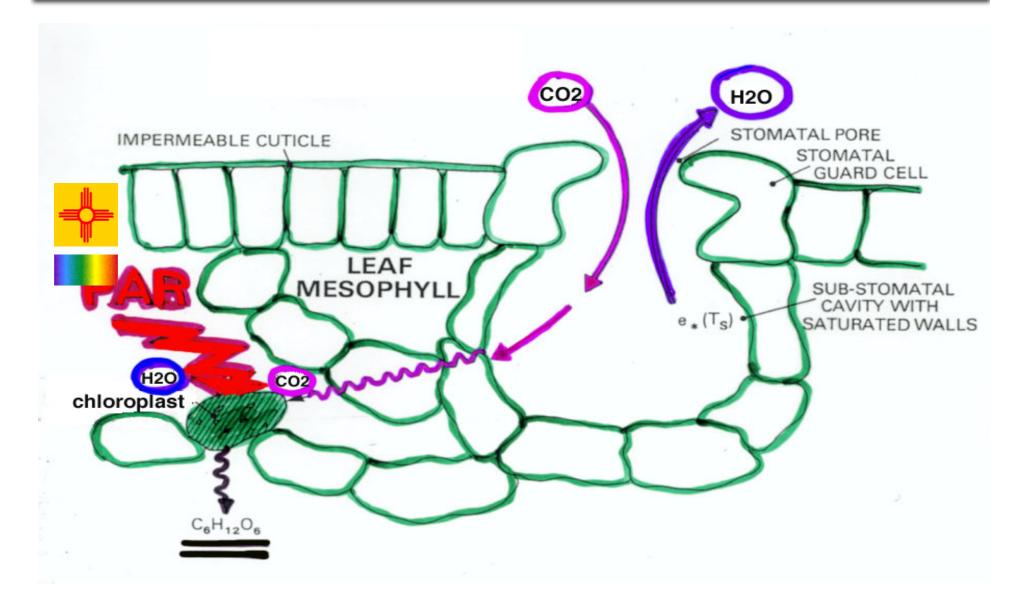
• Growing season net flux ~15-20% of annual NPP

