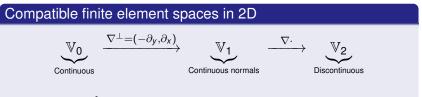
Mimetic finite elements

Mimetic properties

Discrete versions of calculus identities: $\nabla \cdot \nabla^{\perp} \psi = 0, \nabla^{\perp} \cdot \nabla D = 0, \nabla^{\perp} \cdot \boldsymbol{u}^{\perp} = \nabla \cdot \boldsymbol{u}, \text{ where } (\nabla^{\perp} = \boldsymbol{k} \times \nabla).$



- $\nabla \cdot$ maps from \mathbb{V}_1 onto \mathbb{V}_2 .
- ∇^{\perp} maps from \mathbb{V}_0 onto ker $(\nabla \cdot)$ in \mathbb{V}_1 .

Applied to the linearised shallow water equations:

- 1) global energy conservation 2) local mass conservation
- 3) steady geostrophic states 4) no spurious pressure modes

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Application to nonlinear shallow water equations

$$\boldsymbol{u}_t + (\underbrace{\boldsymbol{u}qD}_{\boldsymbol{q}})^{\perp} = -\nabla \left(\frac{1}{2}|\boldsymbol{u}|^2 + gD\right)$$
 (1) $D_t + \nabla (\underbrace{\boldsymbol{u}D}_{\boldsymbol{F}}) = 0$ (2)

where $q = rac{\zeta + f}{D}$ and $\zeta = \nabla^{\perp} \boldsymbol{u}$

 $abla^{\perp}(1) ext{ gives } (qD)_t +
abla \cdot \boldsymbol{Q} = 0$ (3)

Timestepping:

- within each timestep perform multiple quasi Newton iterations: solve the Helmholtz equation for updates to u and D using a hybridized method
- first need to calculate F and Q:
 - solve equation 2 for *D* and calculate mass flux *F*: flux reconstruction
 - solve equation 3 for *q* and diagnose potential vorticity flux
 Q: Taylor-Galerkin methods

Mimetic finite element methods for solving the nonlinear rotating shallow water equations.

Motivation

- Construct numerical schemes that have all the desirable properties of the C-grid finite difference discretisation, without the constraint that the grid be orthogonal
- Devise stable, consistent advection schemes for both layer depth (discontinuous) and potential vorticity (continuous) fields
- Present benchmarking results from standard testcases

Mimetic properties

Discrete analogues of the vector calculus identities

 $\nabla \cdot \nabla^{\perp} \psi \equiv \mathbf{0}, \quad \nabla^{\perp} \cdot \nabla \psi \equiv \mathbf{0}, \quad \nabla^{\perp} \cdot \mathbf{u}^{\perp} \equiv \nabla \cdot \mathbf{u}$ where $\nabla^{\perp} = \mathbf{k} \times \nabla$ and $\mathbf{u}^{\perp} = \mathbf{k} \times \mathbf{u}$ where \mathbf{k} is the normal to the surface.

Applied to the linearised shallow water equations [1]:

satisfy LBB condition

- steady geostrophic modes
- global energy conservation
- no spurious pressure modes
- Iocal mass conservation

Shallow water equations

 $u_t + (\zeta + f)u^{\perp} + \nabla \left(g(D+b) + \frac{1}{2} |u|^2 \right) = 0$ (1) $D_t + \nabla \cdot (uD) = 0$ (2) *u*: horizontal velocity, *D*: layer depth, *b*: height of the lower boundary, *f*: Coriolis parameter,

 ζ : vorticity (= $\nabla^{\perp} \boldsymbol{u}$), g: gravitational acceleration. Combining equation 2 with ∇^{\perp} equation 1 gives Lagrangian conservation of potential

