The Tropical Madden-Julian Oscillations As An Asymptotically-**Nondivergent Nonlinear Rossby Wave**

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What is MJO: Madden-Julian Oscillation?

- :Madden & Julian (1972)
- 30-60 day-period eastward-propagating atmospheric eauqtorial Wave:
- **Mechanism?**



MJO as a Vorticity Dynamics?

(e.g., Yanai et al., 2000)





(b)

400-hPa Temperature (K) and 850-hPa Streamlines



MJO:

- Simulated by a 1-km Deep Global Shallow-Water Model:
- Initialized by: Equivalent-Barotropic Modon for 250m-Deep QC System BATION AT T= 0.00 DAYS





Precipitation Field



-0.24 -0.16

-0.08 0.00

vertical wind (Pa/s

0.08

0.24

15 Apr 2009



70

-0.24

80

-0.16 -0.08

90

100 110 120 130

Longitude

vertical wind (Pa/s

160 170

140 150

0.00 0.08 0.16 0.24 0.32

Zonal Velocity: 7 Apr 2009



15 Apr 2009 b. 100 200 300 400 500 600 700 800 900 100 110 120 130 140 150 160 170 180 d. 12km paran 100 200 300 400 500 600 700 800 906 60 70 80 90 100 110 120 130 140 150 160 170 18 f. 12km 3Dsmag 100 200 300 400









zonal wind (m/s)

How the Global Models Simutate It? (e.g., Monier et al., 2010, MM5):



How the Global Models Simutate It?:

Difficulties in

- simulating <u>Convection (& precipitation)</u>, but
- the <u>Large-Scale Circulations</u> (e.g., zonal winds) are often <u>correct</u>.

i.e.,

The <u>Dynamical Component</u> of MJO can be well simulated <u>without Convection</u>

Dry MJO?

[convection is likely working as damping in simulations]

(a) **Asymptotic-**Nondivergence of the Large-**Scale Tropical** (b) Circulations **Scatter Plots** between **Vorticity and** (C) **Divergence** (TOGA-COARE LSA: Yano et al., 2009)



Asymptotic- Nondivergence

or Quasi-Nondivergence

 Leading-Order Dynamics is purely Rotational (<u>Vortex</u>-Dominant)

No Leading-Order Role of Convection

Dry Leading-Order Dynamics

Asymptotic-Nondivergence of MJO ?

Asymptotic-**Nondivergence** of MJO



q [g kg⁻¹]

Divergent Wind



Divergence/Vorticity (Transient)



Time scale (days)

(Zagar and Franzke 2015)

(Yano et al., 2009)

Asymptotic-Nondivergence of MJO : **Dry Wave Dynamics?: Two Slowest Equatorial Waves:**

•Kelvin Wave: Mondispersive (Not Slow Enough)

•Rossby Wave: ?

Linear Rossby-Wave Dispersion:

 $c_p = -\beta L_p^2$, $1/L_p^2 = k^2 + l^2 + (1/L_R)^2$

for MJO: $c_p = 5 m/s$ for $|\phi| < 30^\circ$, $\beta \sim 2 \times 10^{-11} 1/s/m$

 $iL_p = (c_p/L_p)^{1/2} \sim 500 \text{ km} \sim L_R \sim 10^3 \text{ km}$

- i.e.,
- only the <u>evanescent</u> wave solution can explain MJO
- under linear Rossy-wave dynamics

Just Laterally-Forced? (Vitart and Jung 2010) U850 Relaxed Relaxed to Obs. Contro to analysis Initial Conditions (a) (b) (c) (d) 2/0 3/0 ISE ISW LONGTUDE SE 15W 150E RE 1800 LONGTUDE 406 1505 RE 18W 19Œ 90W ¥4.,

IFS GCM IFS with $\phi > 35^{\circ}N$ **nudging**

<u>Just Laterally-Forced?</u> : Linear Thoery :

with zonal wavenumber : $n_x = 1 - 3$ $k^1 \sim 10^4 \text{ km} >> L_R \sim |L_p|$

latitudinal scale: $I^2 \sim 1/L_p^2 - [k^2 + 1/L_R^2]$ $\sim 1/L_p^2 - 1/L_R^2$ $\sim - 1/L_R^2$ (*il*)⁻¹~ $L_R \sim 10^3$ km : Influence Scale of the lateral forcing

i.e., the effect of the lateral forcing is confined to $L_R \sim 10^3$ km with no Effect seen at the Equator