(3.1) How? First simple toy model: carbon only

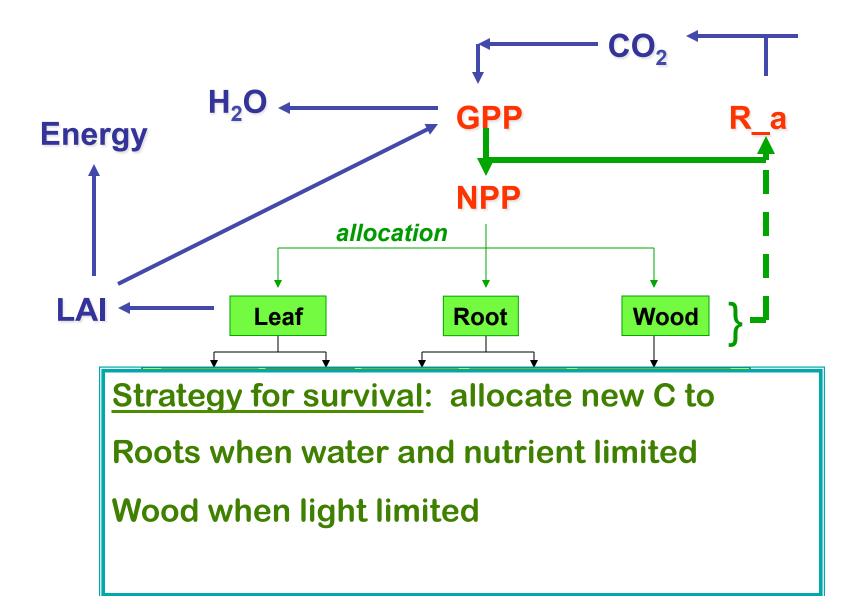




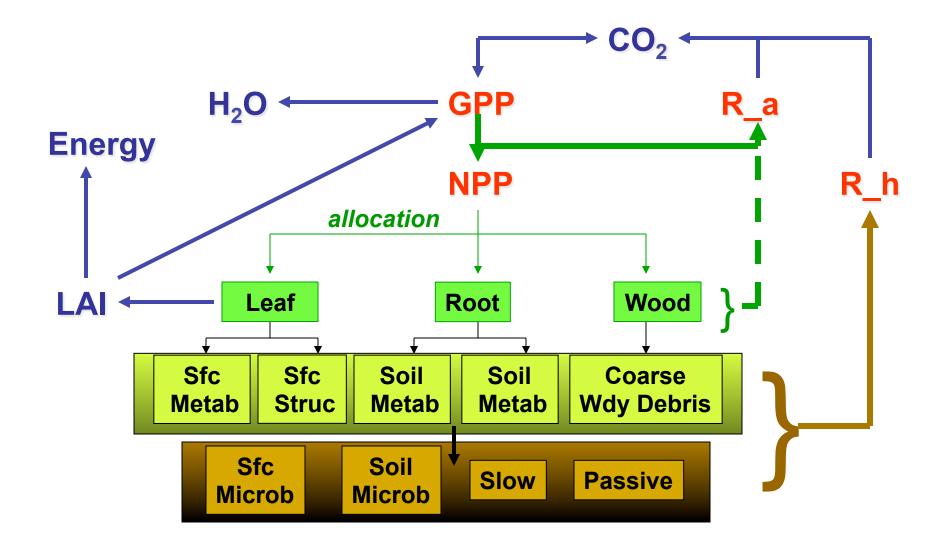
(3.2.1) Leaf Photosynthesis

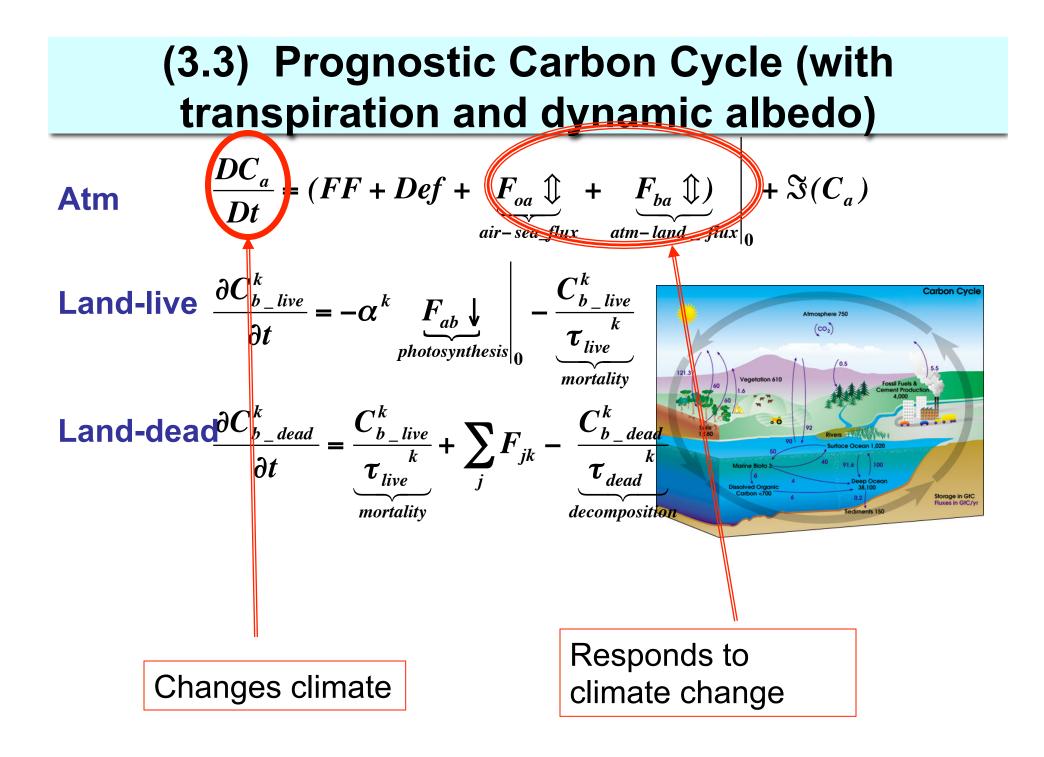


(3.2.2) Photosynthesis --> Growth

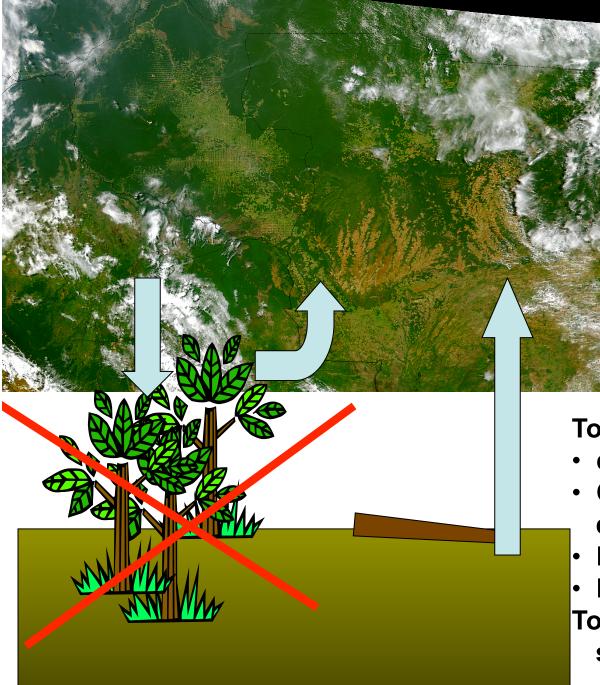


(3.2.3) Photosynthesis, Growth, Mortality & Decay

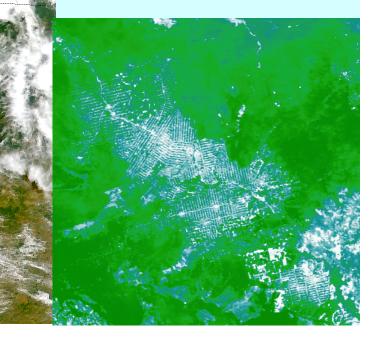




DEFORESTATION



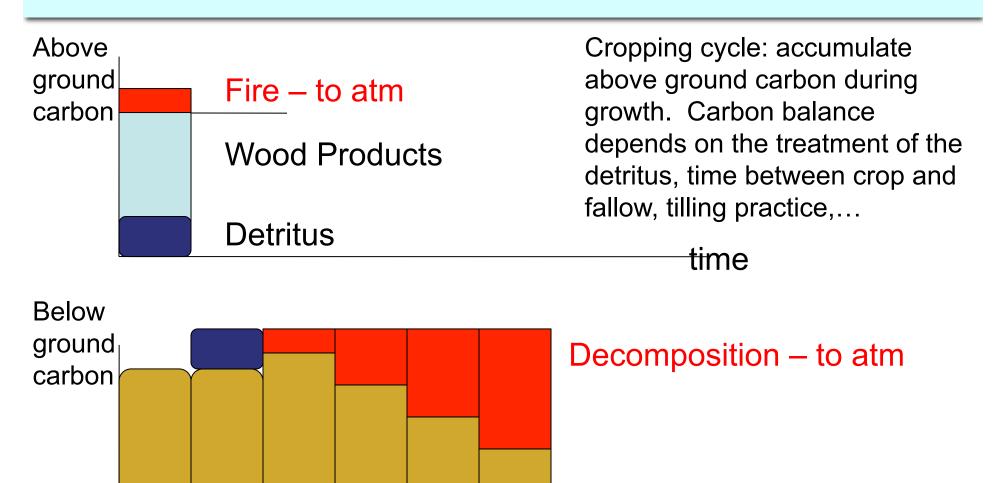
Deforestation