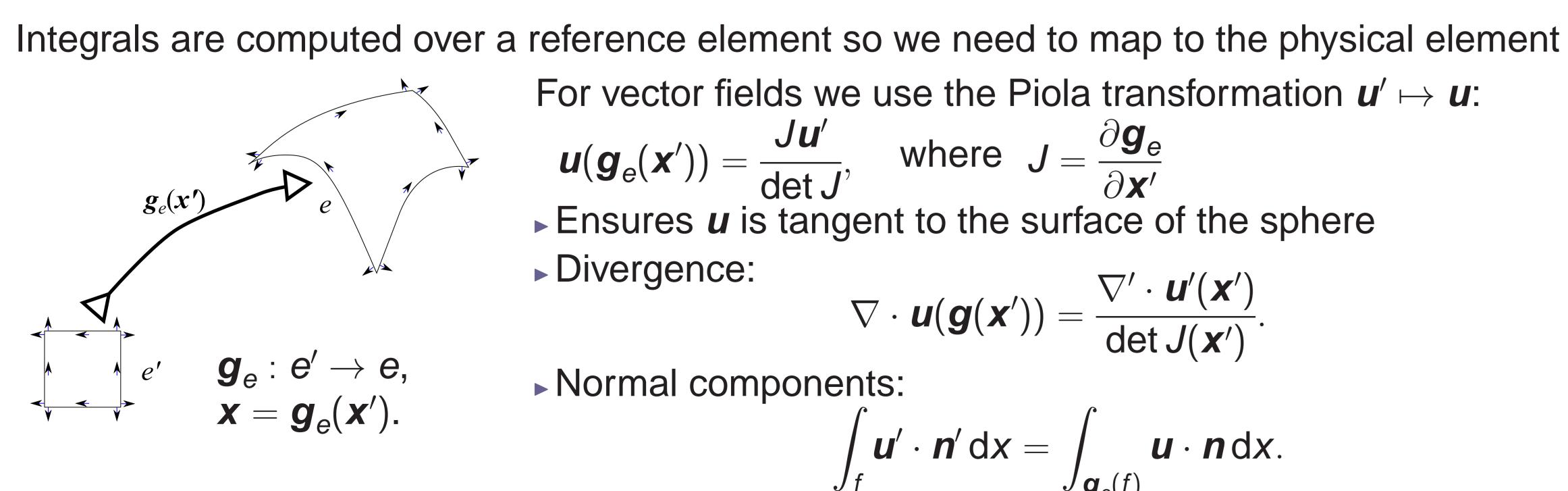
vorticity, q: $q_t + (\boldsymbol{u} \cdot \nabla)q = 0$, where $q = \frac{\zeta + f}{D}$

Compatible finite element spaces

$ \begin{array}{cccccccccccccccccccccccccccccccccccc$	
uadratic (+1 Cubic) Continuous Linear (+2 Quadratic) Cont. normals ∇^{\perp}	Vorticity
$\mathbb{V}_0 = \mathbb{Q}_2 \longrightarrow \mathbb{V}_1 = \mathbb{R}_1 \mathbb{Q}_1$	$\xrightarrow{\text{v}} \mathbb{W}_2 = Q1_{DG} \qquad \bullet \bullet$
Biquadratic Continuous Bilinear/Biquadratic, Cont. normals	Bilinear, Discontinuous

Both satisfy $\dim(\mathbb{V}_1) = 2\dim(\mathbb{V}_2)$, the condition necessary for preventing spurious modes.

