



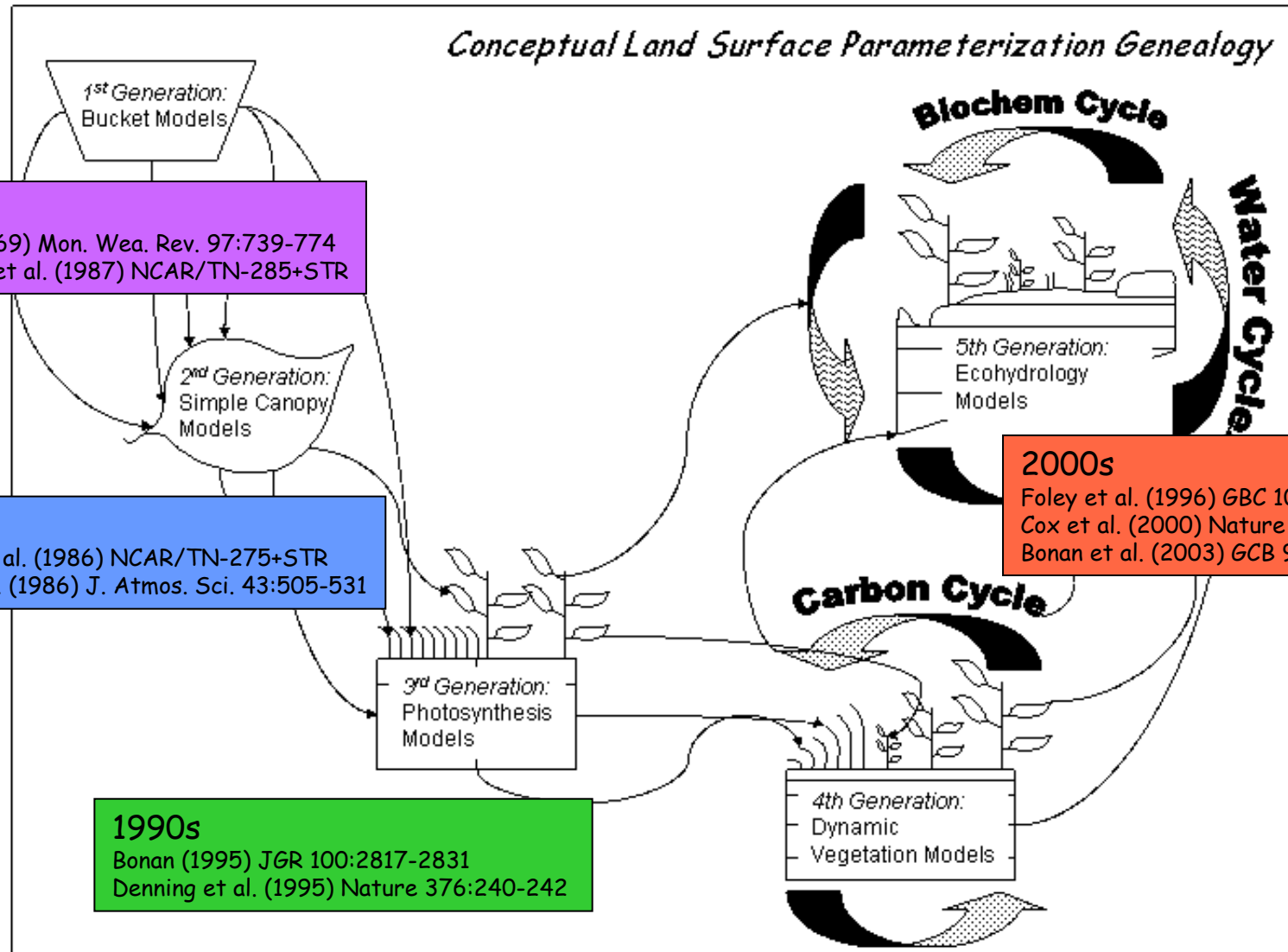
"Advances in land-climate interactions for earth system models: The Community Land Model (CLM) experience"

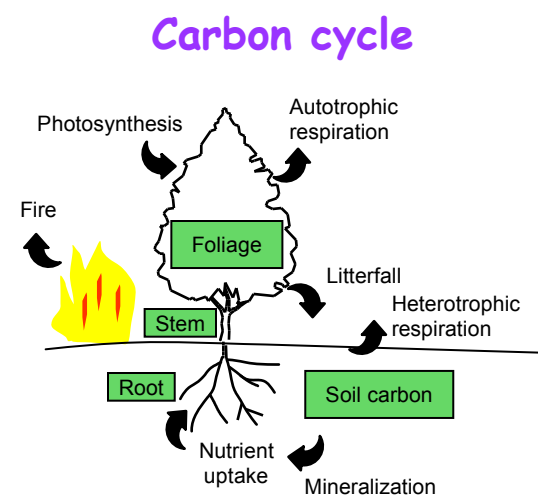
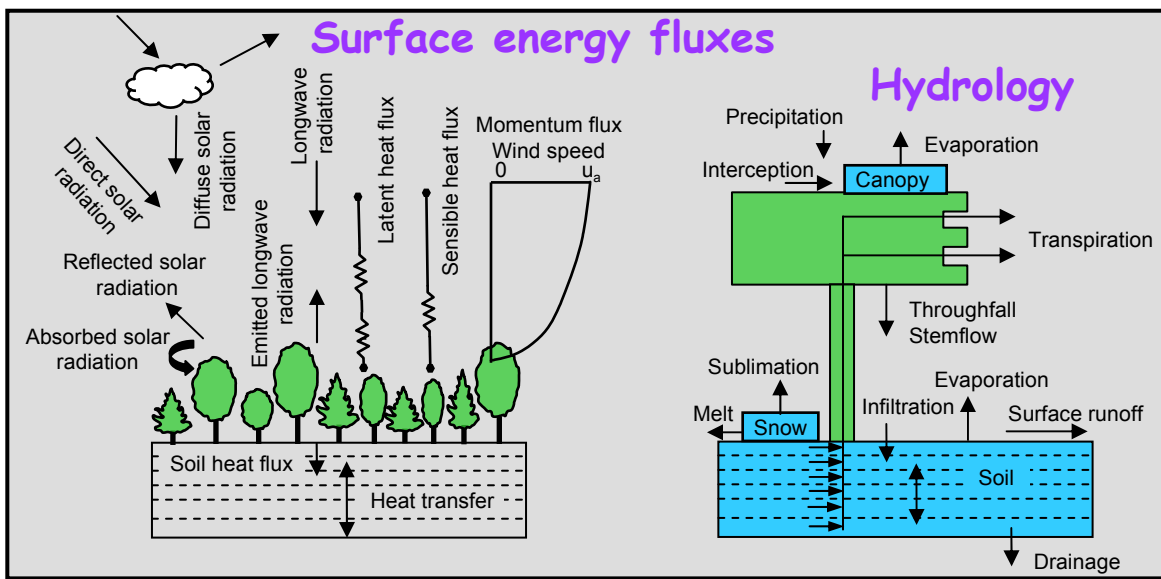
Gordon Bonan
National Center for Atmospheric Research
Boulder, Colorado

With contributions from
Dave Lawrence (NCAR)
Dave Gochis (NCAR)

European Geosciences Union
General Assembly 2008
Vienna, Austria
14 April 2008

Evolution of land surface models





Current-generation models

Disturbance

Establishment

Urbanization



Vegetation dynamics



Deforestation



Land use



Reforestation

Competition

Growth

Land as set of interacting systems

- Biogeochemical systems

- nutrient cycling, trace gas emissions, reactive chemistry, constituent tracing, isotopes

- Water systems

- water resources management, freshwater availability and water quality

- Human systems

- land use, urbanization, energy use

- Ecosystems

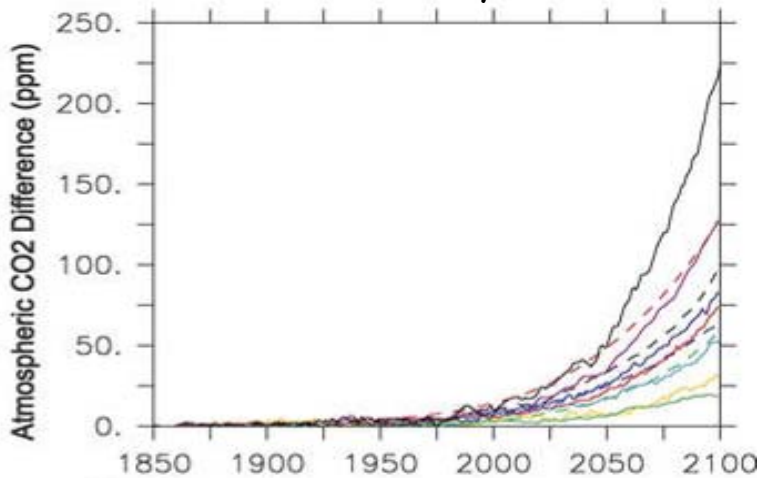
- resilience, good and services

- Ice sheet and glacier systems

- slow and fast ice sheet dynamics

The future of (land) climate change science

Climate-carbon cycle feedback

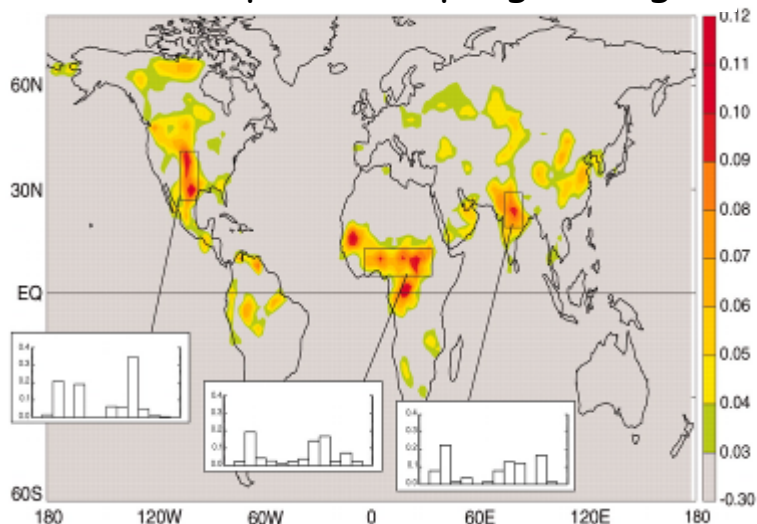


Friedlingstein et al. (2006) *J. Climate* 19:3337-3353

These models are being used to explore forcings and feedbacks in the earth system