Tough to estimate

- deforested area
- Carbon inventory before deforestation
- Fate of removed carbon
- Fate of litter and soil carbon
 Tough to discriminate atm CO2
 signature

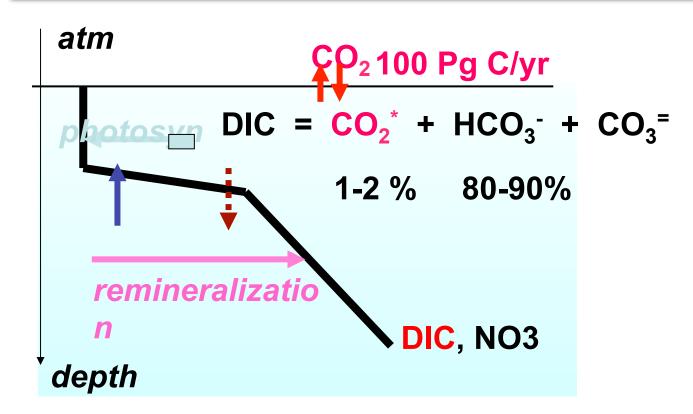
Fate of Carbon



time

OCEAN CARBON

Ocean C from the Atm's Perspective:



$$F_{oa} - F_{ao} = \underbrace{k}_{GasExchRate(m/s)} \times (\underbrace{CO2*_{sfc_ocn}}_{moleC/m3} - \underbrace{\beta}_{so\,lub\,ility(T)} \times pCO2_{atm_sfc})$$