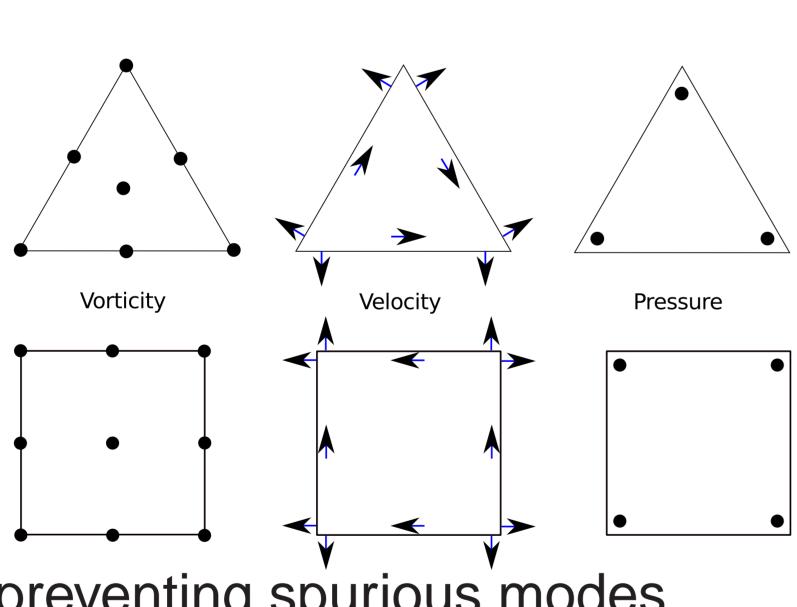
Constructing \mathbb{V}_1 on the sphere



All matrices (except mass matrices) are topological, i.e. they are independent of coordinates.

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ans that *u* must have



$$\frac{\mathbf{X}')}{\mathbf{X}')}.$$

Mixed finite element discretisation

Multiply equations 1-2 by appropriate test functions, $w \in V_1$ and $\phi \in V_2$, and integrate over the domain Ω :

$$\int_{\Omega} \boldsymbol{w} \cdot \boldsymbol{u}_t \, \mathrm{d}V + \int_{\Omega} \boldsymbol{w} \cdot \boldsymbol{Q}^{\perp} \, \mathrm{d}V - \int_{\Omega} \nabla \cdot \boldsymbol{w} \left(g \left(D + b \right) + \frac{1}{2} |\boldsymbol{u}|^2 \right) \, \mathrm{d}V = 0, \tag{3}$$

where we have introduced the mass flux F = uD and potential vorticity flux Q = qF. Require discrete potential vorticity $q \in \mathbb{V}_0$ to satisfy

 $\gamma q D dV = \int -\nabla^{\perp} \gamma u dV$

Differentiate equation 5 and substitute for u_t using equation 3 with $w = -\nabla^{\perp}\gamma$. Since $\nabla \cdot \nabla^{\perp} \equiv 0$ and we assume $f_t = 0$, this gives an advection equation for q:

$$\int_{\Omega} \gamma(qD)_t + \nabla \gamma \cdot \mathbf{Q} \, \mathrm{d} \, V = 0.$$
(6)

Rearranging gives:

$$\int_{\Omega} \gamma (Dq_t + F \cdot \nabla q) \, \mathrm{d}V = - \int_{\Omega} \gamma q \underbrace{(D_t + \nabla \cdot F)}_{=0 \text{ for flat elements [3]}} \, \mathrm{d}V.$$
(7)

i.e. if q is initially spatially constant it will remain so. This consistency can be recovered for higher order curved elements (see Colin's talk).

Timestepping

- discretise equations 3-4 in time using the theta method
- for updates to **u** and D (see following box)
- ▶ first need to calculate **F** and **Q**:
- ▶ solve equation 4 for *D* and calculate mass flux *F* (see flux reconstruction) solve equation 6 for q and diagnose potential vorticity flux Q (see Taylor-Galerkin methods)

Hybridized Helmholtz equation

$$\int_{\Omega} \boldsymbol{w} \cdot \Delta \boldsymbol{u} \, \mathrm{d}\boldsymbol{x} - \Delta t \int_{\Omega} \nabla \cdot \boldsymbol{w} \Delta D \, \mathrm{d}\boldsymbol{x} = \int_{\Omega} \boldsymbol{w} \cdot \boldsymbol{R}_{\boldsymbol{u}} \, \mathrm{d}\boldsymbol{x} \quad \forall \boldsymbol{w} \in \mathbb{V}_{1}$$

$$\int_{\Omega} \phi \Delta D \, \mathrm{d}\boldsymbol{x} + \Delta t \int_{\Omega} \phi \nabla \cdot \boldsymbol{u} \, \mathrm{d}\boldsymbol{x} = \int_{\Omega} \phi \boldsymbol{R}_{D} \, \mathrm{d}\boldsymbol{x} \quad \forall \phi \in \mathbb{V}_{2}$$