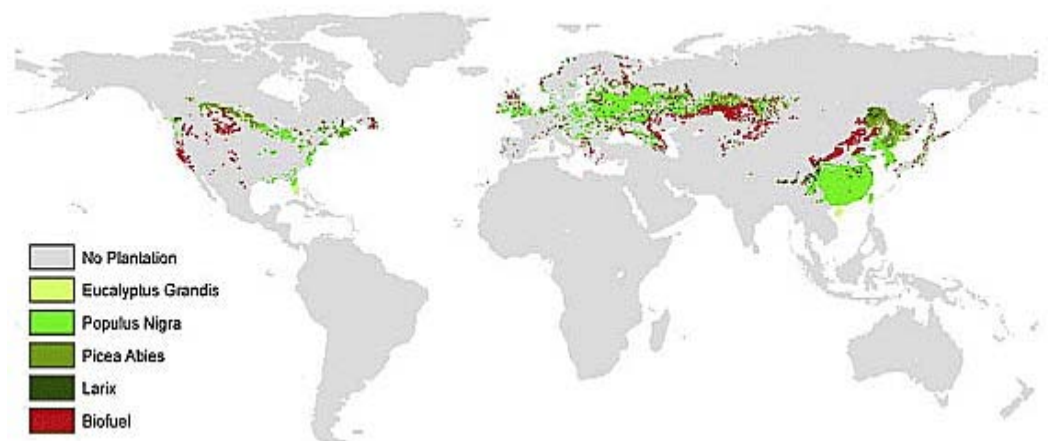
Increasing demand for information on impacts (especially regional), adaptation strategies, and mitigation solutions

Land-atmosphere coupling strength



Koster et al. (2004) *Science* 305:1138-1140

Land management mitigation strategies



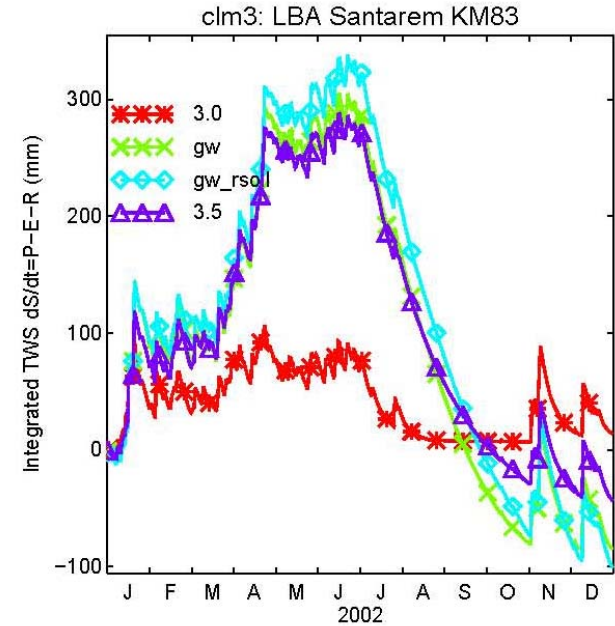
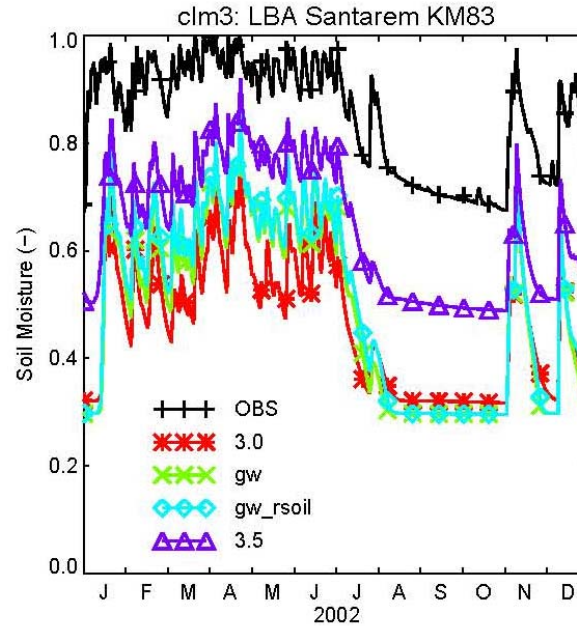
Schaeffer et al. (2006) *GBC*, doi:10.1029/2005GB002581

Hydrometeorology

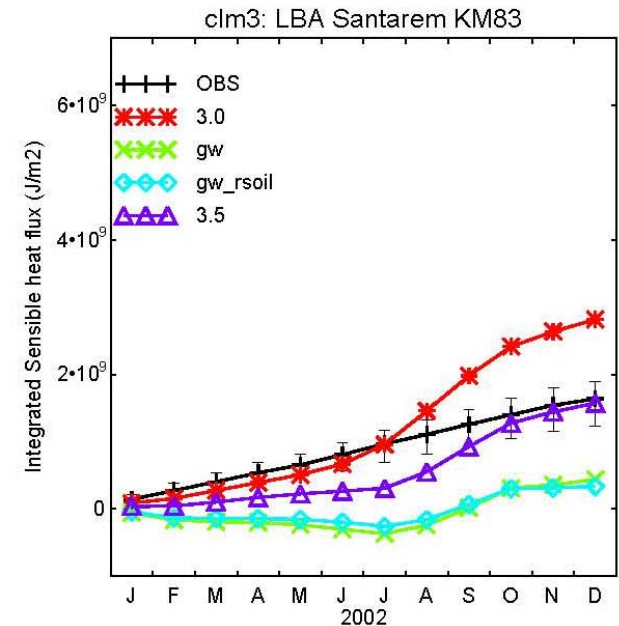
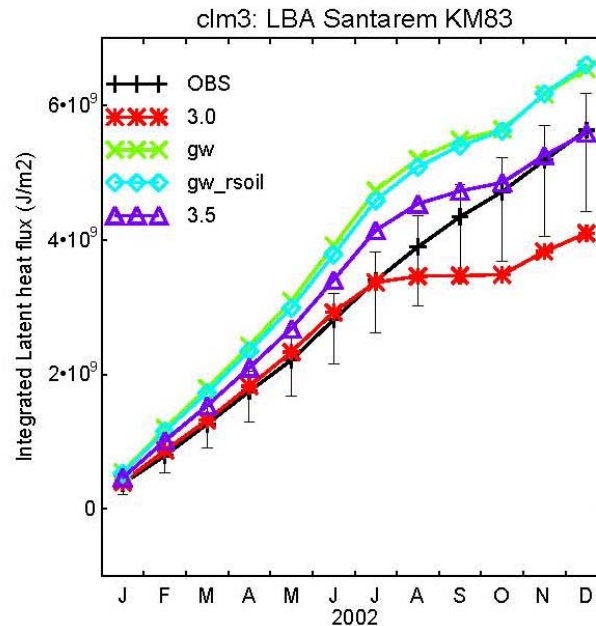
Flux tower observations,
tropical evergreen
forest, Brazil

CLM3 - dry soil, low dry
season latent heat flux,
high dry season sensible
heat flux

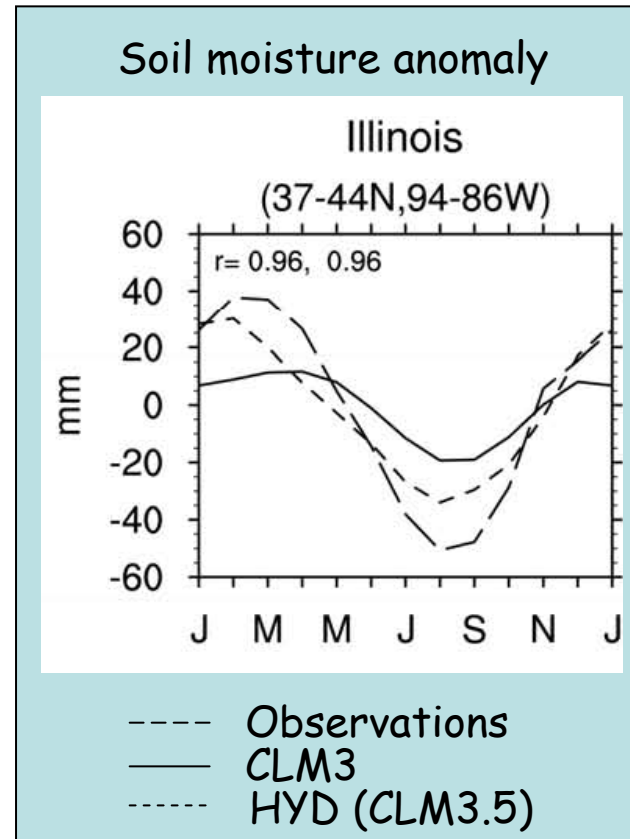
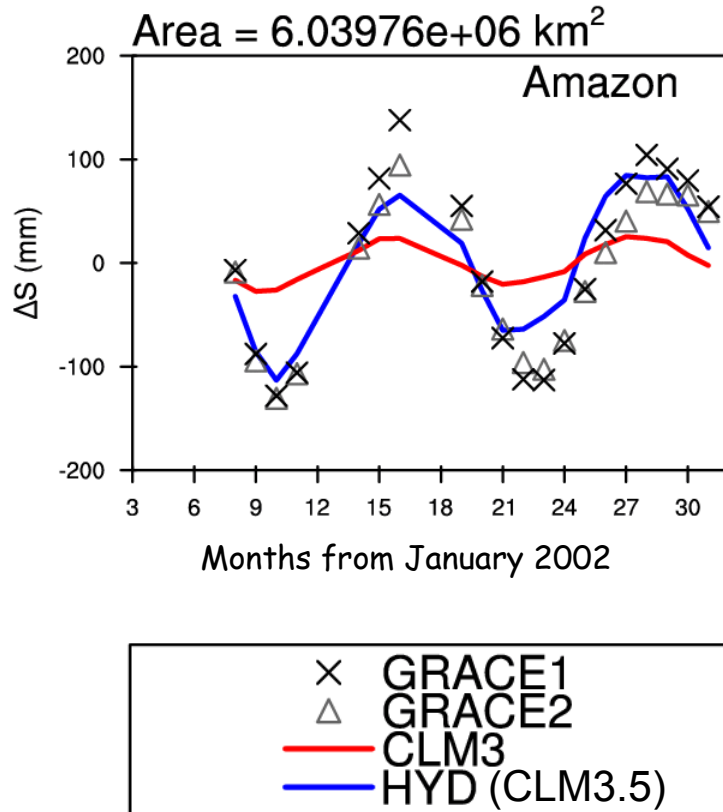
CLM3.5 - wetter soil
and higher latent heat
flux during dry season



b)