(1) Simplified Carbonate Chemistry

Chemical Equilibrium

 $CO_2+H_2O \leftrightarrow H^+ + HCO_3^ CO_3^= + H^+ \leftrightarrow HCO_3^-$

 $K(T,S) = \frac{[HCO3][HCO3]}{[CO2][CO3]}$

Mass and Charge Balance

$\Sigma CO2 = TC = DIC = CO_2 + HCO_3 + CO_3^{=}$

CO2 increases with increasing DIC

$ALK = HCO_3^- + 2[CO_3^-]$

CO2 decreases with increasing ALK

Henry's Law

 $\mathbf{pCO}_2 = \mathbf{k}_{\mathrm{H}} * \mathbf{CO}_2$

pCO₂: partial pressure (CO2 gas), in atm

Atm View: Summary of Carbonate Chemistry

$$\frac{\partial}{\partial t}(h \cdot DIC) = k \cdot solub \cdot pCO2_a - k CO2_{ml} + Prod_bio - Loss_bio$$

$$\frac{1}{CO2^*}d(CO2^*) = d(\ln(CO2^*))$$

$$= \frac{d\ln CO2^*}{d\ln DIC} d\ln DIC \quad \text{Revelle factor}$$

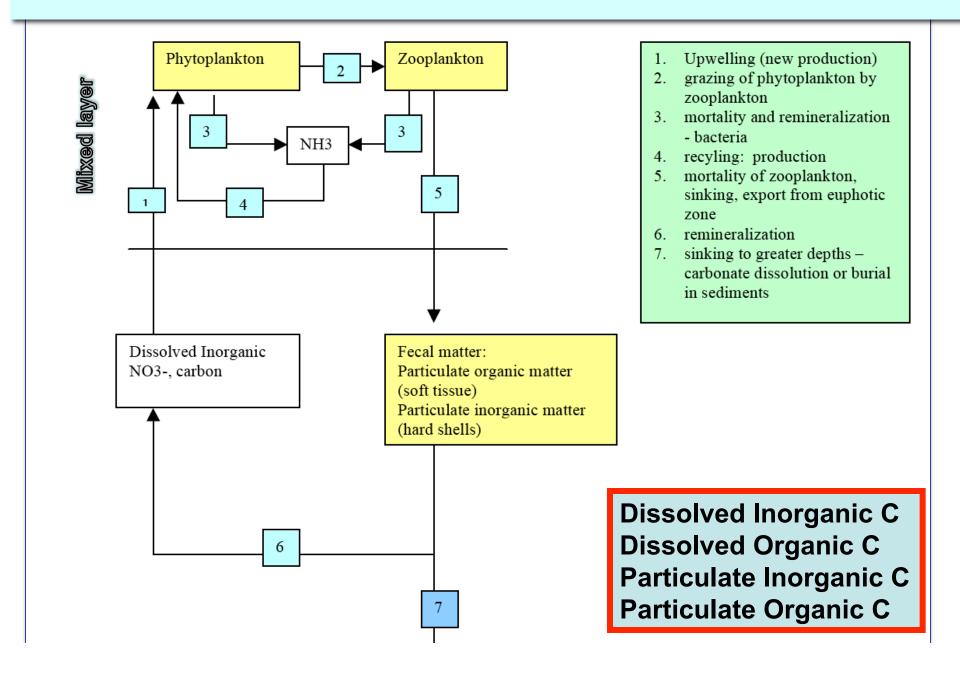
$$+ \frac{d\ln CO2^*}{d\ln TALK} d\ln TALK$$

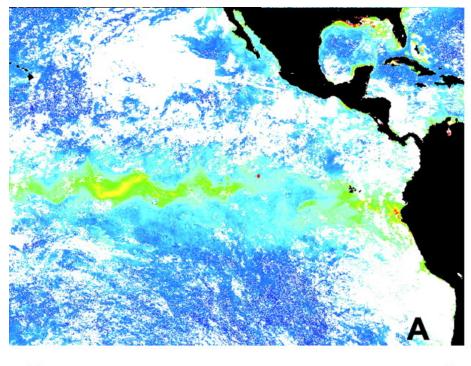
$$\sim -9$$

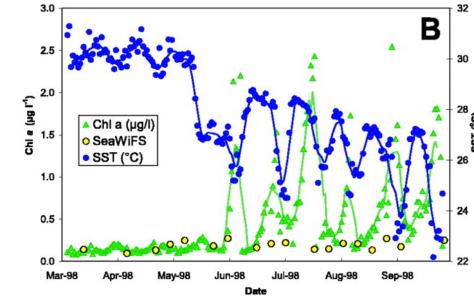
$$+ \frac{d\ln CO2^*}{dT} dT \quad \text{Solubility Pump}$$

$$+ \frac{d\ln CO2^*}{dS} dS$$

(2.1) Marine Biota – Biological Carbon Pump







Marine Productivity

Is possible when upwelling brings:

Nutrients from below to euphotic zone
Cold water

²⁸ Small Flux, small inventory
 ²⁶ of organic C
 ²⁷ But alters DIC(z)

Biological C Pump

Redfield ratio = C:N:P in organic matter ~103: 16: 1

		DIC	TALK	[CO2]
Formation of POC (surface ocean)		- 1 unit	+ (16/103)*1 (Redfield * charge)	•
Formation of CaCO3 (surface)	Ca ⁺⁺ + CO ₃ ⁼ → CaCO3	- 1 unit	- 2 units	^
Remineraliza tion of organic matter		+ 1 unit	- 16/103	^
Dissolution of shells		+ 1 unit	+ 2	¥

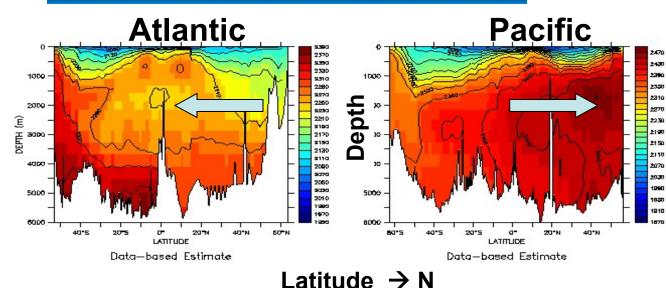
The Global Conveyor Belt

(2.3) OceanC : Mainly Dissolved Inorganic Carbon (DIC)

Conveyor Belt Transport of DIC:

•Southward in Atlantic •Northward in Pacific

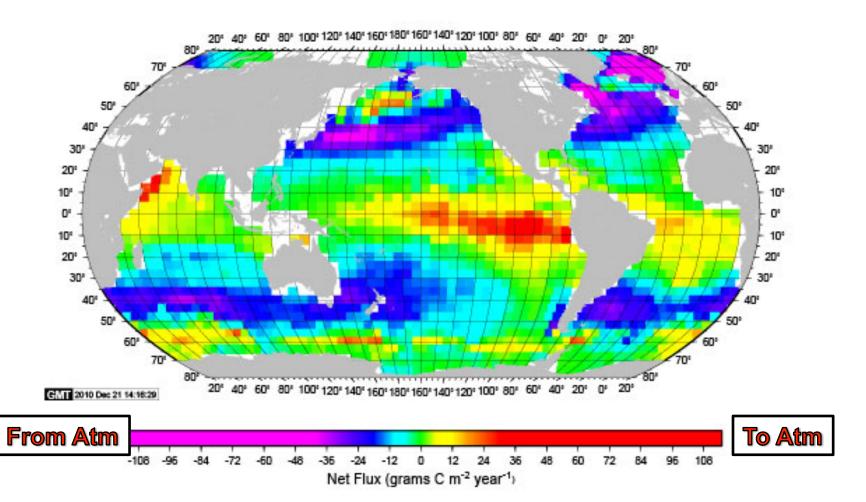
Ocn currents ~ cm/s Time scale ~ 10^3 yr



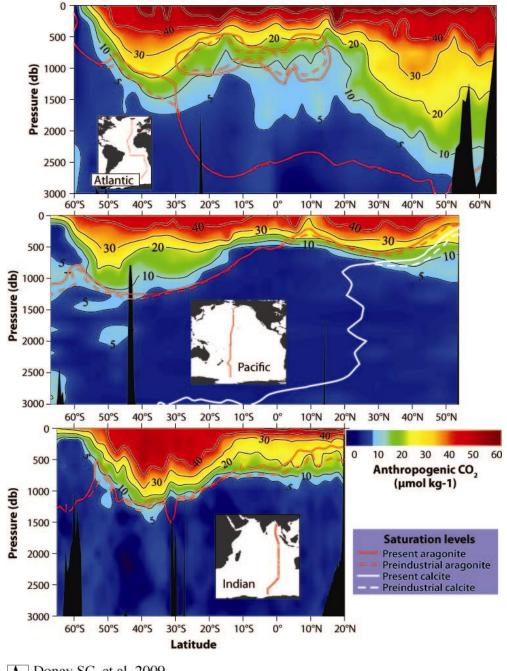
Biology and DIC: •Depletion near sfc •Enrichment at Depth