- Relax continuity constraints and allow u to be fully discontinuous
- Define a set of Lagrange multipliers λ on the trace space of \mathbb{V}_1 (i.e. on the set of element

$$\int_{\Gamma} \mu[\boldsymbol{u} \cdot \boldsymbol{n}] ds = 0, \int_{\Omega} \bar{\boldsymbol{w}} \cdot \Delta \bar{\boldsymbol{u}} dx - \Delta t \int_{\Omega} \nabla \cdot \bar{\boldsymbol{w}} \Delta D dx = \int_{\Omega} \bar{\boldsymbol{w}} \cdot \boldsymbol{R}_{\boldsymbol{u}} dx + \int_{\Gamma} \lambda[\boldsymbol{w} \cdot \boldsymbol{n}] ds \quad \forall \bar{\boldsymbol{w}} \in \bar{\mathbb{V}}_{1}$$

- Eliminate both \boldsymbol{u} and D to give a symmetric, positive definite matrix-vector equation for $\boldsymbol{\lambda}$

Flux reconstruction

Solve the weak form of the mass continuity equation 4 in each element: $\int \phi \Delta D \, dx - \Delta t \int \nabla \phi \cdot u D \, dx$

where \tilde{D} is the value of D on the upwind side of ∂e , using standard DG advection methods in this case, 3rd order SSPRK. Aim: to find mass flux $F \in V_1$ that satisfies $\Delta D = \Delta t \nabla \cdot F$. This mass flux can be constructed in each element by solving the following set of equations:

$$\int_{\delta e} \phi \boldsymbol{F} \cdot \boldsymbol{n} \, \mathrm{d}\boldsymbol{s} = \int_{\delta e} \phi \tilde{\boldsymbol{D}} \boldsymbol{u} \cdot \boldsymbol{n} \, \mathrm{d}\boldsymbol{S}, \quad \int_{e} \nabla \phi \cdot \boldsymbol{F} \, \mathrm{d}\boldsymbol{x} = \int_{e} \nabla \phi \cdot \boldsymbol{u} \boldsymbol{D} \, \mathrm{d}\boldsymbol{V} \quad \forall \phi \in \mathbb{V}_{2}, \quad \int_{e} \nabla^{\perp} \gamma \cdot \boldsymbol{F} \, \mathrm{d}\boldsymbol{V} = \boldsymbol{0} \quad \forall \gamma \in \mathbb{V}_{0}$$

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$$\int_{\Omega} \phi \left(D_t + \nabla \cdot \boldsymbol{F} \right) \mathrm{d} V = \mathbf{0}, \qquad (4)$$

$$\mathbf{I}\mathbf{V} + \int_{\Omega} \gamma f \, \mathrm{d}\mathbf{V} \quad \forall \gamma \in \mathbb{V}_0.$$
 (5)

within each timestep perform multiple quasi Newton iterations: solve the Helmholtz equation

Trace space for RT0 elements

the normal component of *u*.

• *u* and *D* are reconstructed within each element from the values of λ on the element edges. Advantages: can implicitly include Coriolis term and avoid lumping the mass matrix.

$$d\mathbf{x} + \Delta t \int_{\partial e} \phi D \mathbf{u} \cdot \mathbf{n} ds = 0.$$