Hydrologic cycle



Observations - Large annual cycle in soil moisture storage

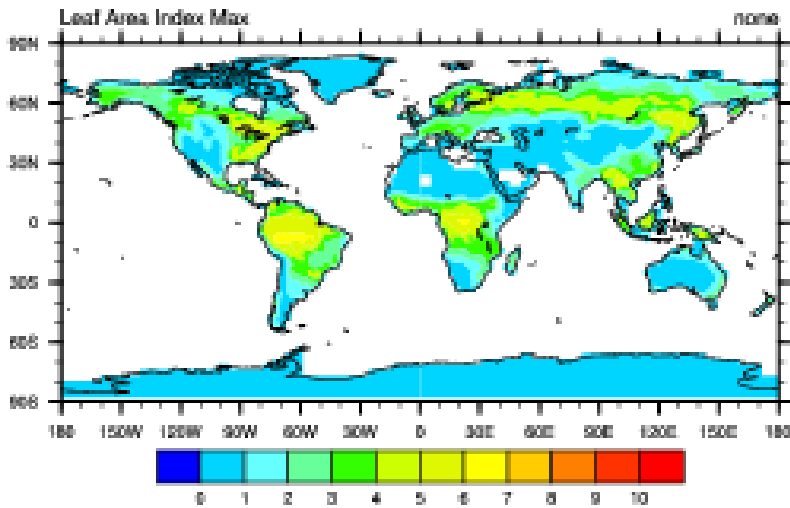
CLM3 - Small change in soil moisture storage

CLM3.5 - Better reproduces observations

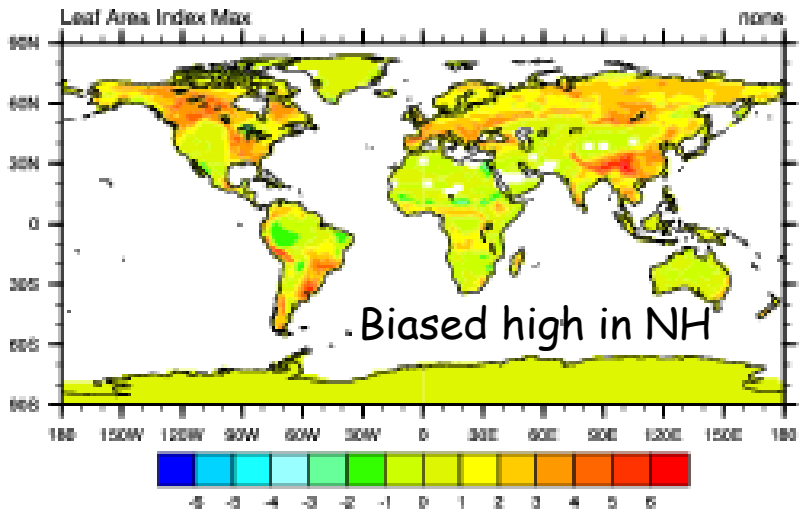
Improved leaf area index bias

Better hydrologic cycle improves carbon cycle

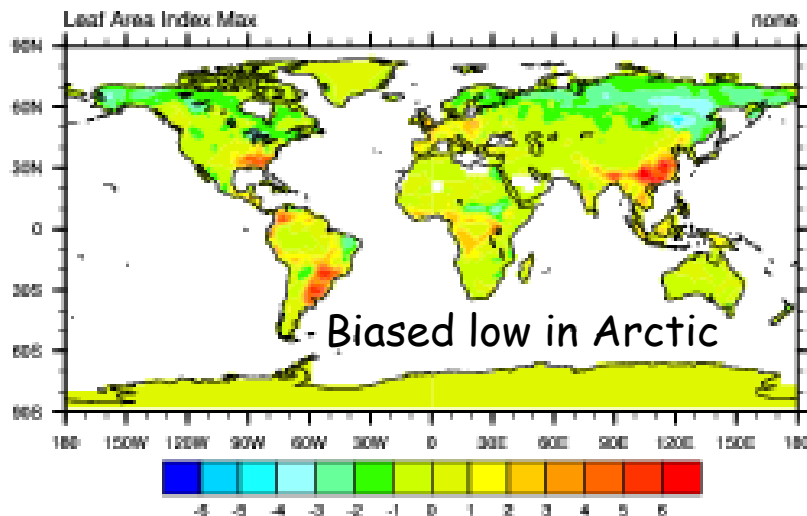
MODIS MOD 15A2 2000-2005



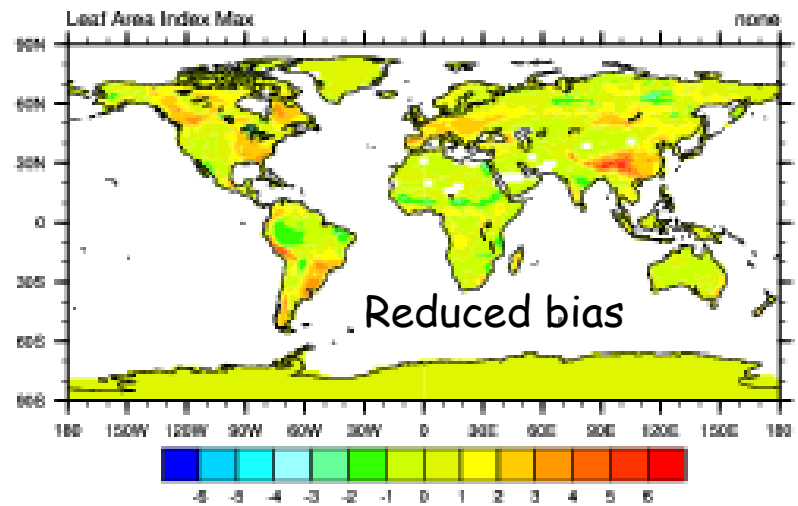
CLM3.5 - MODIS



CLM3.1 - MODIS

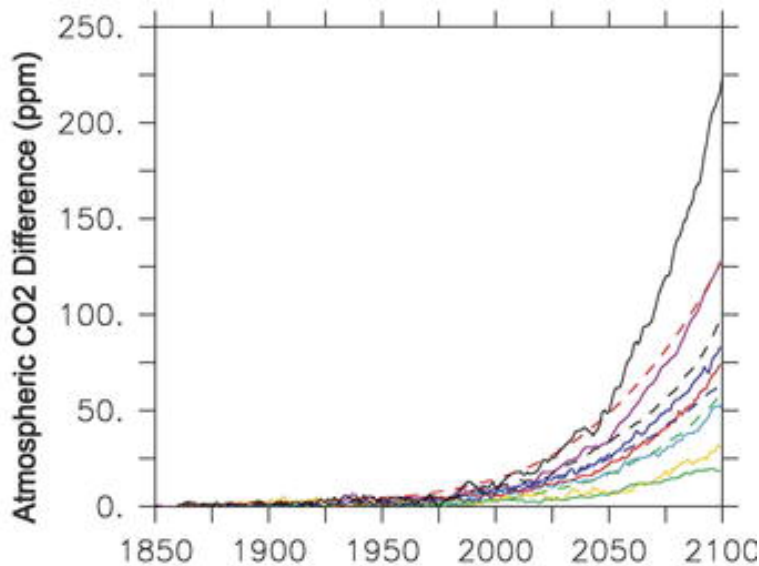


CLM3.6 - MODIS



C4MIP - Climate and carbon cycle

Effect of climate change on carbon cycle

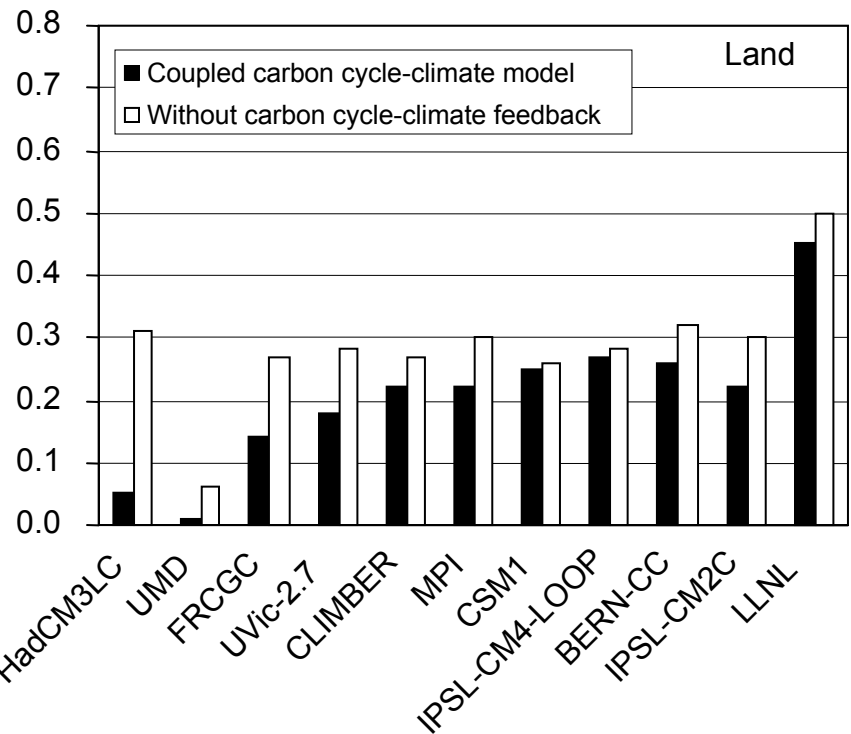


Climate-carbon cycle feedback

All models have a positive climate-carbon cycle feedback. The difference between fully coupled carbon cycle climate simulations and uncoupled simulations (CO_2 has no radiative effect) ranges from 20 ppm to >200 ppm

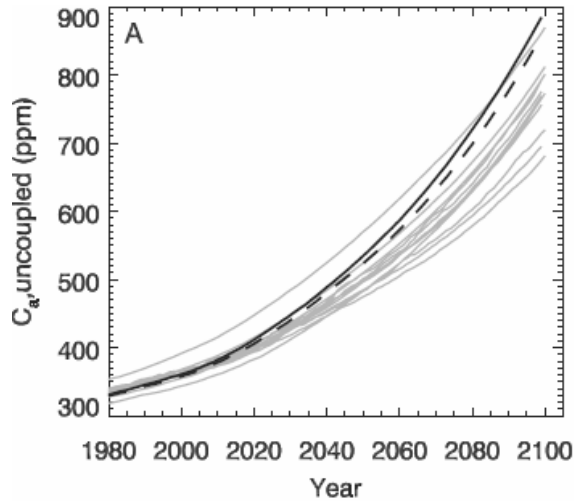
Distribution at 2100 of cumulative anthropogenic carbon emissions

The amount of carbon stored in the atmosphere increases in each model compared with the comparable simulation without climate-carbon cycle feedback, while the land carbon storage decreases.