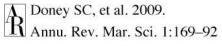
First barrier to Ocean Uptake: Air-Sea CO2 Flux

Mean Annual Air-Sea Flux for 2000 [Rev Dec 10] (NCEP II Wind, 3,040K, Γ=.26)

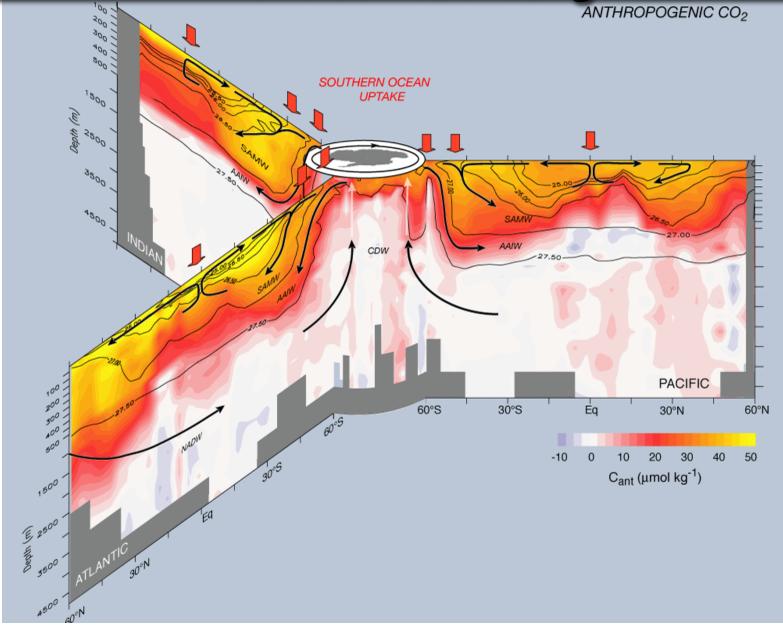


http://www.ldeo.columbia.edu/res/pi/CO2/carbondioxide/pages/air_sea_flux_2010.html





Second Barrier to Ocean Uptake: Thermocline Mixing



SUMMARY

Prognostic Carbon Cycle

$$\begin{aligned} \operatorname{Atm} \qquad & \left. \frac{\partial C_a}{\partial t} + \frac{\partial}{\partial x} (u_a C_a) + \ldots = (FF + Def + \underbrace{F_{oa} \bigoplus_{\operatorname{air-sea} \operatorname{flux}}}_{\operatorname{air-land}_{\operatorname{flux}}} \left| \underbrace{flux}_{atm-land_{\operatorname{flux}}} \right|_0^1 + \underbrace{\mathfrak{I}_{\operatorname{trbulent}_{\operatorname{mixing}}}}_{\operatorname{turbulent_{mixing}}} \\ \operatorname{Ocean} \qquad & \left. \frac{\partial C_o}{\partial t} + \frac{\partial}{\partial x} (u_o C_o) + \ldots = -F_{oa} \oiint_0^1 + \underbrace{P - L}_{\operatorname{biology}} + \mathfrak{I}(C_o) \\ \operatorname{Land-live} \qquad & \left. \frac{\partial C_{b_live}}{\partial t} \right|_0^1 = -\alpha^k \underbrace{F_{ab} \downarrow}_{photosynthesis} \right|_0^1 - \underbrace{\frac{C_{b_live}}{\tau_{live}^k}}_{mortality}; \quad \sum_k \alpha^k = 1 \\ \operatorname{Land-dead} \qquad & \left. \frac{\partial C_{b_dead}}{\partial t} \right|_0^1 = \underbrace{\frac{C_{b_live}}{\tau_{live}^k}}_{photosynthesis} + \sum_j F_{jk} - \underbrace{(1 - \gamma^k) \frac{C_{b_dead}}{\tau_{dead}^k}}_{decomposition} - \gamma^k \frac{C_{b_dead}}{\tau_{dead}^k} \\ & \sum_k \gamma^k \frac{C_{b_dead}^k}{\tau_{dead}^k} = F_{ba} \end{aligned}$$