Taylor-Galerkin methods for PV advection

Solve equation 6 using a multistage method with q at each stage defined as

$$\hat{qD}_i - \eta$$

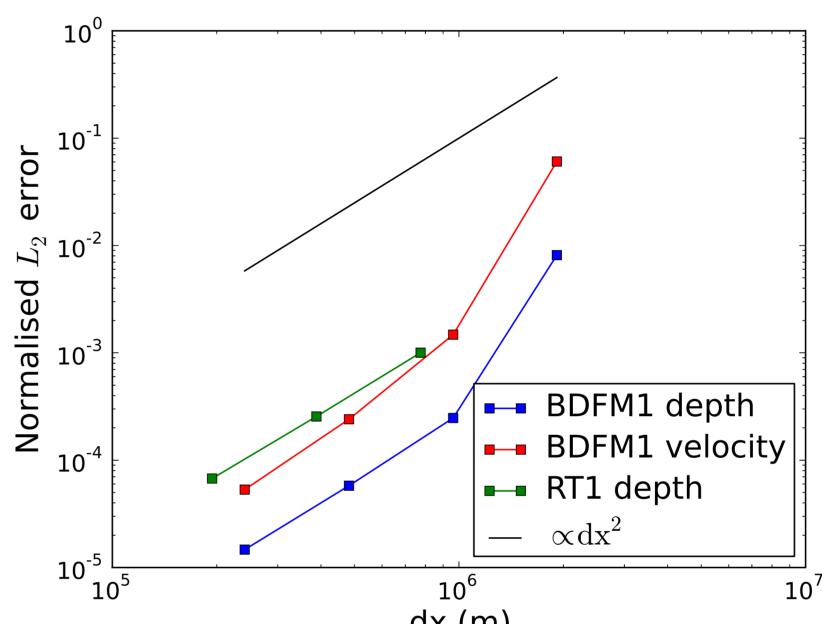
where i = 1, ..., k denotes the stage and the $\{\mu\}_{ij}$ and $\{\nu\}_{ij}$ are coefficients defined in [4]. Using equation 6 to replace temporal derivatives with spatial derivatives, we can write:

$$\int_{\Omega} \gamma \left((qD)_j \right)_t \mathrm{d}V = -\int_{\Omega} \nabla \gamma \cdot \mathbf{F} q \, \mathrm{d}V \quad \text{and} \quad \int_{\Omega} \gamma \left((qD)_j \right)_{tt} \mathrm{d}V = -\int_{\Omega} \frac{\mathbf{F}}{D} \cdot \nabla \gamma \mathbf{F} \cdot \nabla q \, \mathrm{d}V$$

which, comparing with equation 6, we see is exactly the form we require. We use a 2 level, 3rd order in time $\overline{T}(2,3)$ Taylor-Galerkin scheme to solve the continuity equation for q. This is stable for values of $\eta > 0.473$.

Results

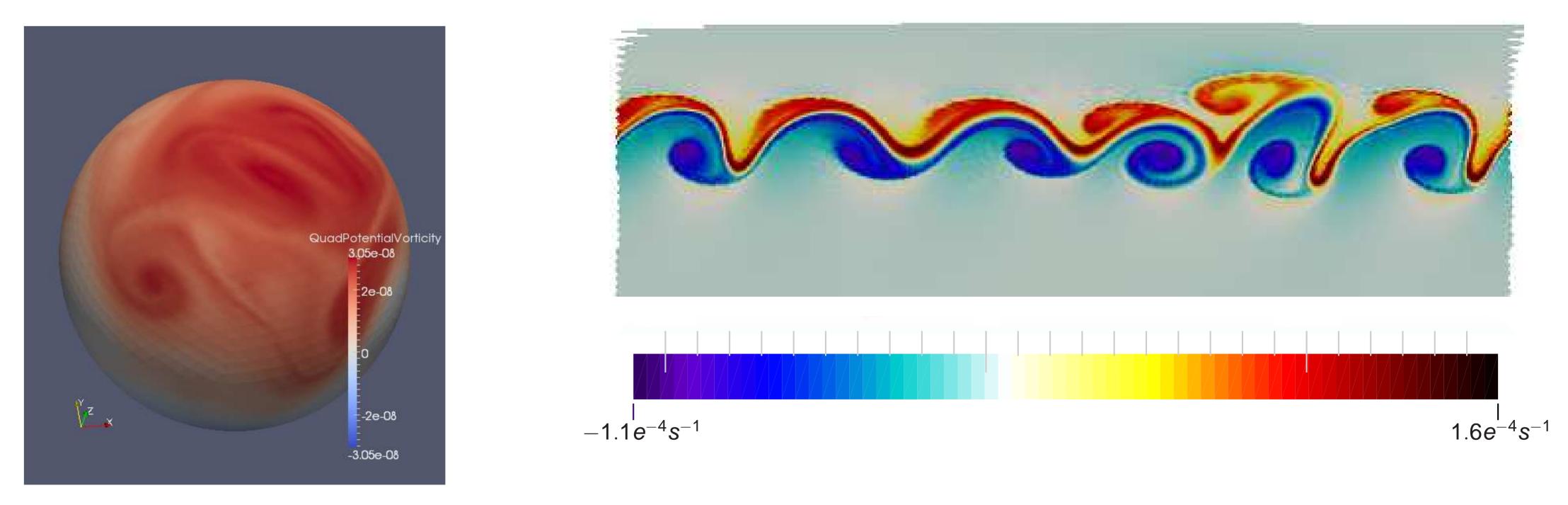
Solid body rotation: Williamson [5], test case 2