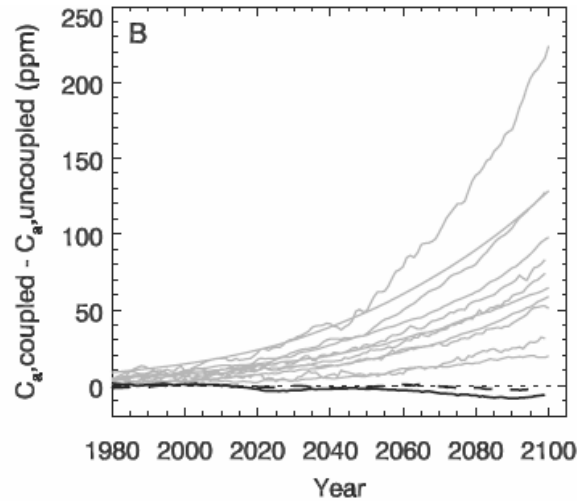


Climate, carbon, and nitrogen cycle

Ca from uncoupled experiments

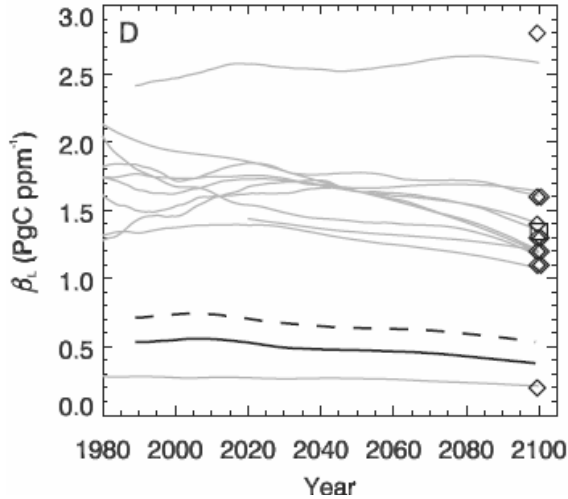


ΔCa due to radiative coupling

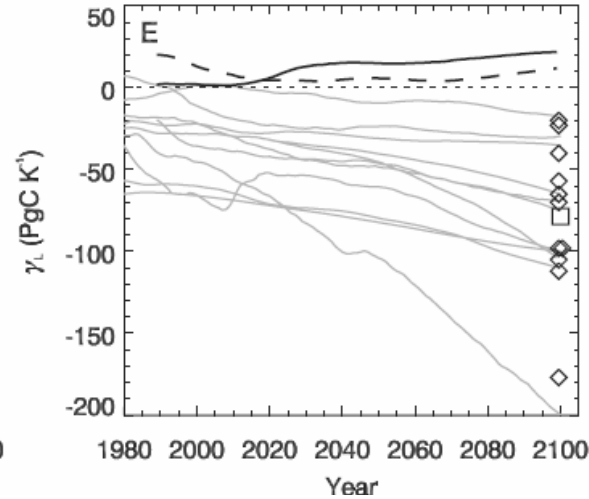


Panel A: Atmospheric CO_2 (C_a) of 884 ppm by 2100, radiatively-uncoupled

Panel B: Radiative coupling reduces C_a by 6 ppm, with a further reduction of 27 ppm due to anthropogenic N deposition



Land biosphere response to increasing atmospheric CO_2



Land biosphere response to increasing temperature

Gray lines show archived results from eleven previous studies

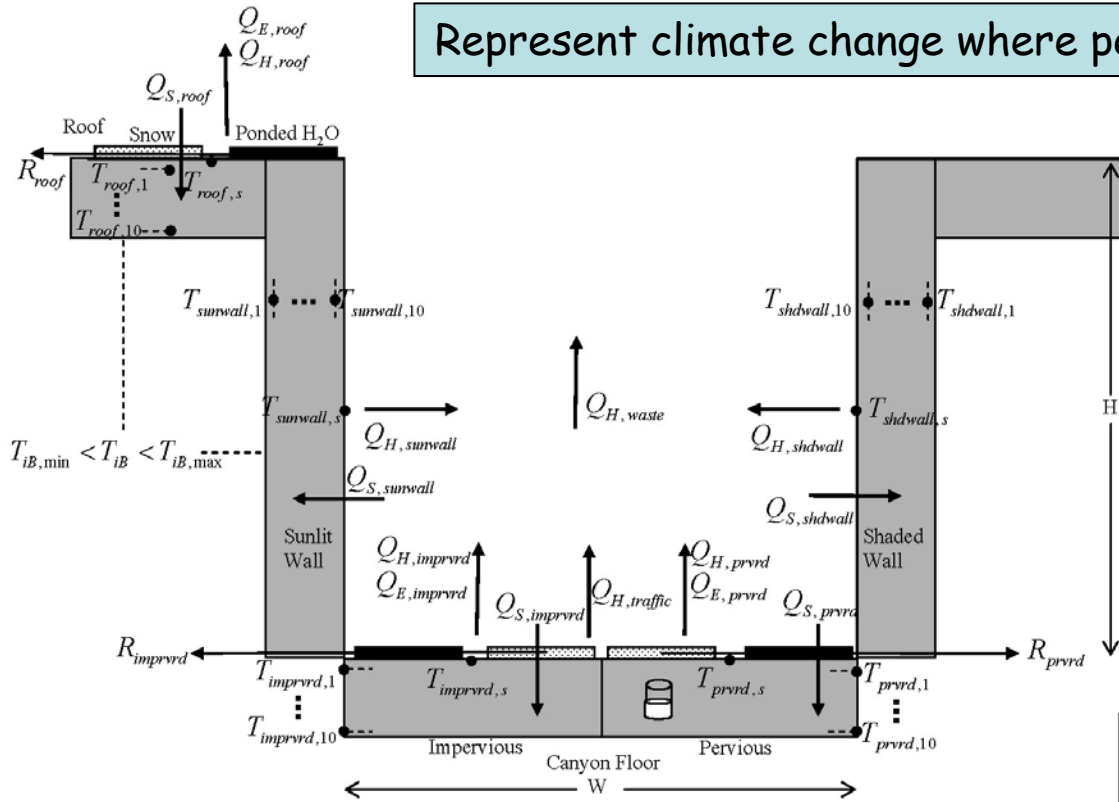
Thick solid line is for CLM with preindustrial N deposition