→ <u>Nonlinear Response</u> (Nontrivial Problem)

Nonlinear Response (Nontrivial Problem):

Nonlinear Solitary Rossby-Wave: Theory for MJO:

- with prescribed latitudinal scale $I < 1/L_R$:
- Why It Propagates Easwards?:
- Because the tail-part is linear and evanescent with the scale : $(ik)^{-1} \sim L_R \sim 10^3 \text{ km}$



Nonlinear Solitary Rossby-Wave Simulation: MJO:



(Wedi and Smolarkiewicz 2010)

Main Conclusion: Preliminary Result:

GEOPOTENTIAL PERTURBATION AT T= Ø.ØØ DAYS



MJO:

Simulated by a 1-km Deep Global Shallow-Water Model:

×/Pi

Initialized by: Equivalent-Barotropic Modon for 250m-Deep QG System

MJO Problem: Analogy with Midlatitude Synoptic Storm Problem		
	Midlatitude Synoptic Storm	MJO
Old Dominant Theory	Controlled by <u>Rain Formation</u> (Sir John Mason, F. H. Ludlam)	Driven by <u>Convection</u> (Y. Hayashi, R. S. Lindzen,)
Modern Theory	Baroclinic Instabilities (Eady, Charney)	Nonlinear Solitary Rossby Wave (Yano, Wedi, Smolarkiewicz)

Further Materials for Questions and Discussions

Linear Free Wave Solutions: RMS of divergence/vorticity



Forced Problem

$\frac{\partial^2 \tilde{v}}{\partial y^2} + \left[\left(\frac{\nu^2}{gh_e} - k^2 - \frac{k}{\nu} \beta \right) - \frac{\beta^2 y^2}{gh_e} \right] \tilde{v} = H_n(\xi) \exp(-\xi^2/2 - i\nu t)$

Linear Forced Wave Solutions(c_g=50m/s): RMS of divergence/vorticity n=0 n=1

(a) $c_g = 50 m/s$, n = 0

(b) $c_s = 50 \text{m/s}, n = 1$



Balanced?

(Free-Ride, Fraedrich & McBride 1989):

Vertical Advection =Diabatic Heating



Vorticity >> Divergence with MJO:



GRADS: CO_A/IDES

Grabs: COLA/IDES



PC2



-0.24 -0.16

-0.08 0.00

vertical wind (Pa/s)

0.08 0.16 0.24 0.32







-0.32 -0.24 -0.16 -0.08 0.09 0.08 0.16 0.24 0.32 vertical wind (Pa/s)

Zonal Velocity: 7 Apr 2009







40 50 60 70 80 90 100 110 120 130 140 150 160 170 180 4km 3Dsmao

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60 90 100 110 120 130 140 150 160 170 180 Longitude



10 Apr 2009 b 100 200 300 400 500 600 700 800 900 1000 40 100 110 120 130 140 150 160 170 180 50 60 70 - An 901 d. 12km paran 100 200 300 400



500 600

700

800

900

1000

100

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600

700

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200

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400

500

600

700

800

900

1000

1.





50 60 70 80 90 100 110 120 130 140 150 160 170 180 4km 3Dsmao



70 80 90 100 110 120 130 140 150 160 170 180 50 60 ECMWF



Longitude



15 Apr 2009







f.



70 80 100 110 120 130 140 150 160 170 60 - 90 40 -50 h. 4km 2Dsmag



60 70 80 90 100 110 120 130 140 150 160 170 180 40 50 4km 3Dsmag



40 50 -AD 90 L ECMWF







(Adames et al 2014)





(Adames et al 2014)

Asymptotic Tendency for Non-Divergence: Divergence/Vorticity(Transient)



Weakly-Nonlinear Theory Does Not Explain MJO:

Rossby Wave Dispersion: $\omega = -\beta k/(k^2 + l^2 + F)$

Weakly-Nonlinear Approximation: Longwave Limit:

$$\omega \sim -\beta k/(l^2 + F)[1 - k^2/(l^2 + F)]$$

or

or

$$c_{\rho} \sim \beta k/(l^2 + F)[1 - k^2/(l^2 + F)] < 0$$
 always!

 $-d/dt = -\beta/(l^2 + F)(d/dx)[1 + 1/(l^2 + F)(d^2/dx^2)]$

cf., KdV Equation: $u_t + uu_x + u_{xxx} = 0$

Equatorial-Rossy Solition Dispersion (Boyd 1980)