 L^2 error of depth and velocity fields after 5 days, versus mesh size dx, for BDFM1 and RT1 finite element spaces.

Flow over mountain:

Snapshot of potential vorticity at 50 days for Williamson test case 5, for the BDFM1 finite element space



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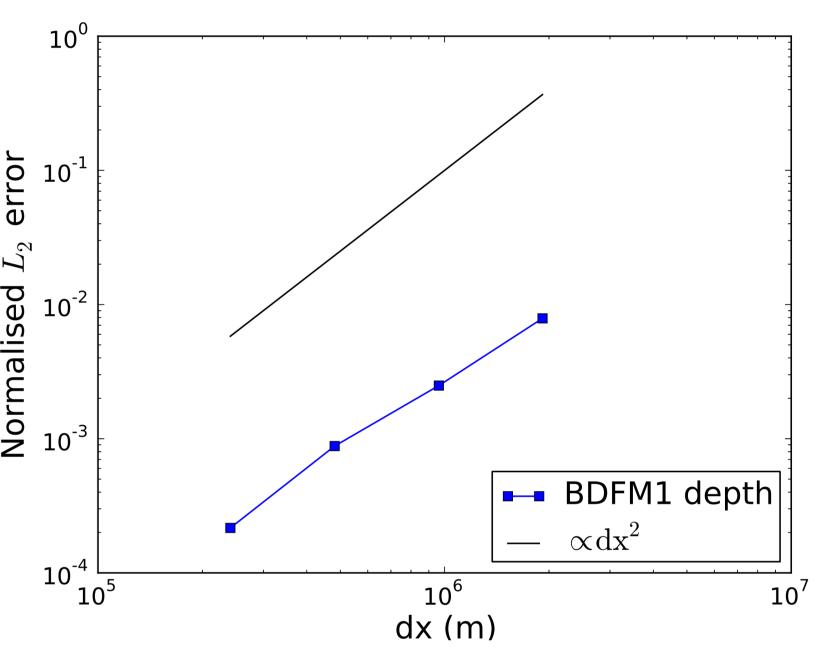
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$$(\Delta t)^{2}((qD)_{i})_{tt} = \Delta t \sum_{j=0}^{i-1} \mu_{ij}((qD)_{j})_{t} + (\Delta t)^{2} \sum_{j=0}^{i-1} \nu_{ij}((qD)_{j})_{tt}$$



Flow over mountain: Williamson [5], test case 5



L² error of depth field, compared to reference solution, at 15 days, versus mesh size dx, for the BDFM1 finite element space.

Barotropically unstable jet (Galewsky [2])

Vorticity field at 6 days for the barotropically unstable jet from Galewsky et. al. [2], using BDFM1 on a grid with 184320 DOFs, corresponding to an average mesh size of 240901m.

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