Thick dashed line is for CLM with anthropogenic N deposition

— low Ndep
- - - high Ndep

Urban systems

Represent climate change where people live

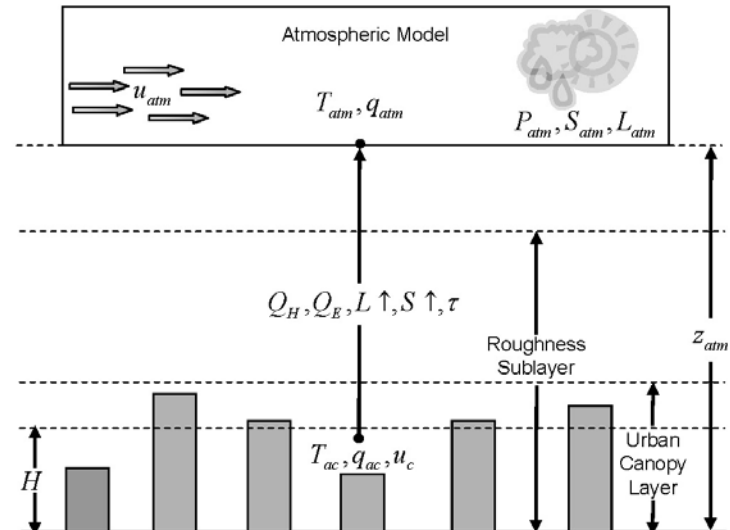


Urban canyon

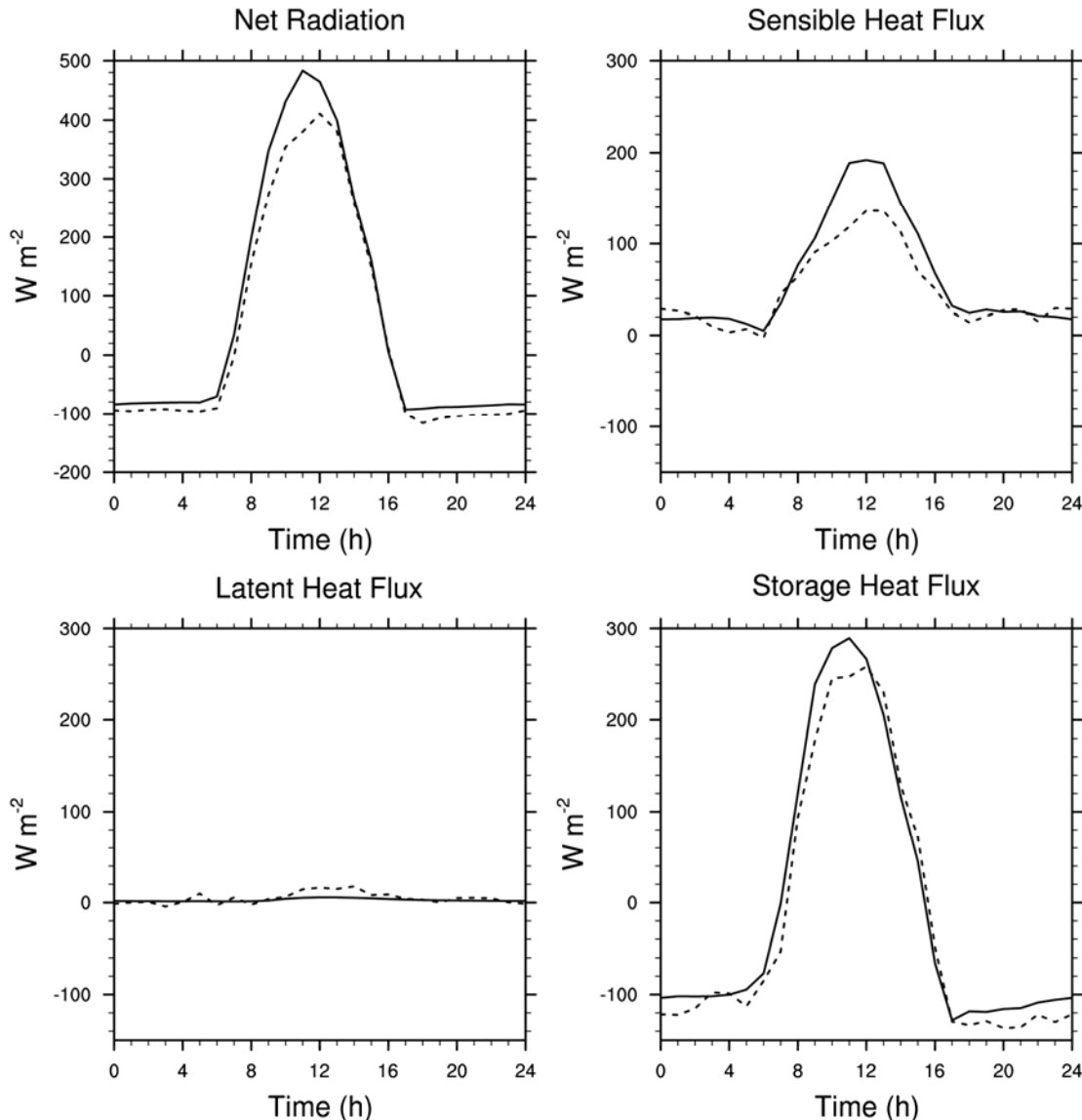


Energy fluxes are modeled for an idealized urban canyon. Key model parameters are:

- H/W - ratio of building height and street width
- W_{roof} - fractional roof area
- $f_{pervious}$ - fractional greenspace
- Building materials - thermal and radiative properties



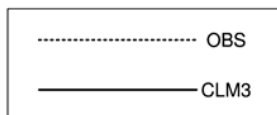
Simulated urban energy balance



Average diurnal cycle of simulated and observed heat fluxes for the Mexico City site (Me93) for 2-7 Dec 1993

Key features

- Diurnal cycle is well represented
- Simulated net radiation is too high (model ignores pollution), which drives high sensible heat
- Negligible latent heat flux
- Large storage heat flux



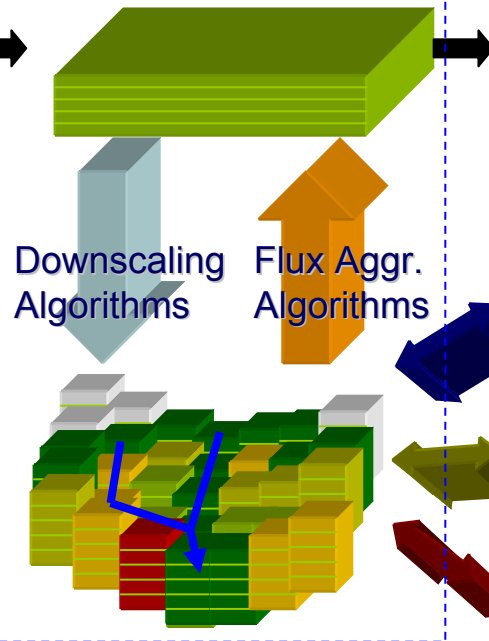
Fine-mesh Community Land Model

2.5° in longitude × 1.9° in latitude (~200 km)



Community Land Model

0.5° resolution



Scale-relevant information for impact studies

Effects of topography on hydroclimate

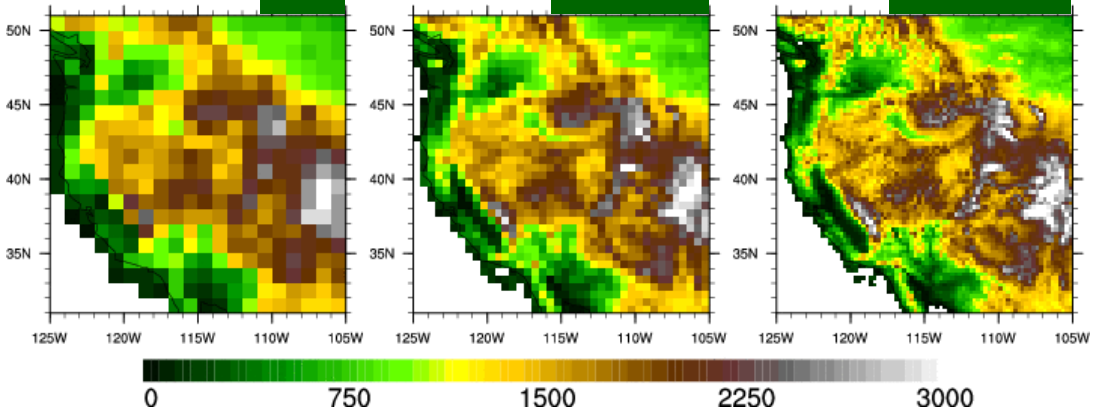
Precipitation downscaling



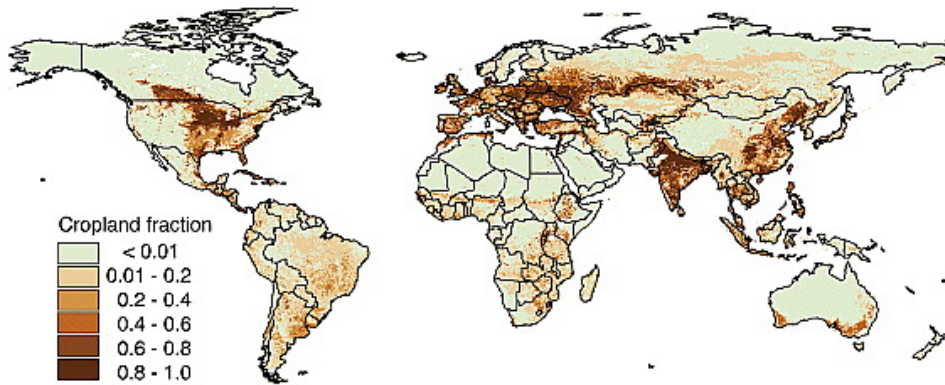
1°

1/2°

1/4°



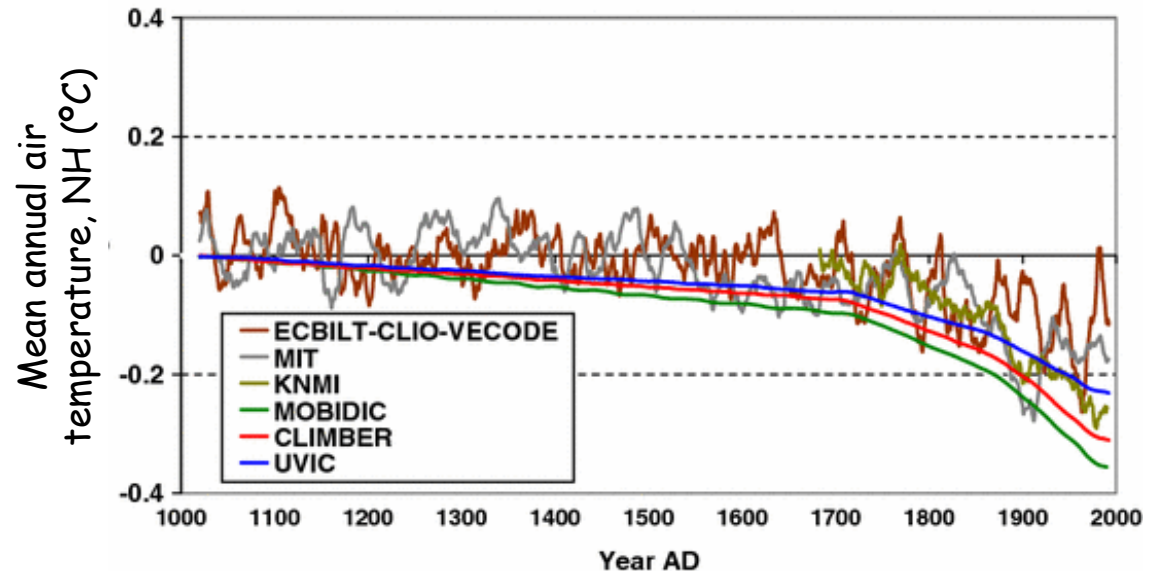
Historical land use forcing of climate



The emerging consensus is that land cover change in middle latitudes has cooled the Northern Hemisphere (primarily because of higher surface albedo in spring)

Northern Hemisphere annual mean temperature decreases by 0.19 to 0.36 °C relative to the pre-industrial era

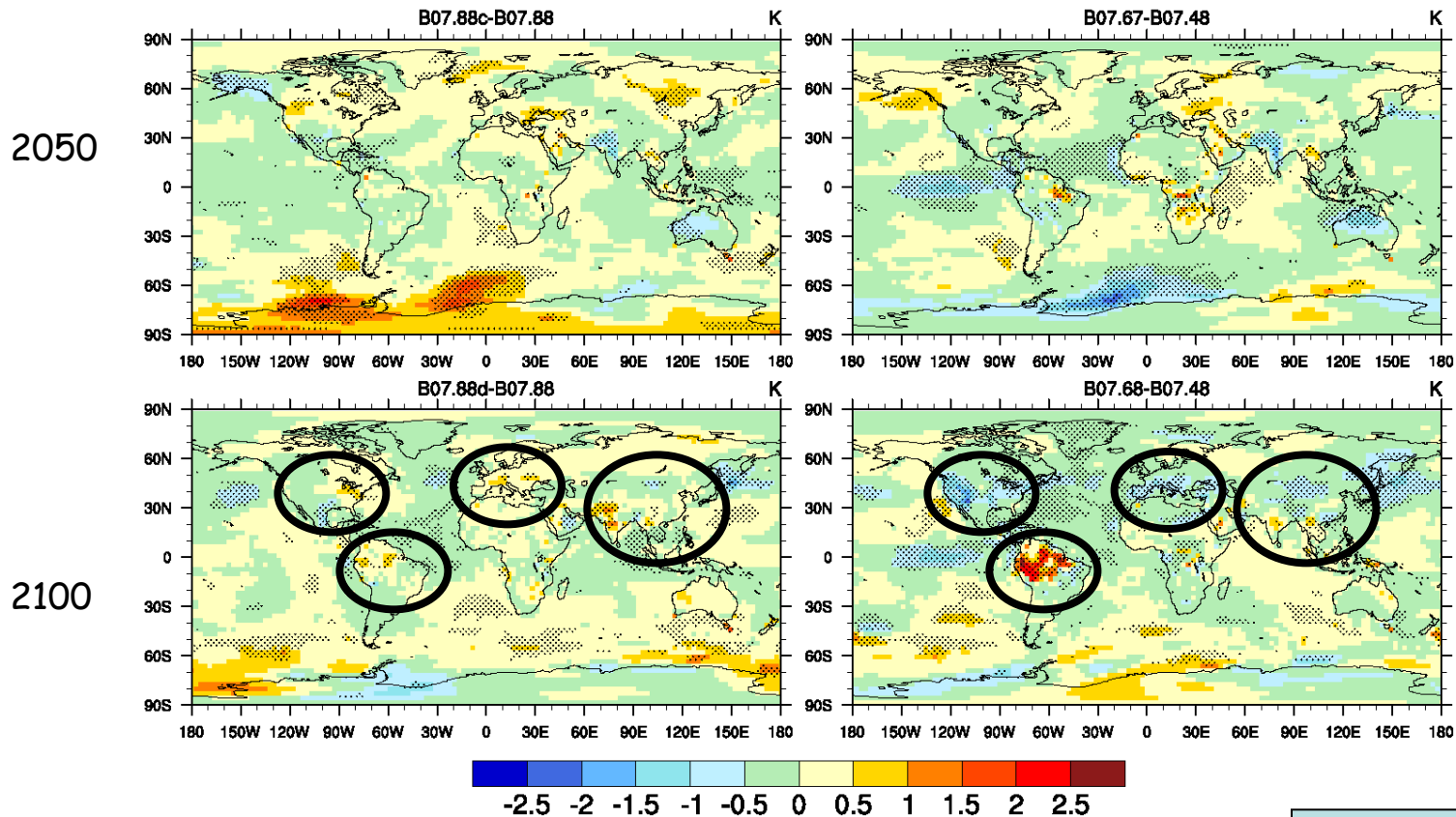
Comparison of 6 EMICs forced with historical land cover change, 1000-1992



Future land cover change as a climate forcing

(SRES land cover + SRES atmospheric forcing) - SRES atmospheric forcing

SRES B1 JJA reference height temperature SRES A2



A2 - Widespread agricultural expansion with most land suitable for agriculture used for farming by 2100 to support a large global population

B1 - Loss of farmland and net reforestation due to declining global population and farm abandonment in the latter part of the century

Land use choices affect climate

Integrated biogeophysical and carbon effects

Land cover change impact on 2100 climate

Biogeophysical

A2 - cooling with widespread cropland

B1 - warming with temperate reforestation

Biogeochemical

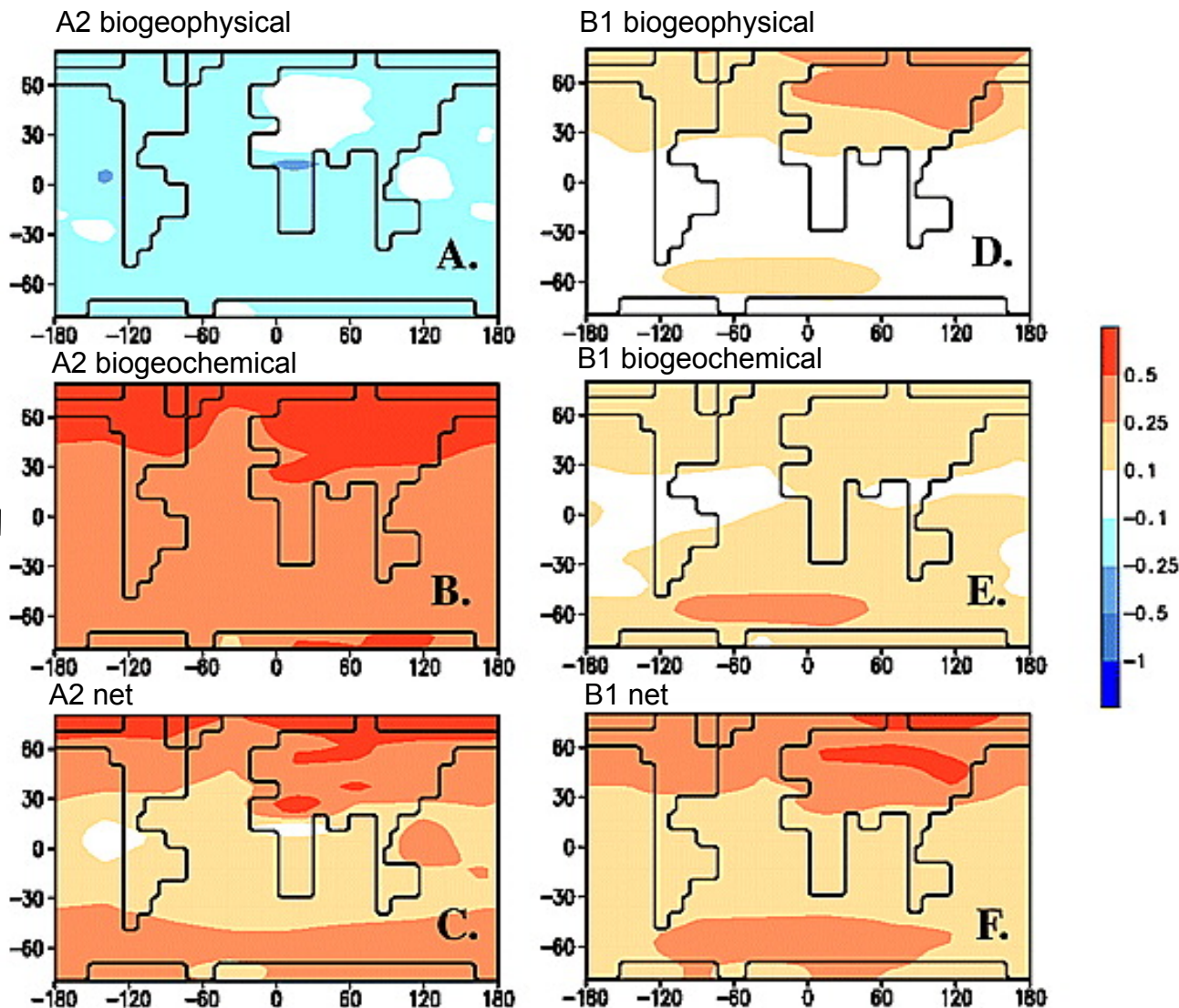
A2 - large warming; widespread deforestation

B1 - weak warming; less tropical deforestation, temperate reforestation

Net effect similar

A2 - BGC warming offsets BGP cooling

B1 - moderate BGC warming augments weak BGP warming



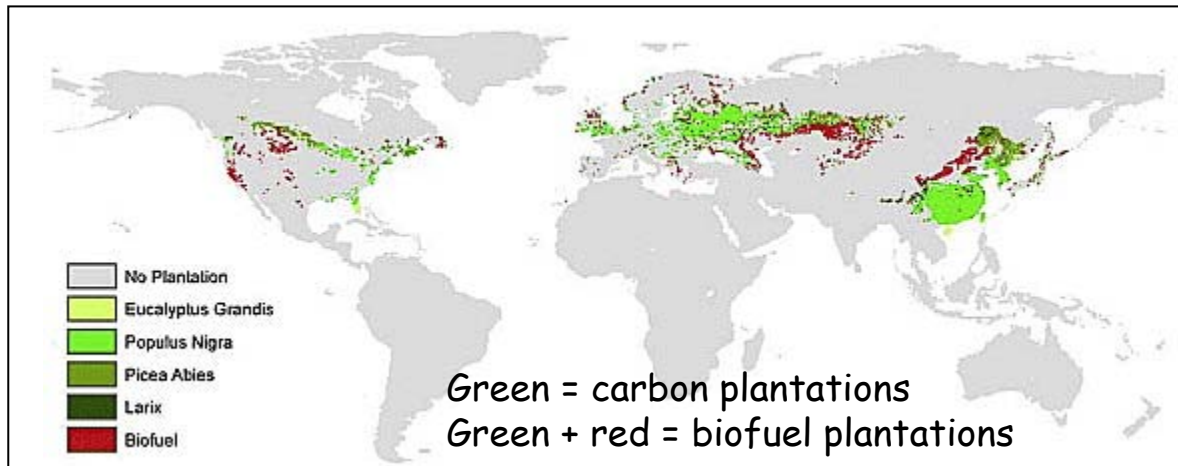
Land management policies to mitigate climate change

Reforestation might be chosen as an option for the enhancement of terrestrial carbon sequestration or biofuel plantations may be used as a substitute for fossil fuels



Ethanol from maize Hybrid poplar plantation

2100 land management, IPCC A1B scenario

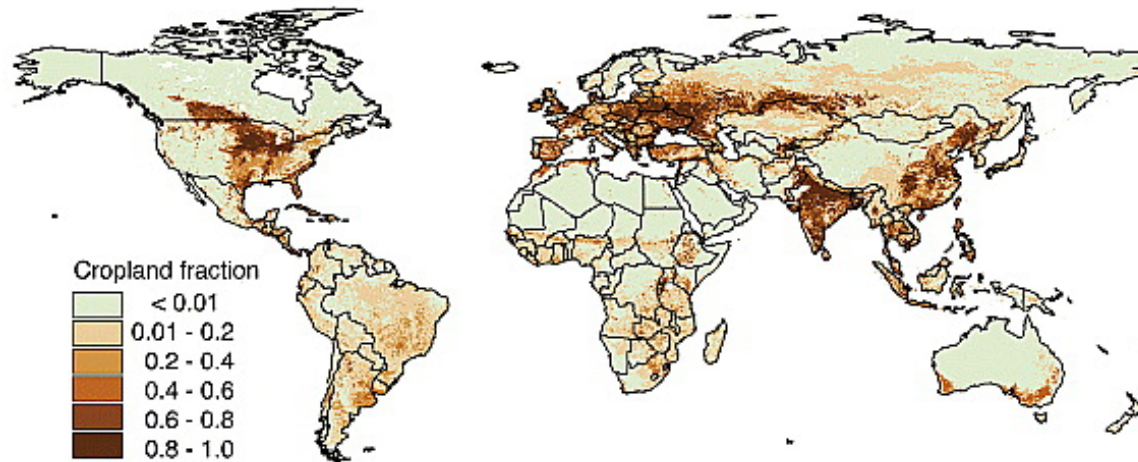


Carbon plantations and biofuel plantations reduce atmospheric CO_2 , leading to cooling

Reduced surface albedo leads to warming, but carbon plantations have lower albedo than biofuels

Excess agricultural land converted to carbon storage or biofuels

Multi-model ensemble of global land use climate forcing



Models

Atmosphere - CAM3 and CAM3.5 (differ in convection, PBL, and clouds)

Land - CLM3.5

Ocean - Prescribed SSTs and sea ice

Experiments

30-year simulations (using SSTs for the period 1972-2001)

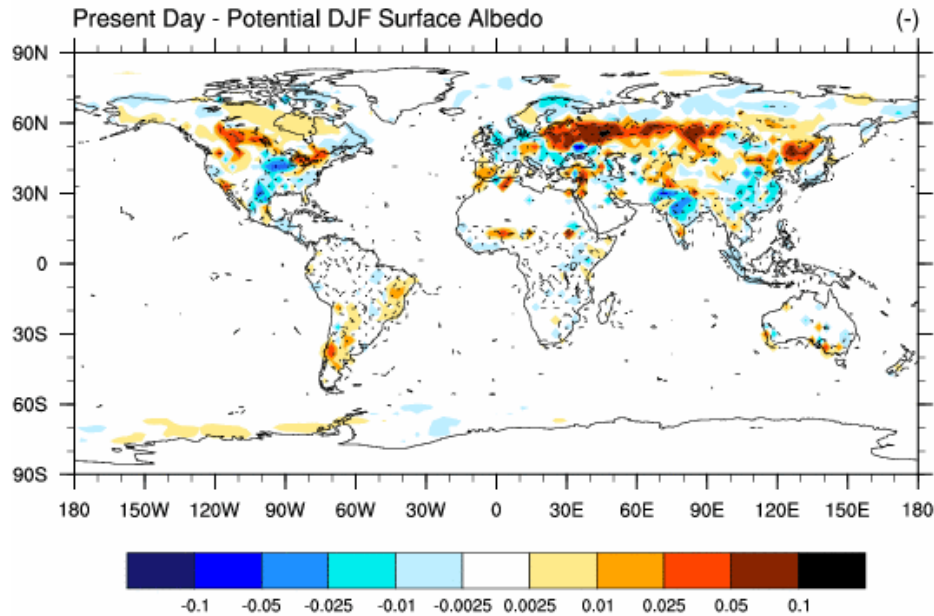
Two land covers: present-day vegetation and potential vegetation

5-member ensembles

Total of 20 simulations and 600 model years

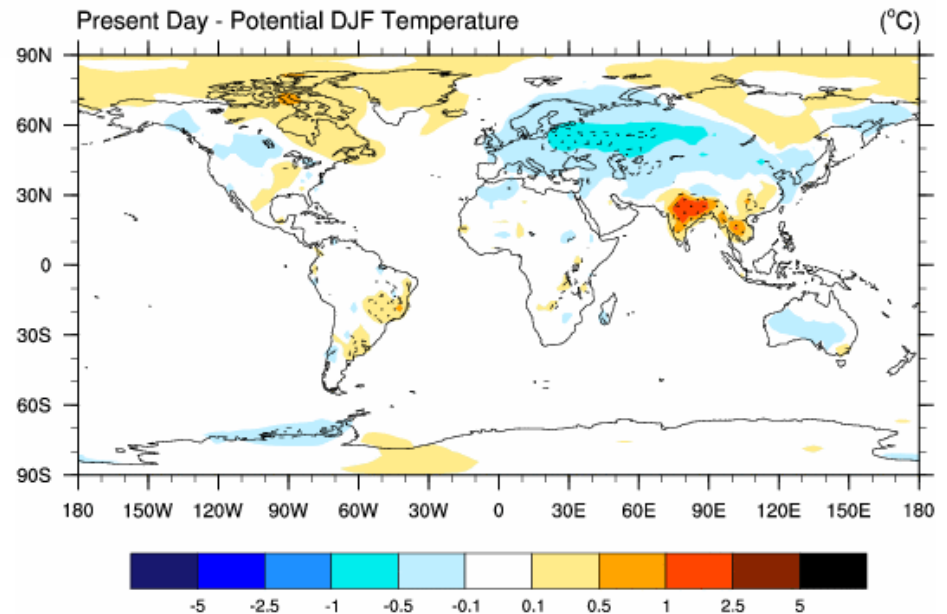
Its not just albedo. The hydrologic cycle affects the land use forcing.

Vegetation-snow albedo feedback



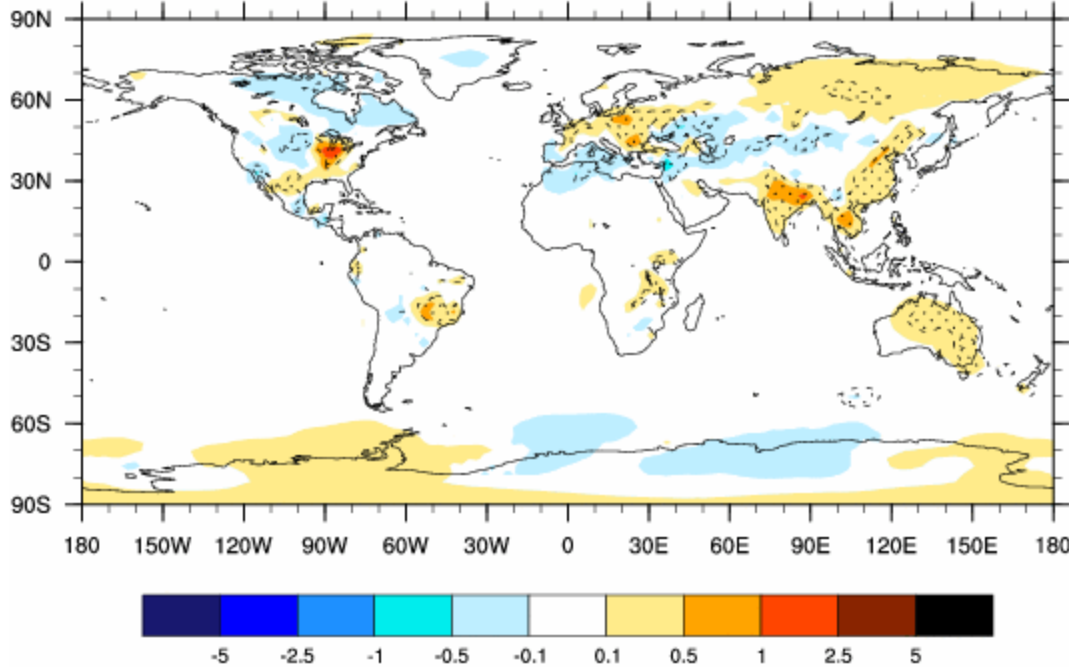
CAM3/CLM3.5 ensemble average

Increased surface albedo leads to surface cooling



Mid-latitude summer

Present Day - Potential JJA Temperature (°C)

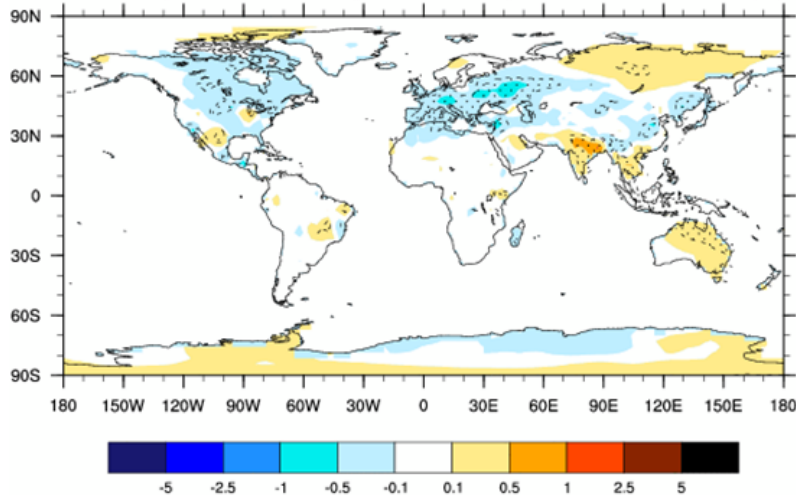


CAM3/CLM3.5 ensemble average

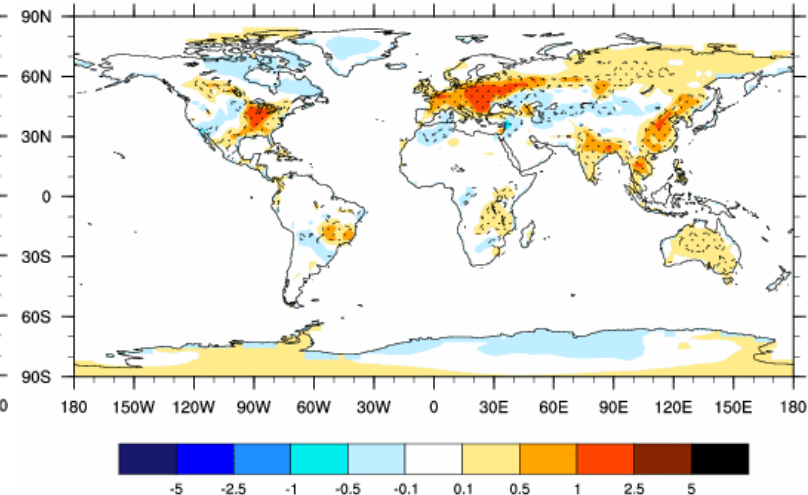
Decrease in daily maximum temperature is offset by increase in daily minimum temperature

Importance of feedbacks with clouds and precipitation

Present Day - Potential JJA Daily Maximum Temperature (°C)



Present Day - Potential JJA Daily Minimum Temperature (°C)



Biofuels - research needs

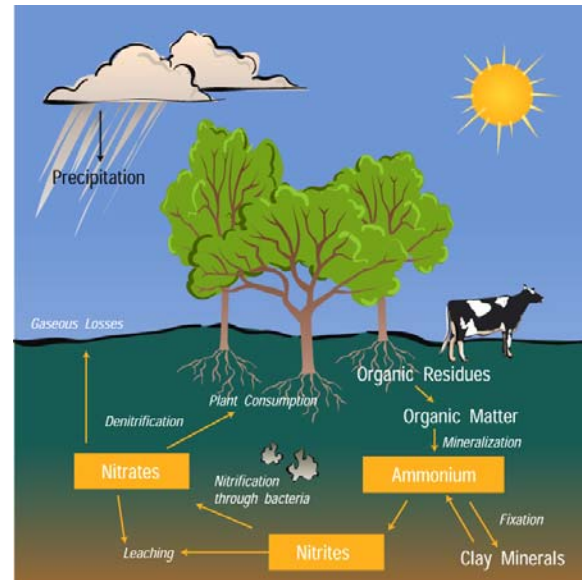


Carbon cycle
Crop management
Land cover change



Maize/Poplar/Switchgrass

Nitrogen cycle
Beth Holland, NCAR

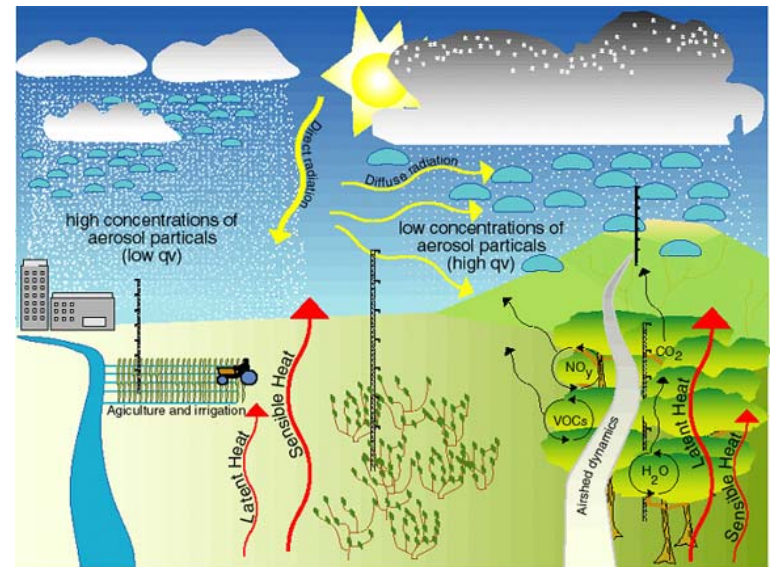


Aerosols and biogenic emissions

Water systems



Center pivot irrigation, northern California



BEACCHON project, NCAR

How to integrate ecological studies with earth system models?

Environmental Monitoring



Eddy covariance flux tower
(courtesy Dennis Baldocchi)

Experimental Manipulation



Soil warming, Harvard Forest

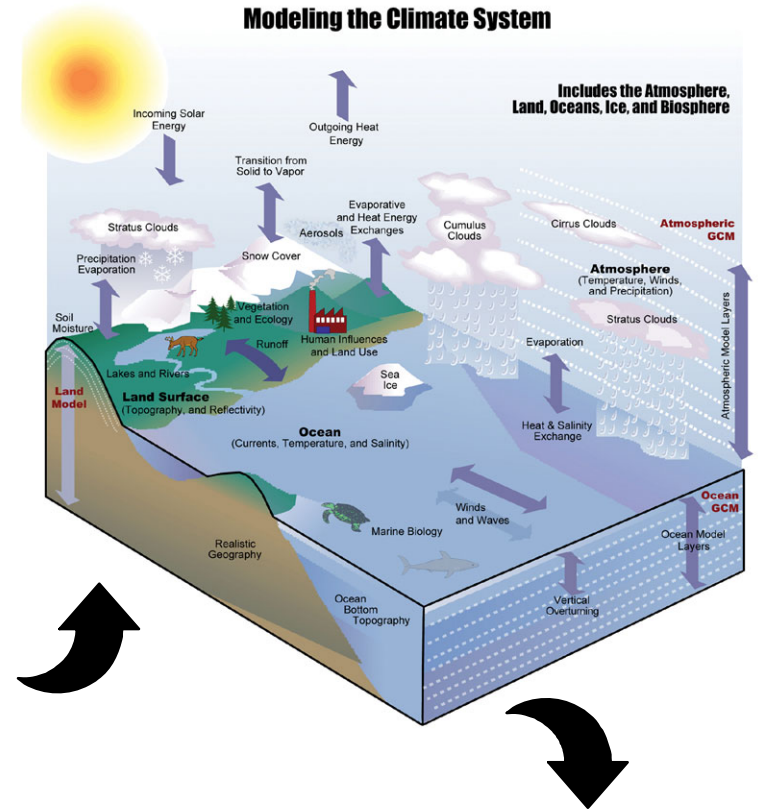


CO₂ enrichment, Duke Forest



Hubbard Brook
Ecosystem Study

Modeling the Climate System



Planetary energetics
Planetary ecology
Planetary metabolism



C-LAMP: Annual net primary production

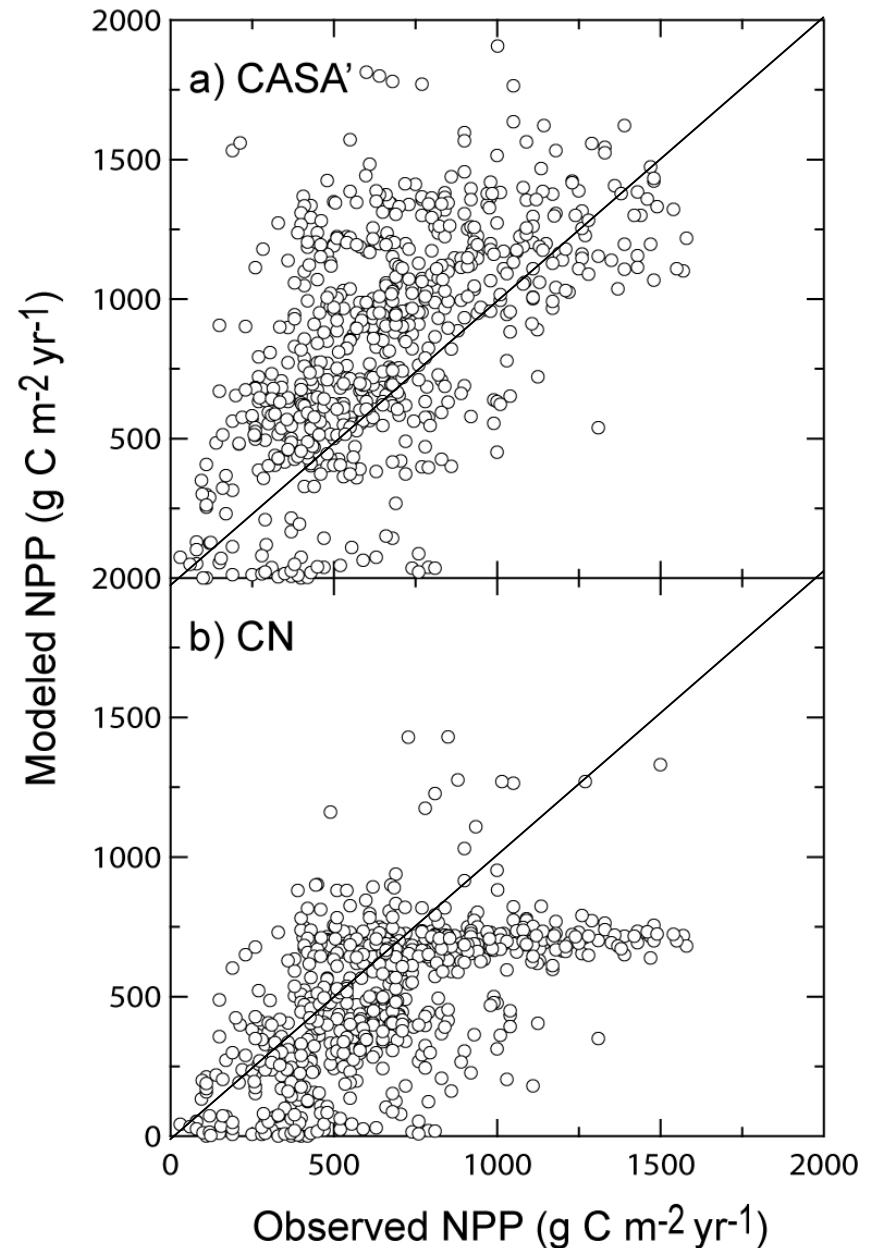
CLAMP is a project of the Community Climate System Model (CCSM) biogeochemistry working group

Ecosystem Model-Data Intercomparison (EMDI) compilation of observations

Class A and Class B observations used

NPP extracted for each model grid cell corresponding to a measurement location

Comparisons are for CLM3 coupled to two biogeochemical carbon models



Comparison with FACE experiments

Experiment	Latitude (°N)	CO ₂ initial	CO ₂ final	Initial NPP	CN		Beta	CASA'		Beta
					final NPP	Initial NPP		final NPP		
DukeFACE	35.6	283.2	364.1	661	733	0.43	1091	1241	0.55	
AspenFACE	45.4	283.2	364.1	358	397	0.43	524	595	0.54	
ORNL-FACE	35.5	283.2	364.1	828	901	0.35	1090	1248	0.58	
POP-EUROFACE	42.2	283.2	364.1	235	253	0.30	397	453	0.56	
Mean:						0.38			0.56	

Observed mean β : **0.60**

CN model mean β : **0.38**

CASA' model mean β : **0.56**

Observed NPP increase (376 → 550ppm): 23%

CN predicted (376 → 550ppm): 14%

CASA' predicted (376 → 550ppm): 21%

$$\beta = \frac{\left(\frac{NPP(f)}{NPP(i)} - 1 \right)}{\ln\left(\frac{CO_2(f)}{CO_2(i)} \right)}$$

$$NPP(t) = NPP(i) \cdot \left[\beta \cdot \ln\left(\frac{CO_2(t)}{CO_2(i)} \right) + 1 \right]$$

Summary

The land surface is a critical interface through which humans are impacted and can affect climate change

Land ESM development:

- Expand capability to simulate forcings and feedbacks in earth system
- Increased emphasis on suitability for impacts, adaptation, and mitigation research
- Requires an integrated systems approach:
 - Biogeochemical systems
 - Water systems
 - Ecosystems
 - Human systems
 - Glacier systems

(IPCC 2007)

